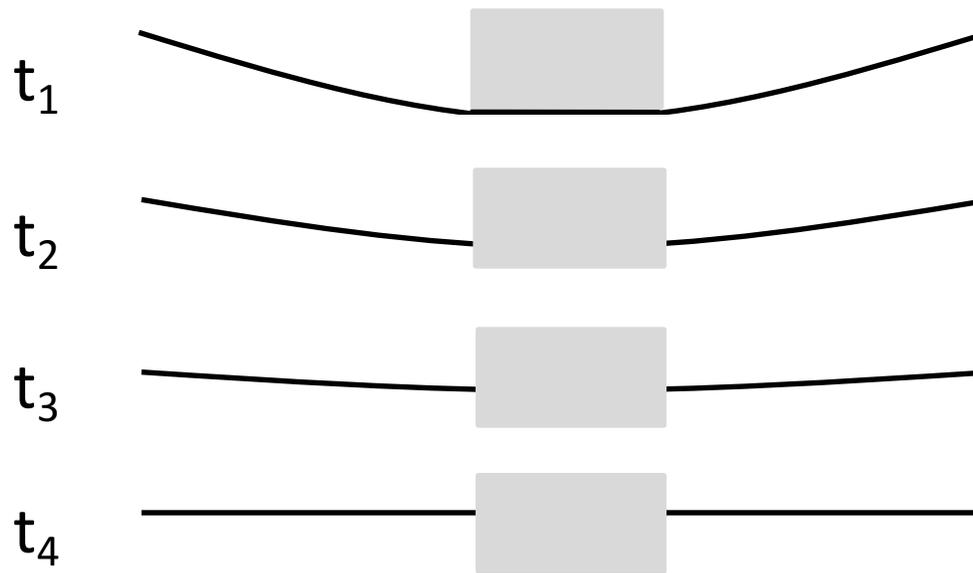


EPSC510 Module 2

Lecture 2: Sea Level Change

James Bay, Ontario, photo credit: Natalya Gomez 2008

Last Class: Isostasy and Rheology



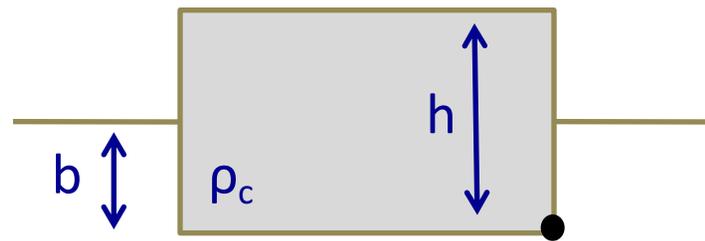
Rheology
(The macroscopic
response of a material to
an applied stress)



Isostasy

Isostasy:

Archimedes Principle /
Hydrostatic Equilibrium
applied to the Earth.



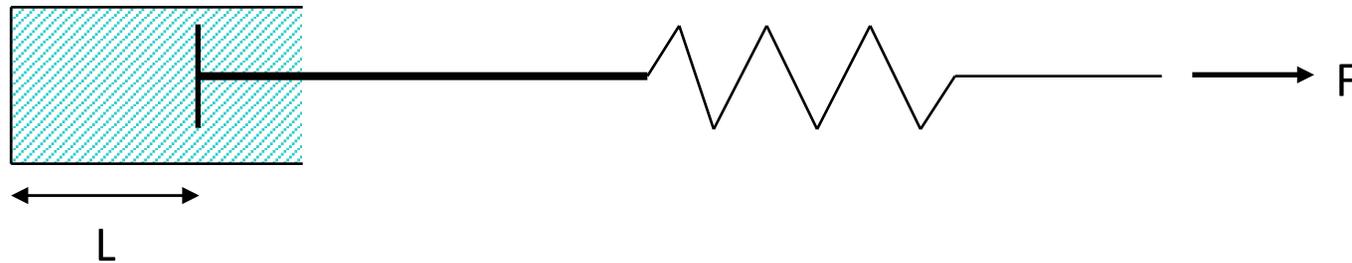
$$\frac{b}{h} = \frac{\rho_c}{\rho_m}$$

Last Class: Isostasy and Rheology

1-D MODELS OF RHEOLOGY

3) Viscoelastic Body

- many ways to define such a material
- assume the following 1-D linear response: “Maxwell Body”



- constitutive equation:

$$\dot{\epsilon} = \frac{1}{k} \dot{F} + \frac{1}{\eta} F$$

End members: $k \rightarrow \infty$ Viscous body
 $\eta \rightarrow \infty$ Elastic body

Last Class: Isostasy and Rheology

ENVIRONMENT DEPENDENCIES

- Both k & η are functions of pressure P & temperature T
- ν is a much stronger function of P, T (in general)
- to a first approximation: $\eta = \eta(P, T) = \eta(T)$

$$\dot{\epsilon} = \frac{1}{k} \dot{F} + \frac{1}{\eta} F$$

- experiments have shown: $\eta(T) = \eta_0 \exp(-a[T - T_0])$
 $\eta(T_0) = \eta_0$
- “ a ” is a complex function of thermodynamic properties
- thus, small changes in temperature produce exponentially large changes in viscosity!!!

$$\dot{\epsilon} = \frac{1}{k} \dot{F} + \frac{1}{\eta_0 \exp(-a[T - T_0])} F$$

Cases:

(1) Cold ... means η large so: $\dot{\epsilon} \cong \frac{1}{k} \dot{F}$ Elastic response

(2) Hot ... means η small so: $\dot{\epsilon} \cong \frac{1}{\eta} F$ Viscous response

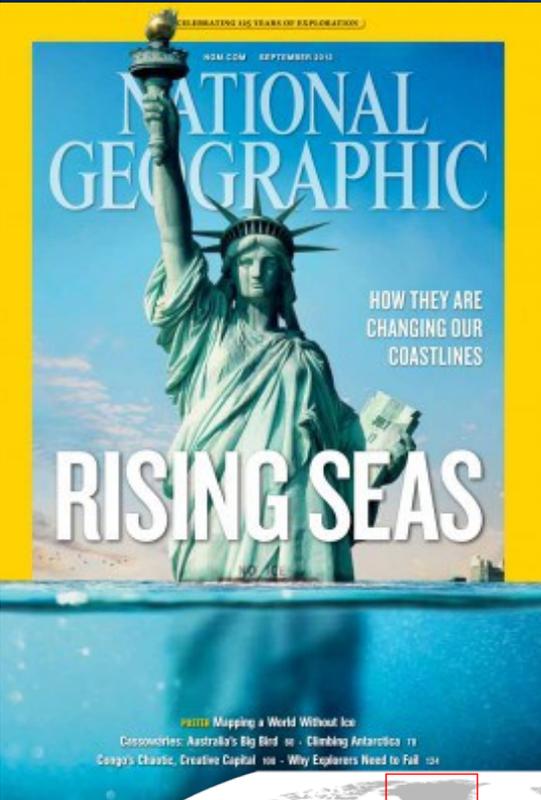
(3) Warm ... intermediate T so: viscoelastic response ... similar to tar

Outline

This Class: Sea Level Change

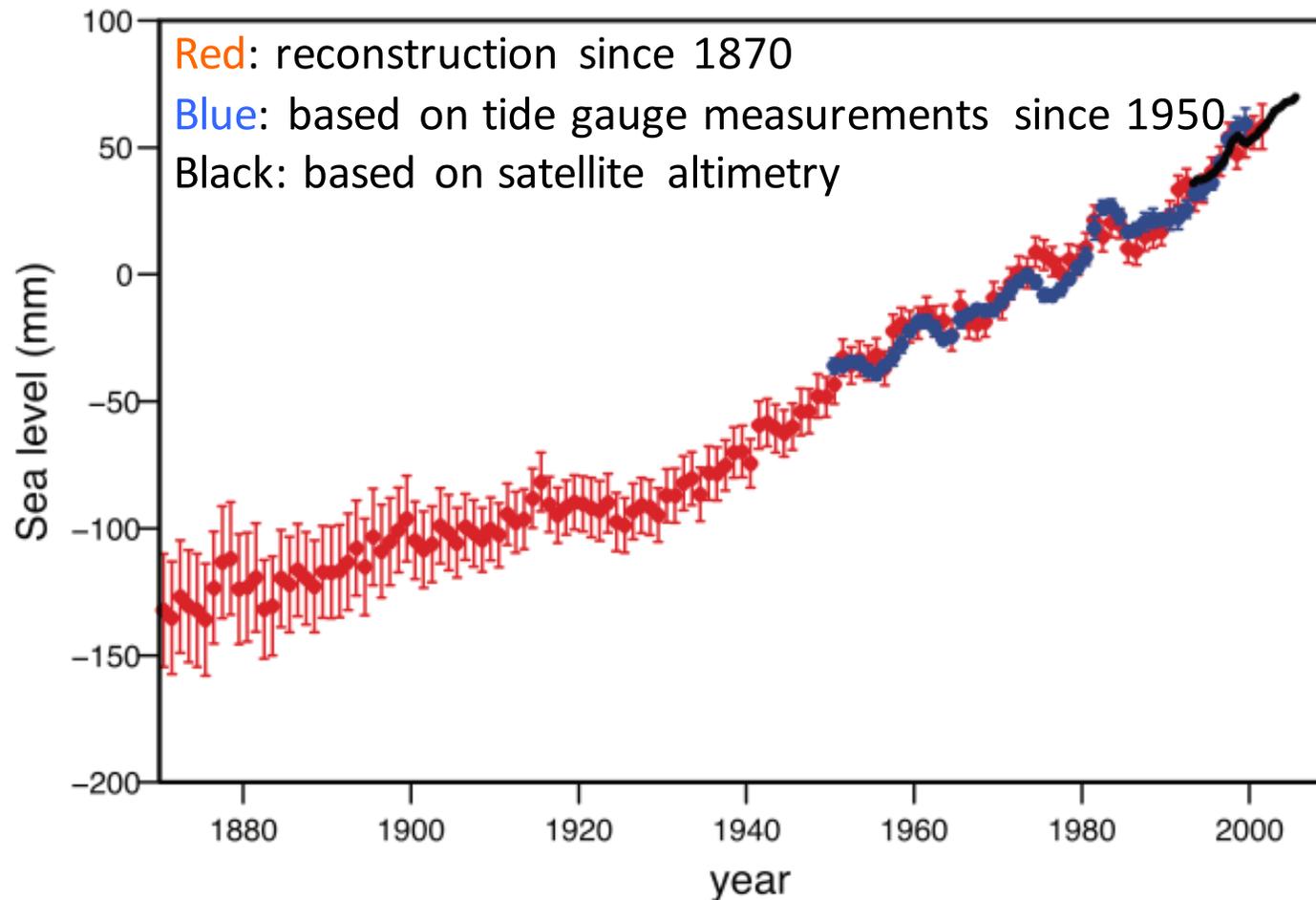
1. Sea-level records and definitions
2. Sea-level changes in response to ice cover variations
 - a. The sea-level equation
 - b. Physics
3. Example

Future Sea-Level Rise?



1. Sea – level records

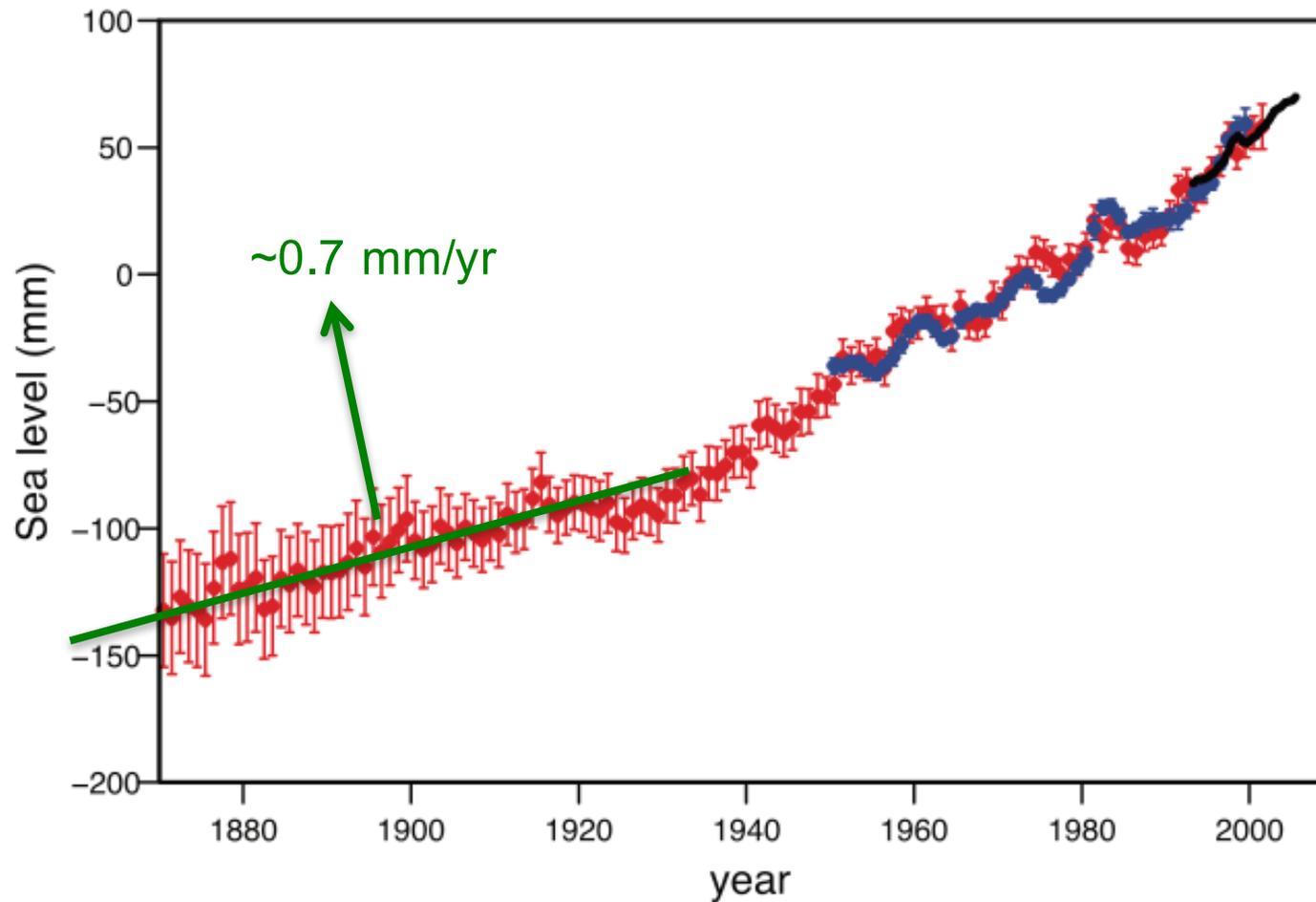
Annual averages of global sea level



Bindoff et al., (2007). Observations: Climate Change and Sea Level. In: *Climate Change 2007: The Physical Science Basis. Contribution of WG1 to the 4th Assessment Report of the IPCC*. USA

1. Sea – level records

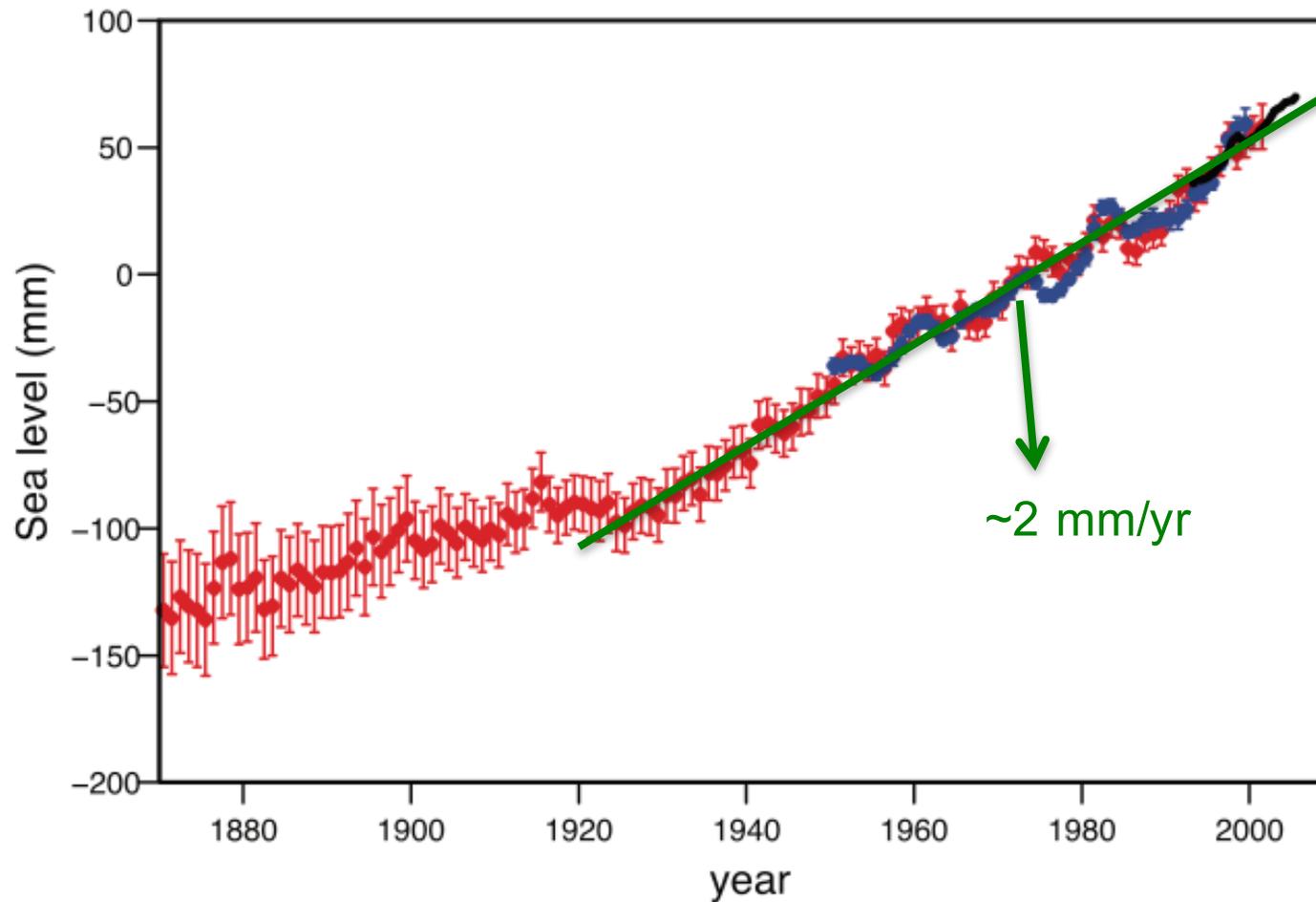
Annual averages of global sea level



Bindoff et al., (2007). Observations: Climate Change and Sea Level. In: *Climate Change 2007: The Physical Science Basis. Contribution of WG1 to the 4th Assessment Report of the IPCC*. USA

1. Sea – level records

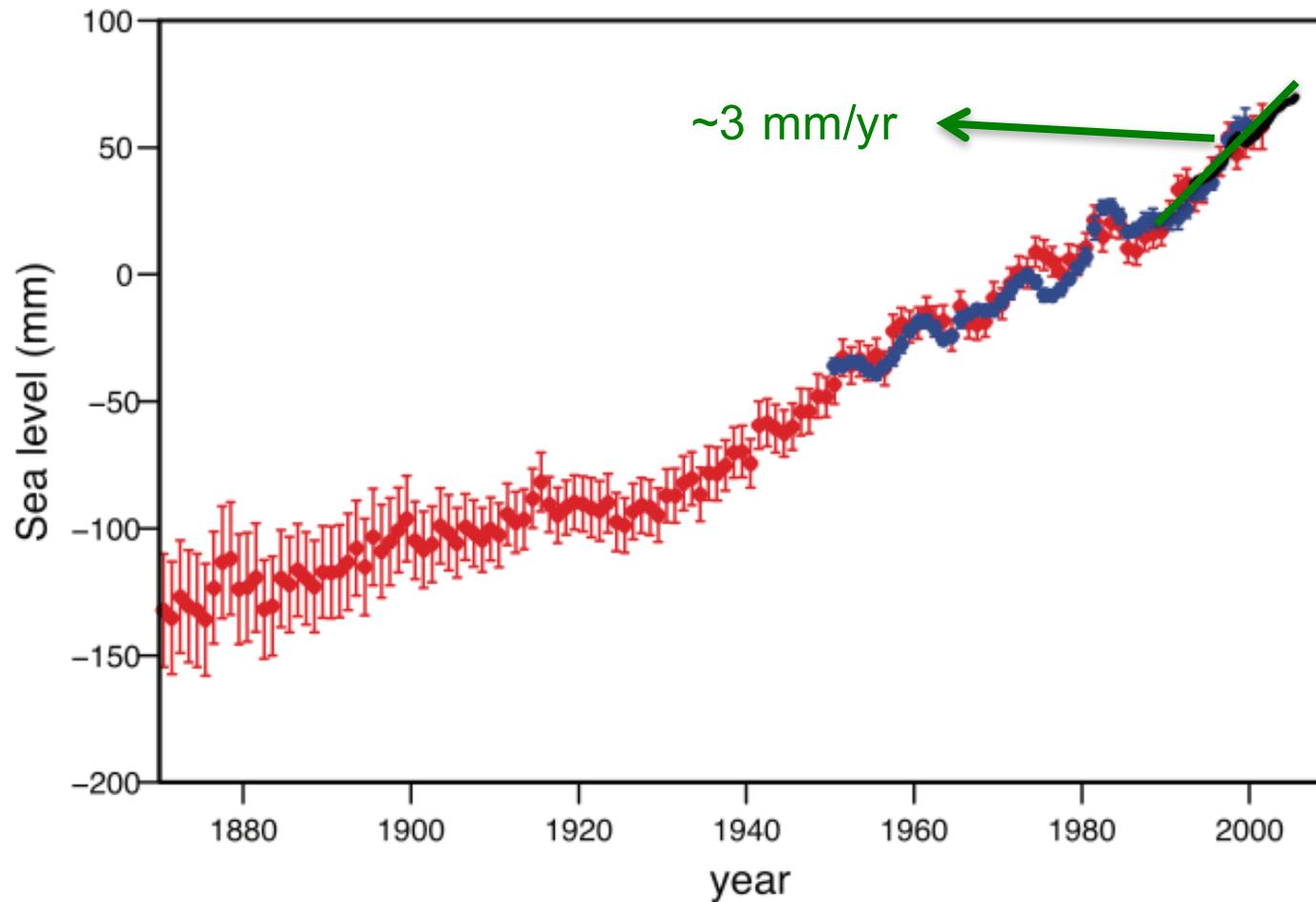
Annual averages of global sea level



Bindoff et al., (2007). Observations: Climate Change and Sea Level. In: *Climate Change 2007: The Physical Science Basis. Contribution of WG1 to the 4th Assessment Report of the IPCC*. USA

1. Sea – level records

Annual averages of global sea level

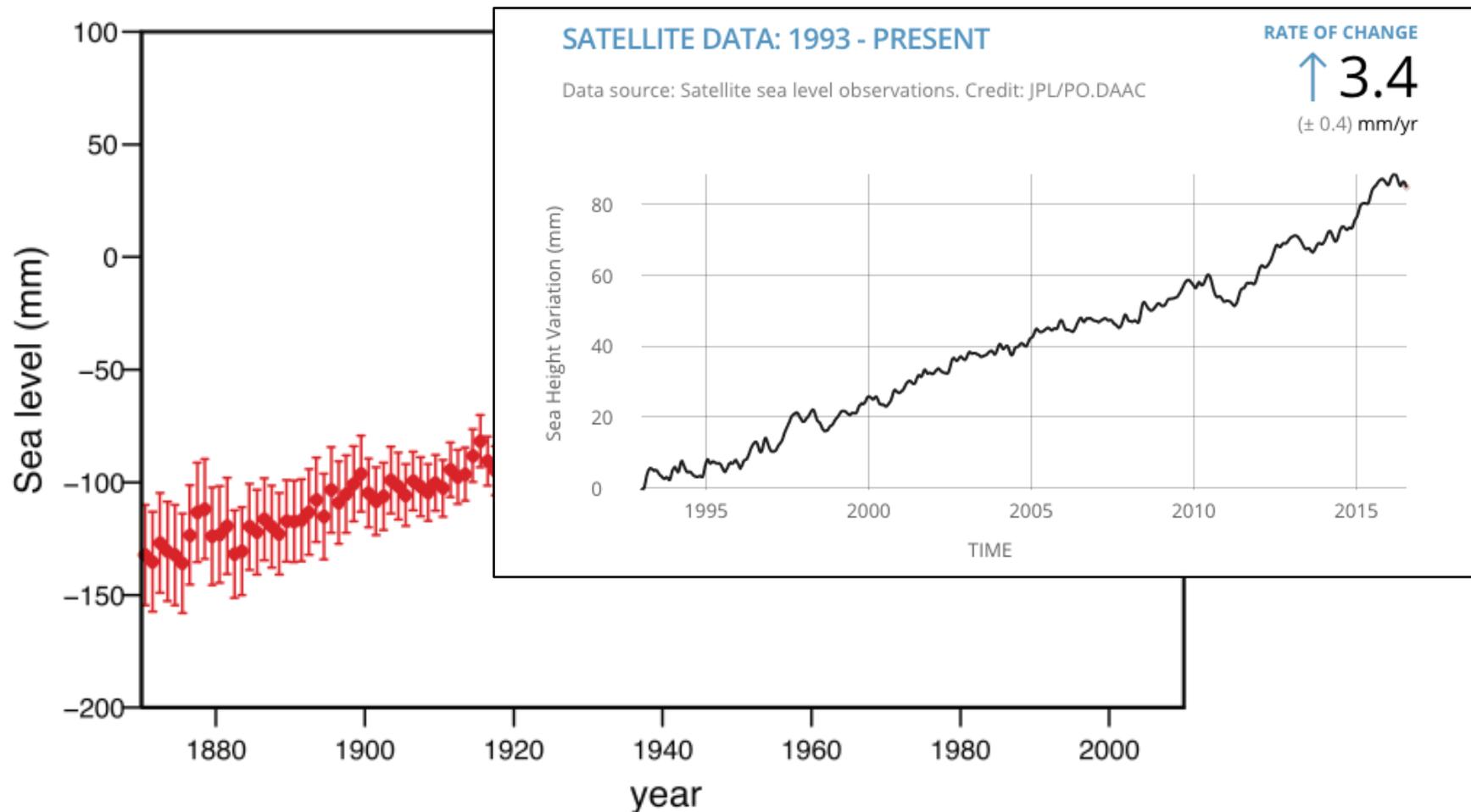


Bindoff et al., (2007). Observations: Climate Change and Sea Level. In: *Climate Change 2007: The Physical Science Basis. Contribution of WG1 to the 4th Assessment Report of the IPCC*. USA

1. Sea – level records

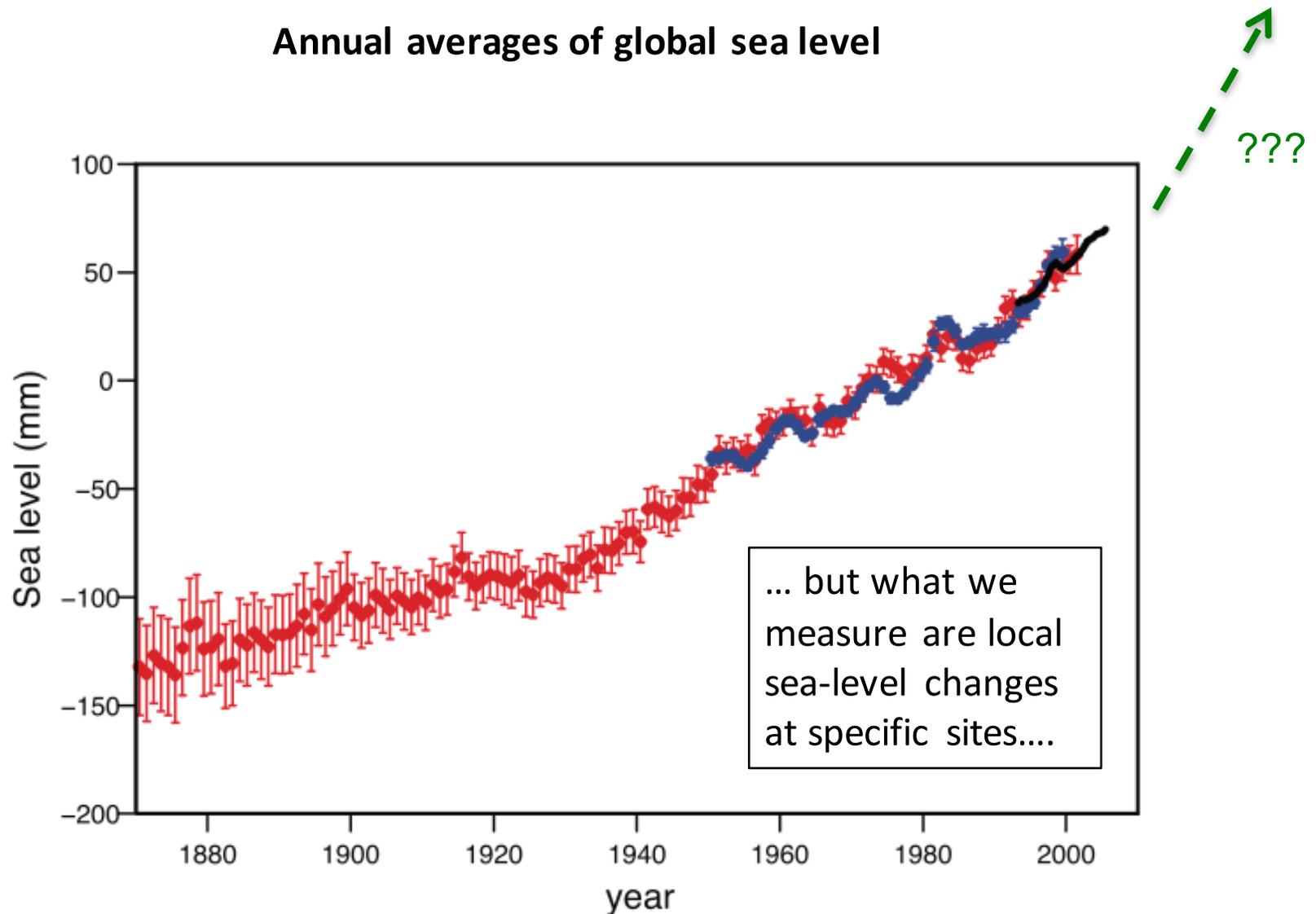
Annual averages of global sea level

<https://sealevel.nasa.gov/understanding-sea-level/key-indicators/global-mean-sea-level>



Bindoff et al., (2007). Observations: Climate Change and Sea Level. In: *Climate Change 2007: The Physical Science Basis. Contribution of WG1 to the 4th Assessment Report of the IPCC*. USA

1. Sea – level records



Bindoff et al., (2007). Observations: Climate Change and Sea Level. In: *Climate Change 2007: The Physical Science Basis. Contribution of WG1 to the 4th Assessment Report of the IPCC*. USA

1. Sea – level records

Tamisiea et al. (2014):
Sea Level Records and Definitions

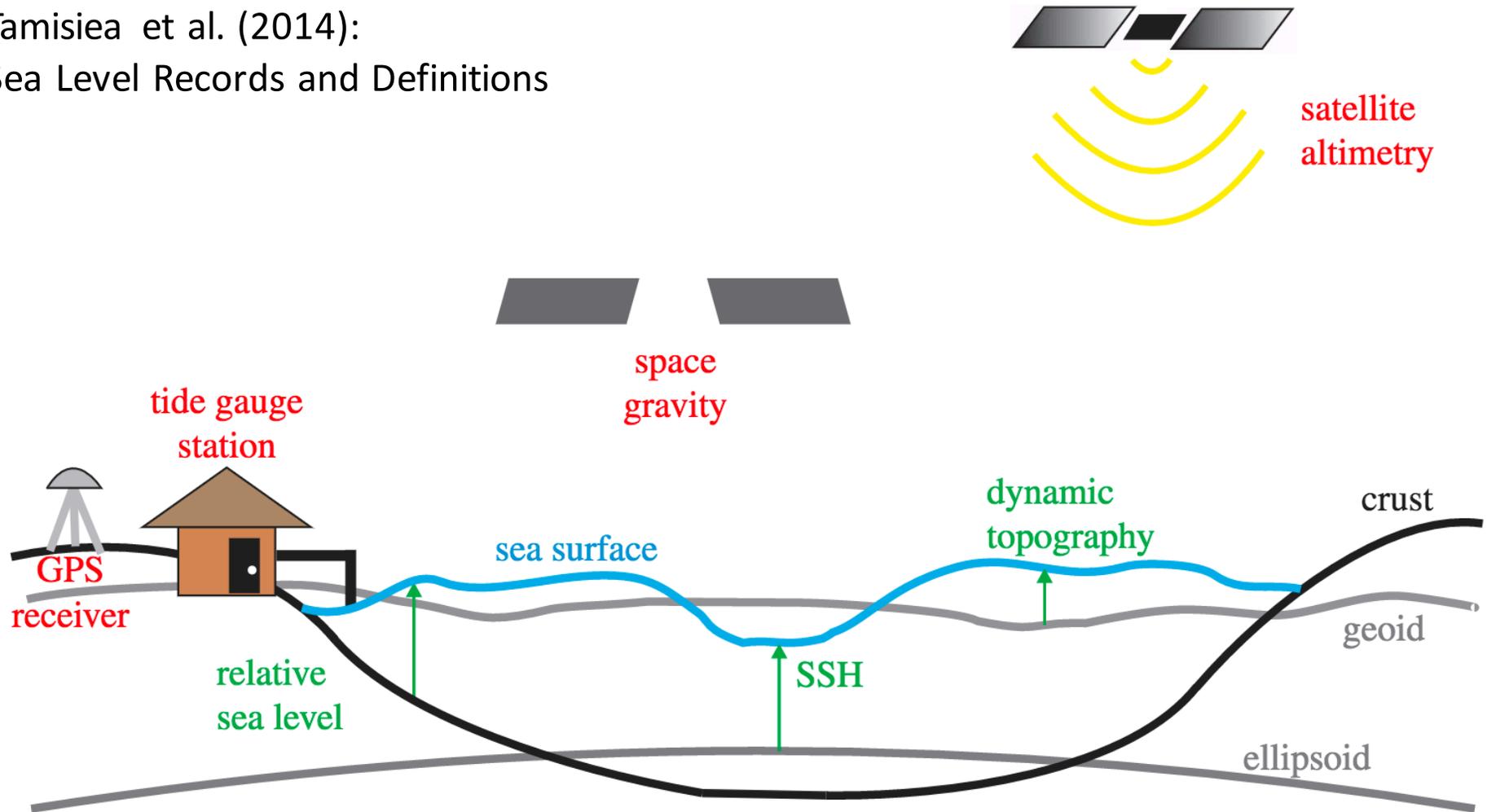


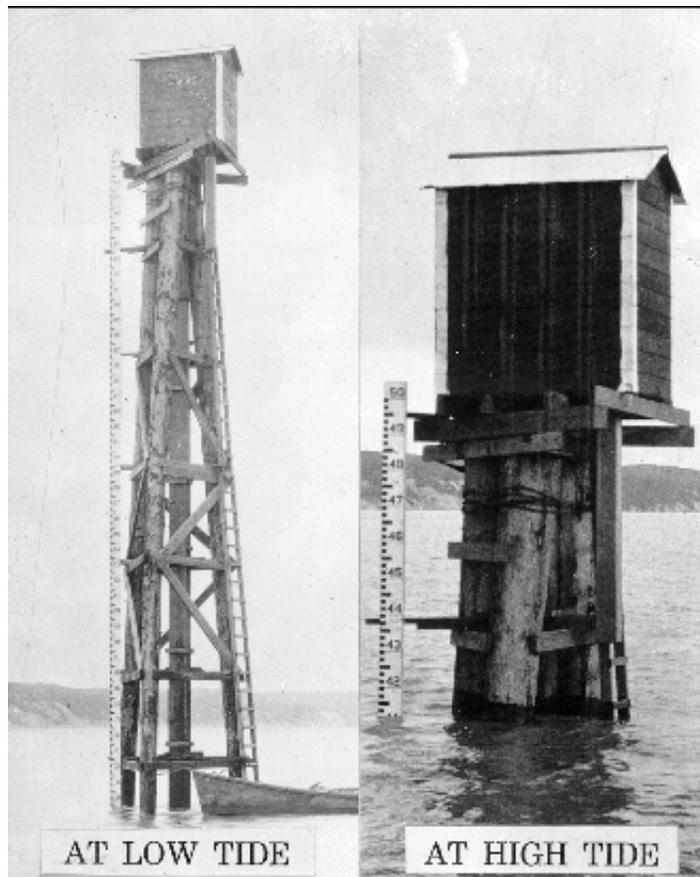
Figure 1. Simple schematic illustrating the relationship between sea surface height (SSH), the geoid, and dynamic topography. Included on the figure are representations of different components of the observing system and their respective measurement: GPS (or GNSS) for crustal deformation, satellite gravity for the geoid, altimetry for SSH and tide gauges for relative sea level.

1. Sea – level records

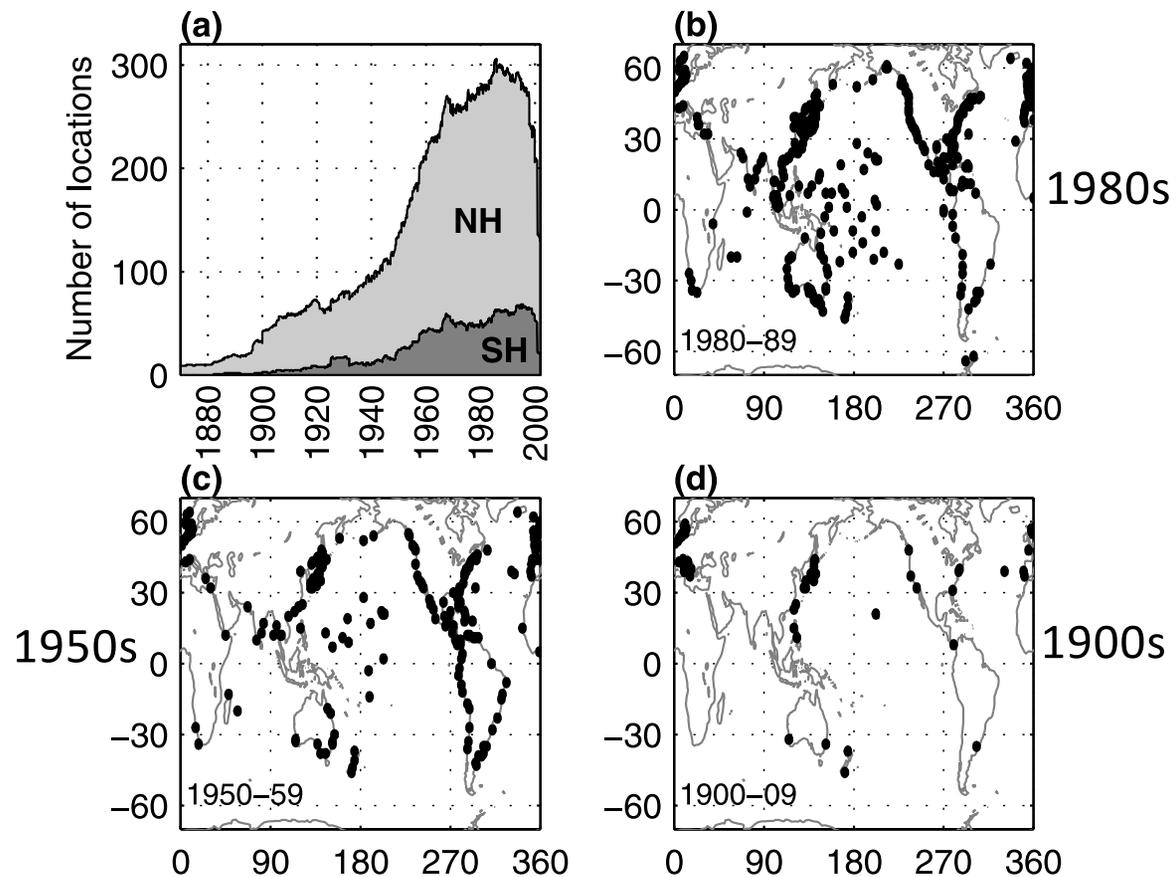
Tide Gauges measure:

(Relative) Sea Level = ocean surface – ocean floor

*caution: “relative sea level” can also refer to the sea level at some time in the past relative to the present.



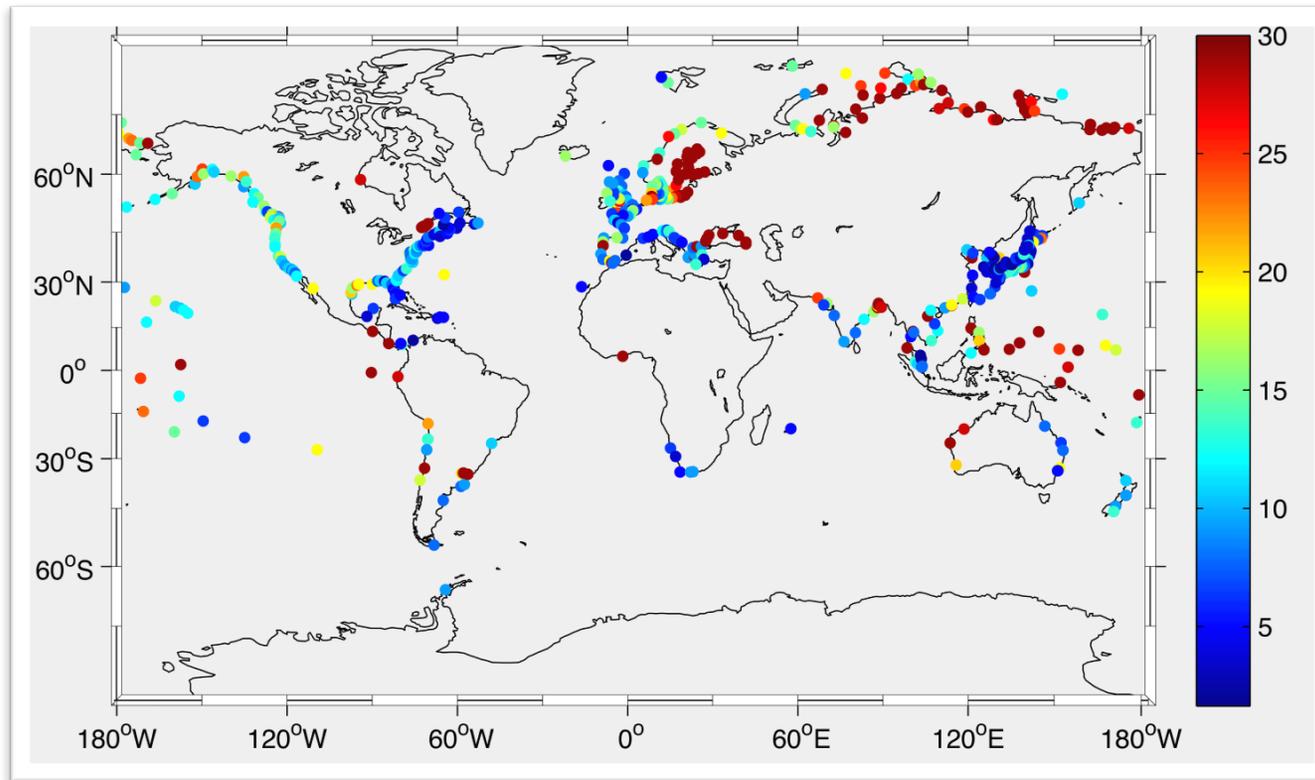
<http://co-ops.nos.noaa.gov/about2.html>



Church and White (2006)

1. Sea – level records

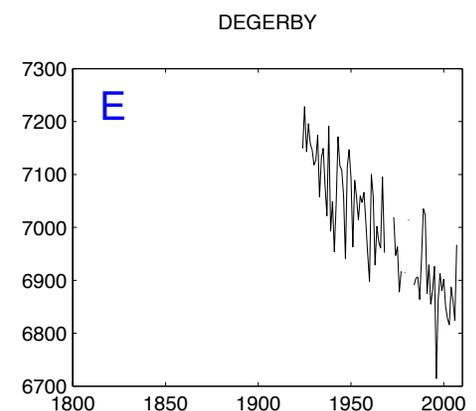
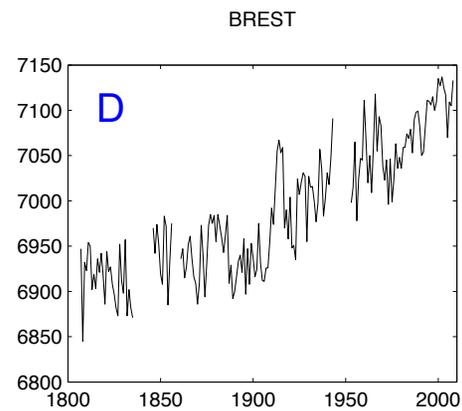
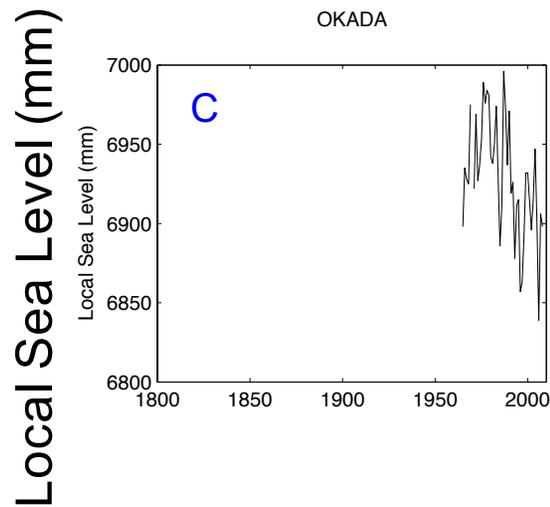
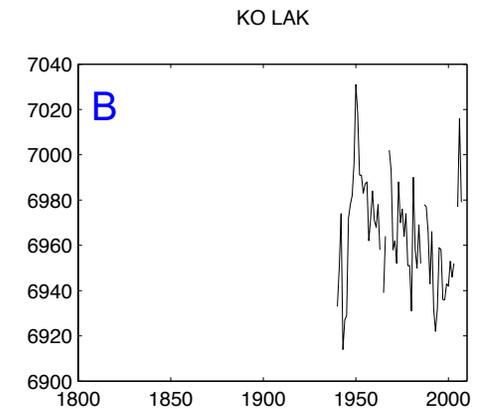
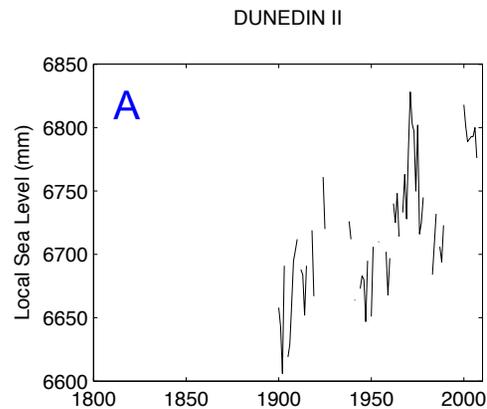
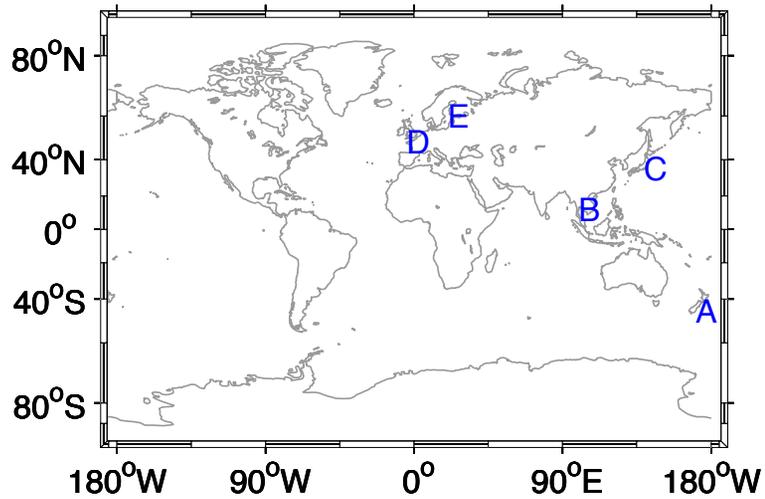
Variance (cm^2) in Annual Tide Gauge Records



1157 tide gauges in the Permanent Service for Mean Sea Level (PSMSL)
Revised Local Reference (RLR) database

1. Sea – level records

You can get these yourself from:
<http://www.psmsl.org/>



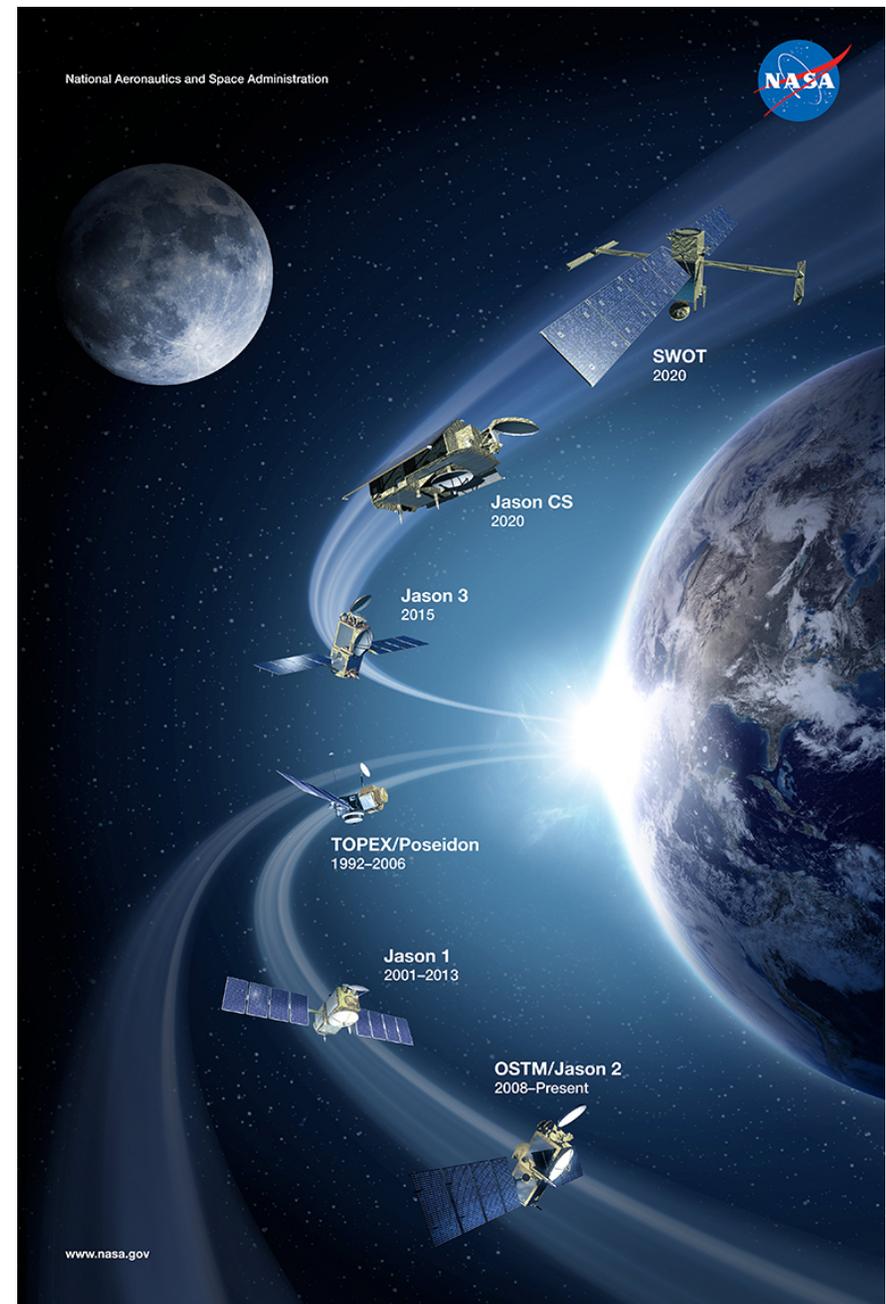
Calendar Year

1. Sea – level records

Satellite Altimetry

“Since 1992 NASA, NOAA and European partners have been tracking global ocean surface topography with joint ocean altimeter satellite missions from an orbit 1,336 km above the ocean surface. The spacecrafts' radar altimeters measure the precise distance between the satellite and sea surface. This record began with TOPEX/Poseidon, followed by Jason-1 and the Ocean Surface Topography Mission on Jason-2, and will be continued by Jason-3.” - NASA

<http://www.eumetsat.int/jason/print.htm>

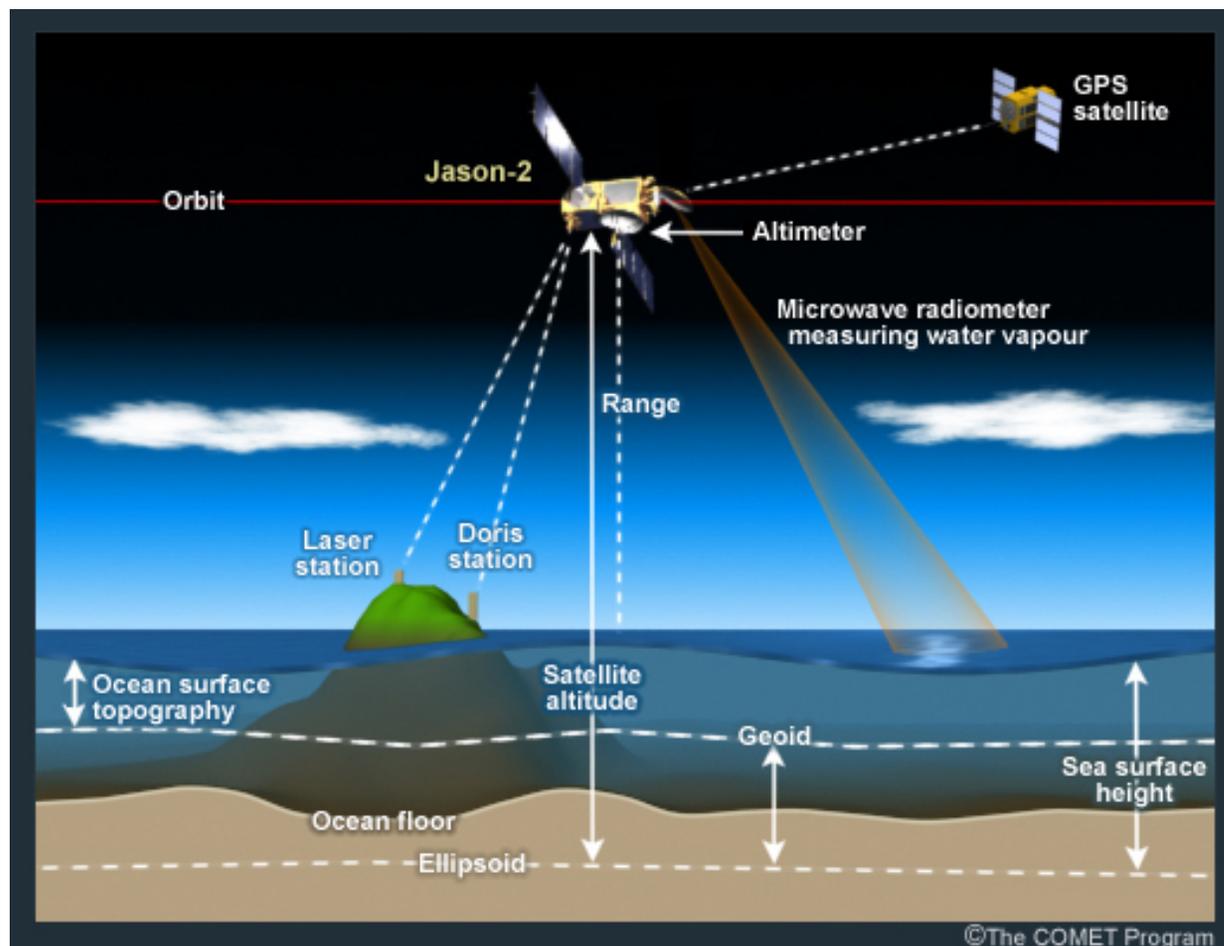


1. Sea – level records

Satellite Altimetry measures:

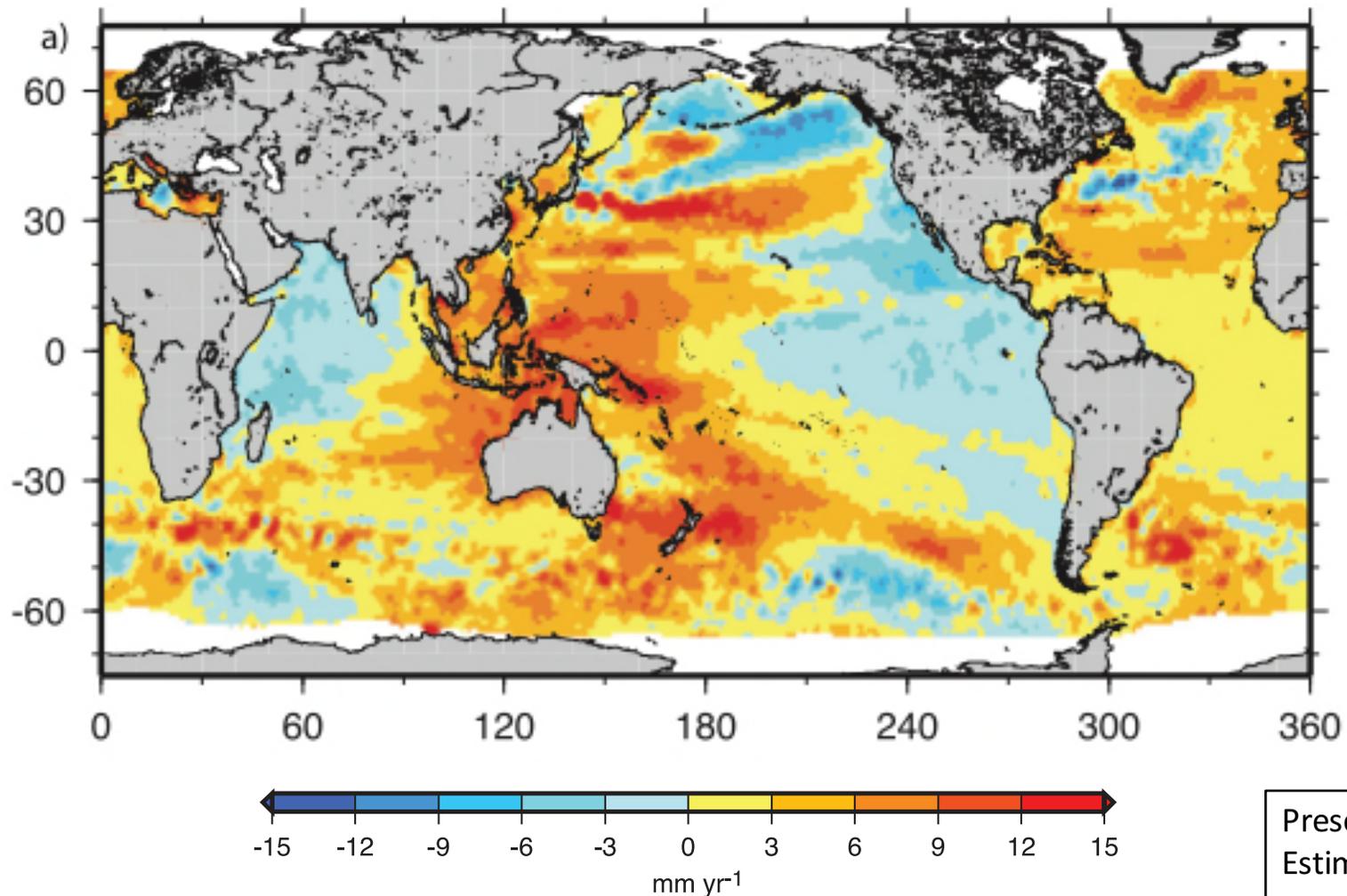
Sea Surface Height = ocean surface – reference ellipsoid

*caution: “sea level” and “sea surface height” are used interchangeably in some literature. In this talk, they are different.



1. Sea – level records

Sea Level trend (mm/year) from 1993-2003 from satellite altimetry



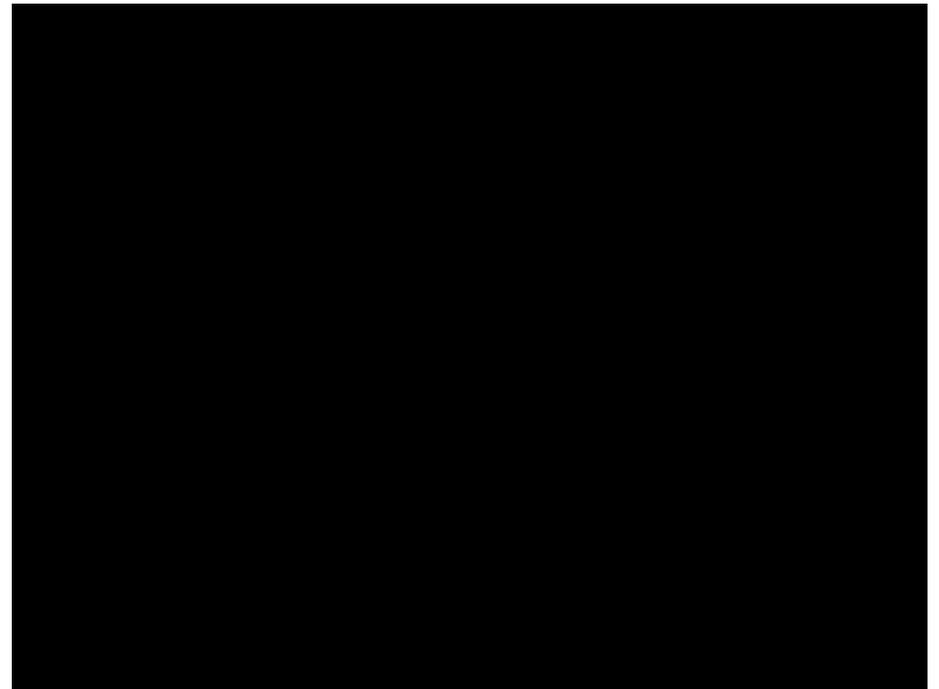
Bindoff et al., (2007). Observations: Climate Change and Sea Level. In: *Climate Change 2007: The Physical Science Basis. Contribution of WG1 to the 4th Assessment Report of the IPCC*. USA

Presentation topic:
Estimating
contributions to
this signal

1. Sea – level records

GRACE – Gravity Recovery and Climate Experiment (since 2002)

2 satellites orbit the Earth together, and their speed changes according to the mass below them. The distance between satellites is measured to produce global maps of mass variations.



1. Sea – level records

GRACE – Gravity Recovery and Climate Experiment (since 2002)

2 satellites orbit the Earth together, and their speed changes according to the mass below them. The distance between satellites is measured to produce global maps of mass variations.

Challenges:

Low Resolution

- Combine with other datasets, develop data processing techniques (e.g. Slepian functions - <https://eos.org/project-updates/a-suite-of-software-analyzes-data-on-the-sphere-2>)

Separating sources - Observed changes mass distribution across the Earth's surface are associated with a combination of:

- Solid Earth changes
- Ocean circulation
- Groundwater storage
- Exchange of mass between ice sheets and mountain glaciers and the ocean
- Atmospheric circulation
- ... other effects, e.g., forest fires!

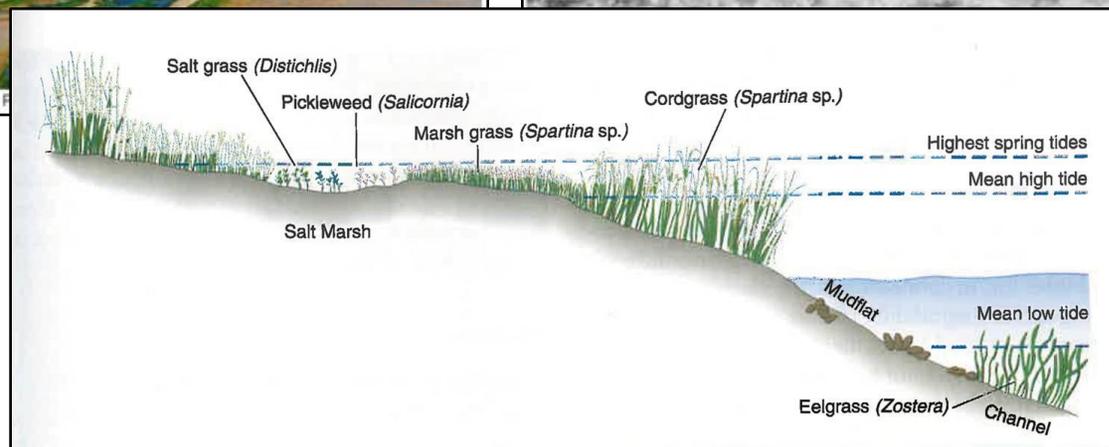
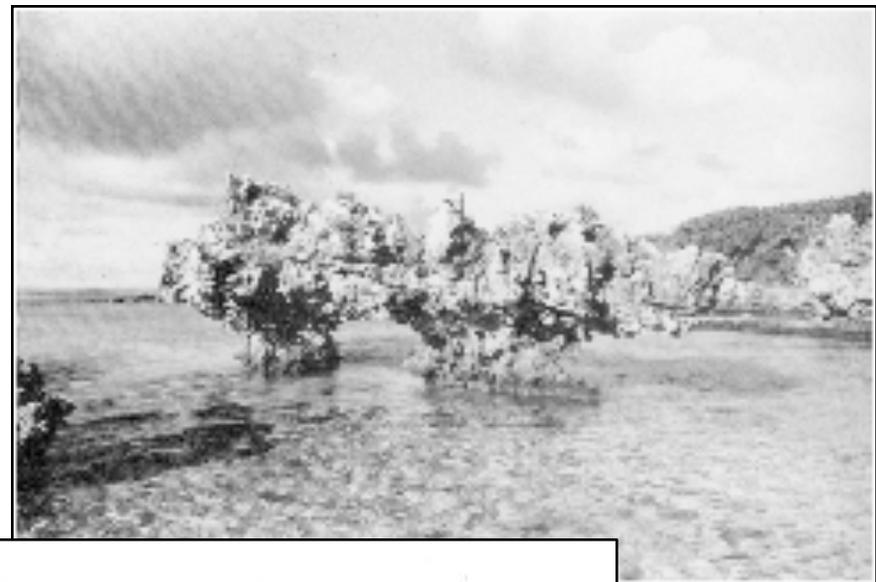


Data, info and related studies available here: <http://www.csr.utexas.edu/grace/>

1. Sea – level records

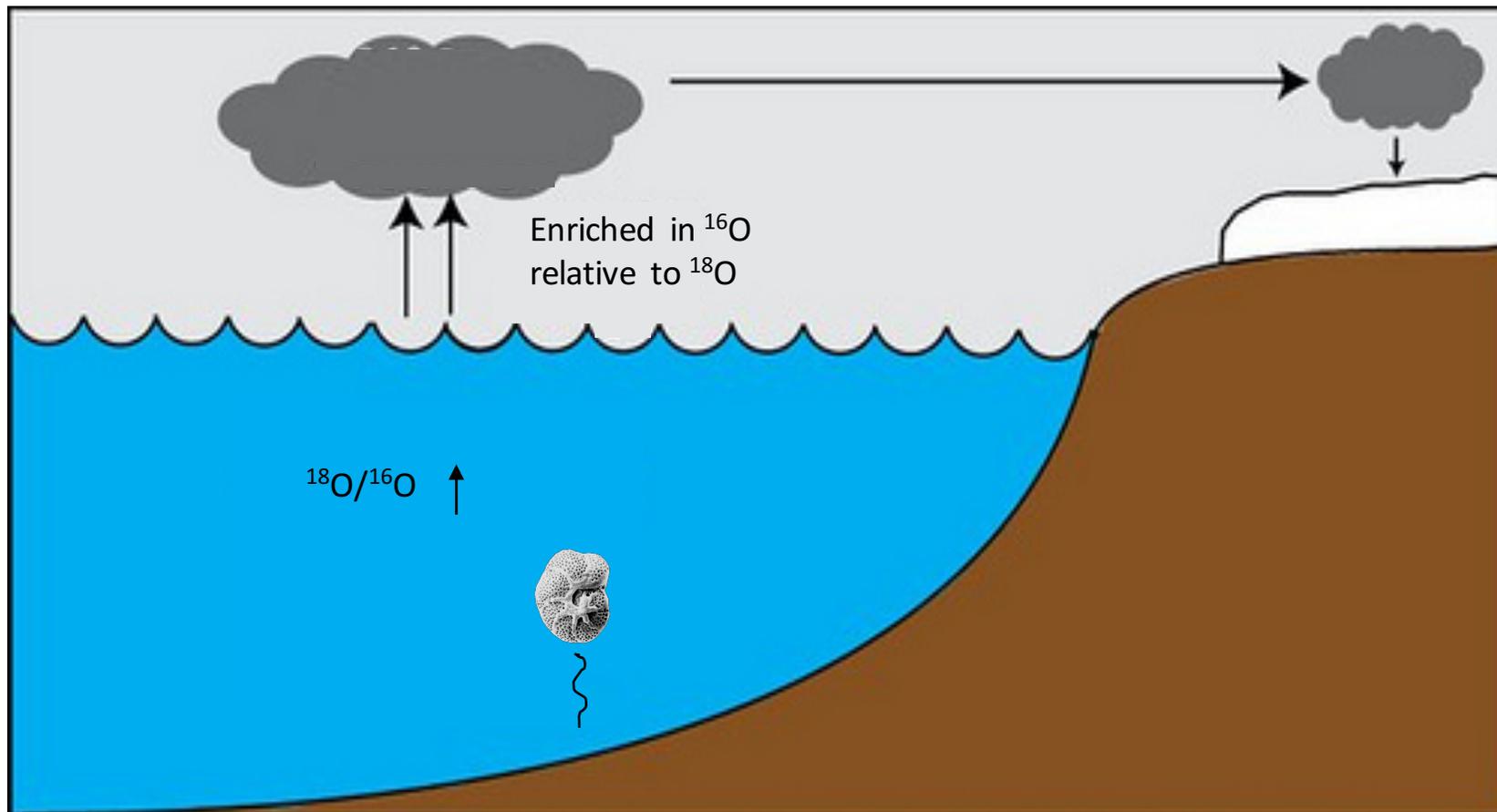
Paleo sea-level records – biological and geological, e.g.,

- Past shoreline markers
- Corals (grow up to just below the surface, then are exposed to sunlight above water when sea level falls.)
- Salt Marshes (last ~ky)



1. Sea – level records

Oxygen Isotope Fractionation (in 1 slide): Ice has a lower $d^{18}\text{O}$ value than seawater, so oxygen isotope values provide a proxy for contribution from ice to sea level changes.



1. Sea – level records

MORE
ICE

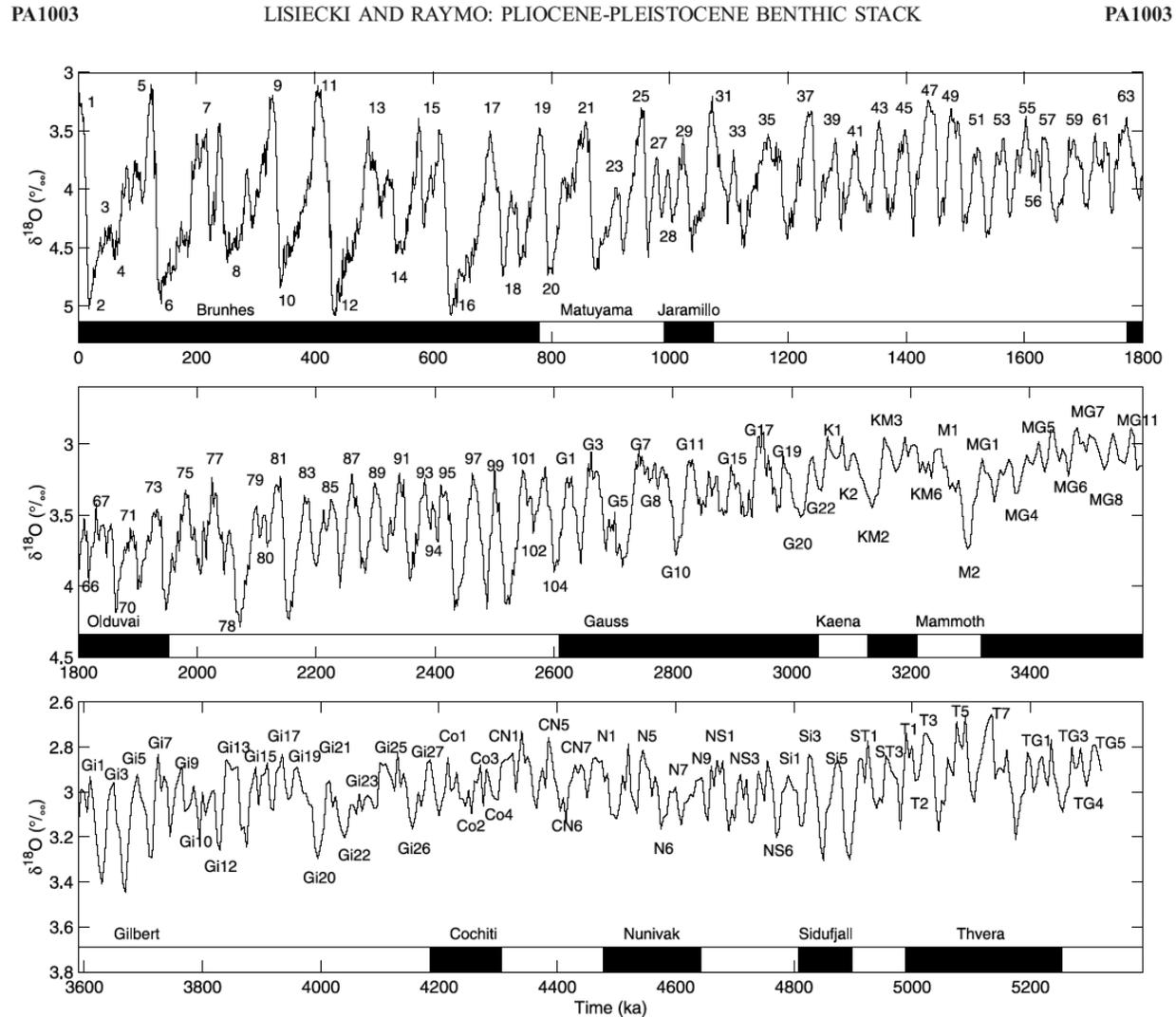
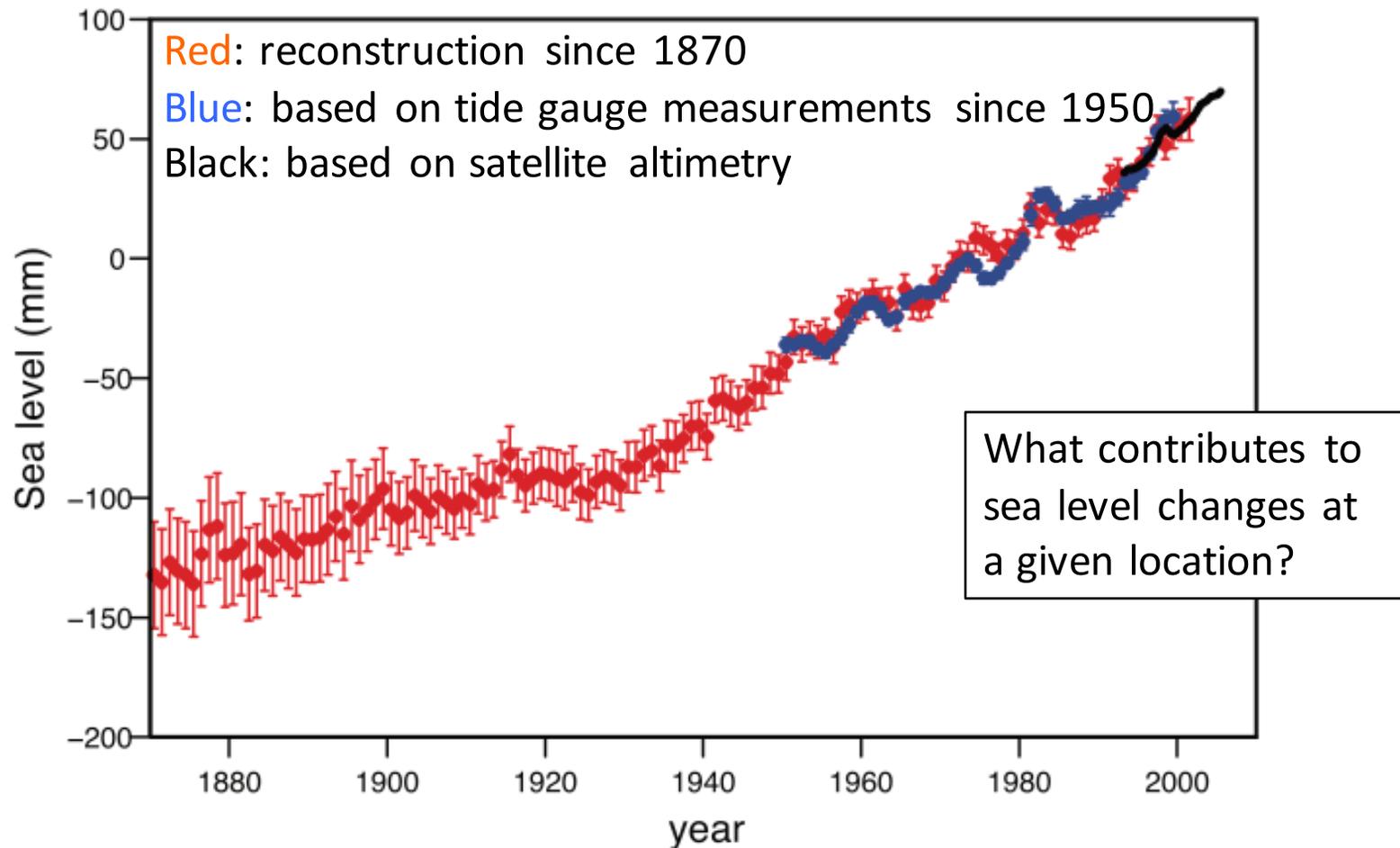


Figure 4. The LR04 benthic $\delta^{18}\text{O}$ stack constructed by the graphic correlation of 57 globally distributed benthic $\delta^{18}\text{O}$ records. The stack is plotted using the LR04 age model described in section 5 and with new MIS labels for the early Pliocene (section 6.2). Note that the scale of the vertical axis changes across panels.

We will come
back to this...

1. Sea – level records

Annual averages of global sea level



Bindoff et al., (2007). Observations: Climate Change and Sea Level. In: *Climate Change 2007: The Physical Science Basis. Contribution of WG1 to the 4th Assessment Report of the IPCC*. USA

2a. The Sea – level Equation

Local sea-level changes are the result of the combination of various effects:

1) Dynamic Effects

long term: thermal expansion of the ocean

short term: ocean dynamics, air-sea interactions, temperature and salinity variations, tides...

2) Static Effects

the advance and retreat of ice sheets and mountain glaciers produce distinct patterns of sea level change called “sea level fingerprints”

3) Glacial Isostatic Adjustment

the ongoing adjustment of the land and sea surface due to ice cover changes that occurred in the past

4) Long term effects (as we discussed last week, we may come back to this)



Sea – level changes associated with ice!

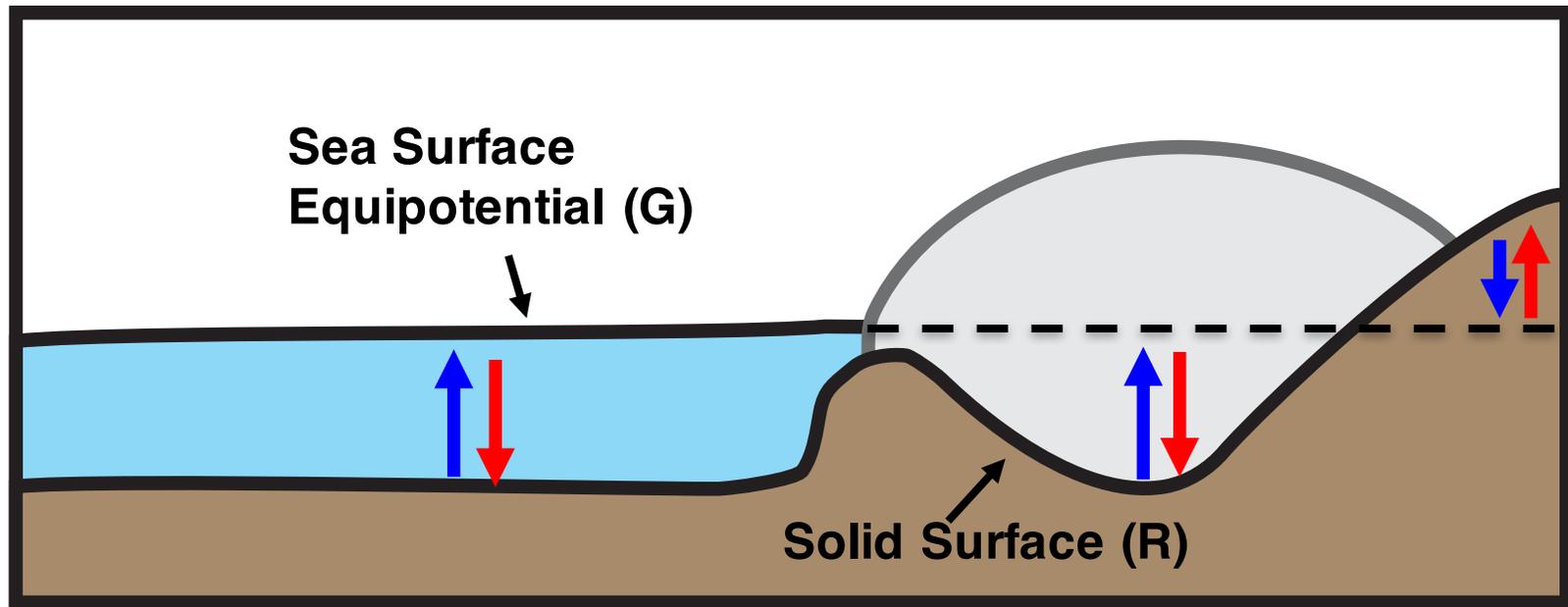
Quick Mental Break!



King Penguins,

Photo credit: Leslie Frost, Environmental Manager at Australia Antarctic Division
(Acquired in an NYC taxi cab)

2a. The sea level equation



 Static Sea Level

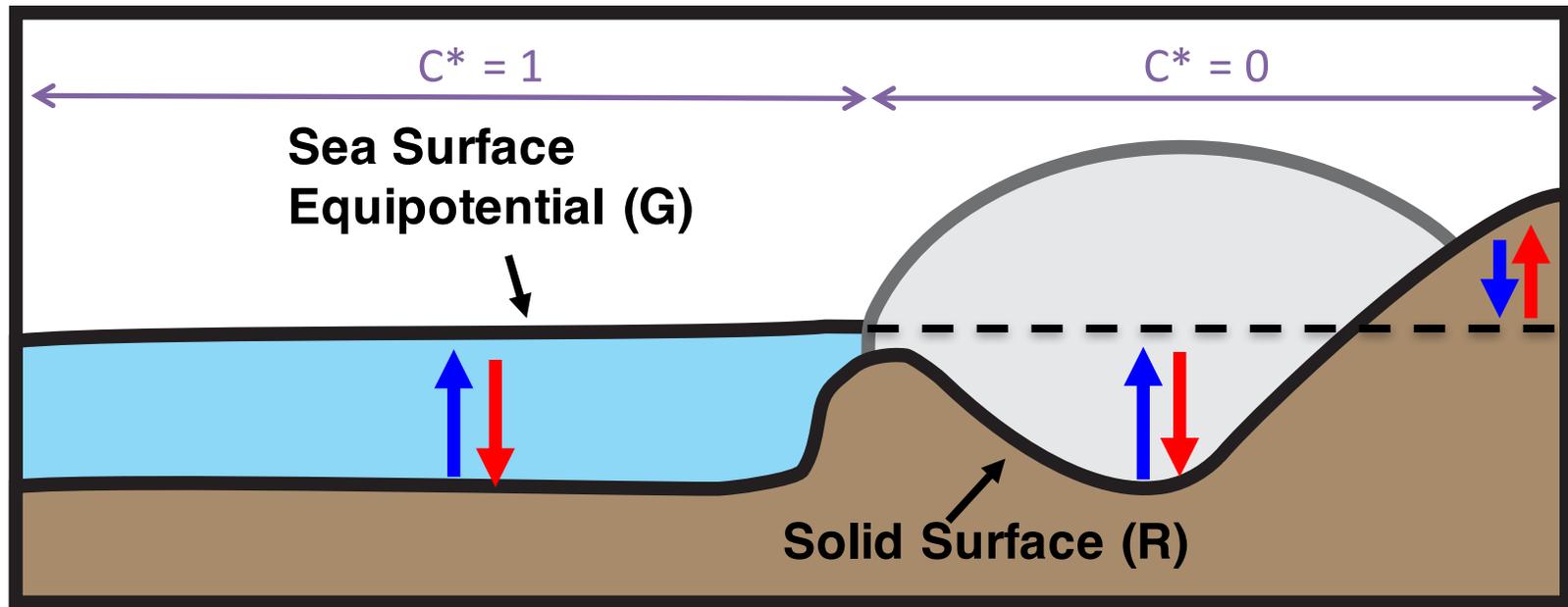
$$SL(\theta, \psi, t) \equiv G(\theta, \psi, t) - R(\theta, \psi, t).$$

 Land Elevation
(Topography)

$$T(\theta, \psi, t) \equiv R(\theta, \psi, t) - G(\theta, \psi, t), \\ = -SL(\theta, \psi, t).$$

One more definition....

2a. The Sea Level Equation



 Static Sea Level

$$SL(\theta, \psi, t) \equiv G(\theta, \psi, t) - R(\theta, \psi, t).$$

 Land Elevation
(Topography)

$$T(\theta, \psi, t) \equiv R(\theta, \psi, t) - G(\theta, \psi, t),$$

$$= -SL(\theta, \psi, t).$$

Ocean Depth: $S(\theta, \psi, t_j) = SL(\theta, \psi, t_j) C^*(\theta, \psi, t_j)$

Ocean Function: $C^*(\theta, \psi, t_j) = \begin{cases} 1 & \text{if } SL(\theta, \psi, t_j) > 0 \text{ and there is no grounded ice} \\ 0 & \text{elsewhere,} \end{cases}$

2a. The Sea Level Equation

Consider a change from the initial state at time t_0 to a new state at time t_j :

$$G(\theta, \psi, t_j) = G(\theta, \psi, t_0) + \Delta G(\theta, \psi, t_j),$$

$$R(\theta, \psi, t_j) = R(\theta, \psi, t_0) + \Delta R(\theta, \psi, t_j),$$

$$SL(\theta, \psi, t_j) = SL(\theta, \psi, t_0) + \Delta SL(\theta, \psi, t_j),$$

$$T(\theta, \psi, t_j) = T(\theta, \psi, t_0) + \Delta T(\theta, \psi, t_j),$$

With some algebra, we can find an expression for ΔS ...

▽

2a. The Sea Level Equation

Generalized Sea-Level Equation:

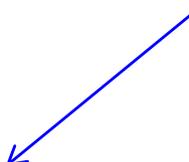
$$\Delta S(\theta, \psi, t_j) = \Delta SL(\theta, \psi, t_j)C(\theta, \psi, t_j) - T(\theta, \psi, t_0)[C(\theta, \psi, t_j) - C(\theta, \psi, t_0)],$$

$$C^*(\theta, \psi, t_j) = \begin{cases} 1 & \text{if } SL(\theta, \psi, t_j) > 0 \text{ and there is no grounded ice} \\ 0 & \text{elsewhere,} \end{cases}$$

2a. The Sea Level Equation

Generalized Sea-Level Equation

Projection of sea-level change at t_j onto the ocean function at time t_j



$$\Delta S(\theta, \psi, t_j) = \boxed{\Delta SL(\theta, \psi, t_j)C(\theta, \psi, t_j)} - T(\theta, \psi, t_0)[C(\theta, \psi, t_j) - C(\theta, \psi, t_0)],$$

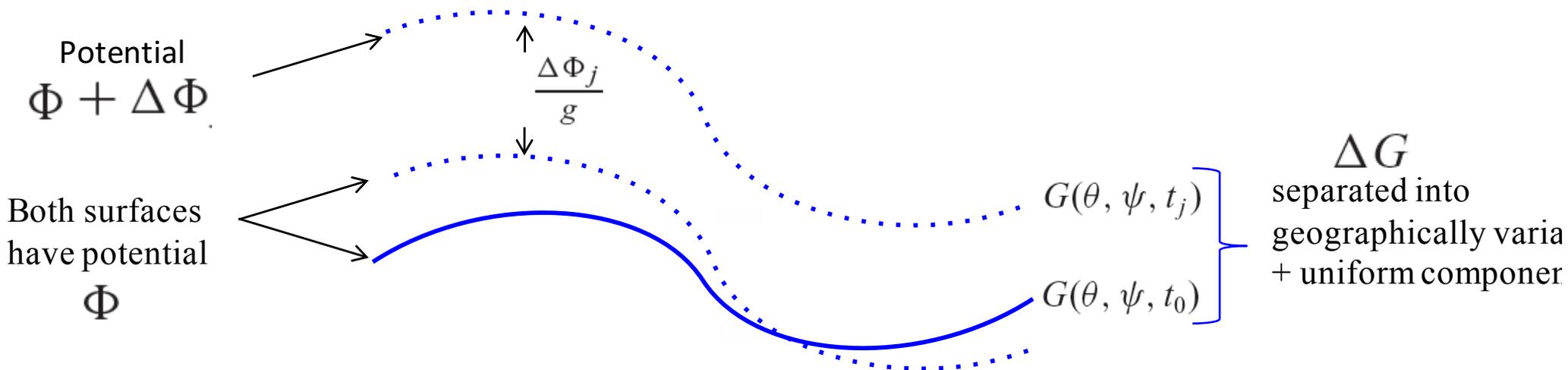
$$C^*(\theta, \psi, t_j) = \begin{cases} 1 & \text{if } SL(\theta, \psi, t_j) > 0 \text{ and there is no grounded ice} \\ 0 & \text{elsewhere,} \end{cases}$$

2a. The Sea Level Equation

Generalized Sea-Level Equation

Projection of sea-level change at t_j onto the ocean function at time t_j

$$\Delta S(\theta, \psi, t_j) = \left[\Delta \mathcal{S}\mathcal{L}(\theta, \psi, t_j) + \frac{\Delta \Phi(t_j)}{g} \right] C(\theta, \psi, t_j) - T(\theta, \psi, t_0) [C(\theta, \psi, t_j) - C(\theta, \psi, t_0)]$$

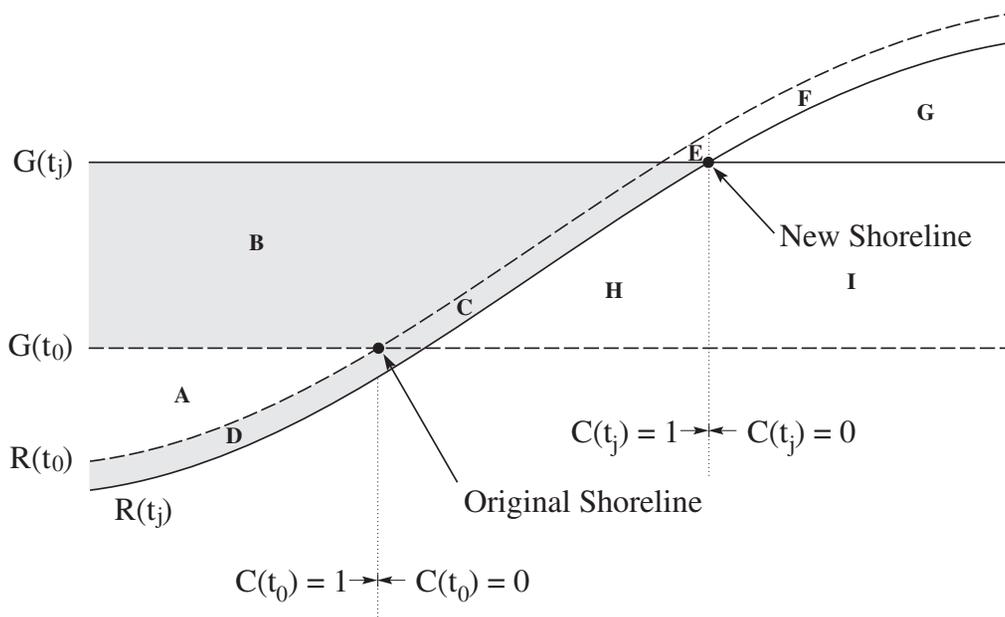


2a. The Sea Level Equation

Generalized Sea-Level Equation

$$\Delta S(\theta, \psi, t_j) = \left[\Delta \mathcal{S}\mathcal{L}(\theta, \psi, t_j) + \frac{\Delta \Phi(t_j)}{g} \right] C(\theta, \psi, t_j) - T(\theta, \psi, t_0) [C(\theta, \psi, t_j) - C(\theta, \psi, t_0)]$$

Projection of sea-level change at t_j onto the ocean function at time t_j



A topographic correction term to account for shoreline migration

2a. The Sea Level Equation

Generalized Sea-Level Equation

$$\Delta S(\theta, \psi, t_j) = \left[\Delta \mathcal{S}\mathcal{L}(\theta, \psi, t_j) + \frac{\Delta \Phi(t_j)}{g} \right] C(\theta, \psi, t_j) - T(\theta, \psi, t_0) [C(\theta, \psi, t_j) - C(\theta, \psi, t_0)]$$

Integral equation – RHS depends on ocean depth changes :

 $\Delta \mathcal{S}\mathcal{L}(\theta, \psi, t) = \Delta \mathcal{S}\mathcal{L}(\Delta I, \Delta S, \Delta \omega)$

Ocean depth changes



2a. The Sea Level Equation

SEA LEVEL MODEL INGREDIENTS

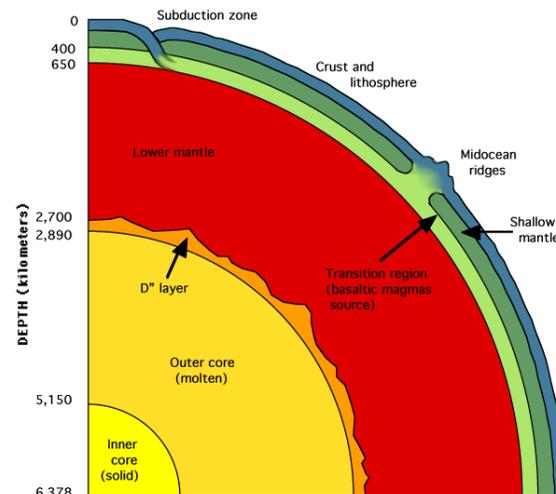
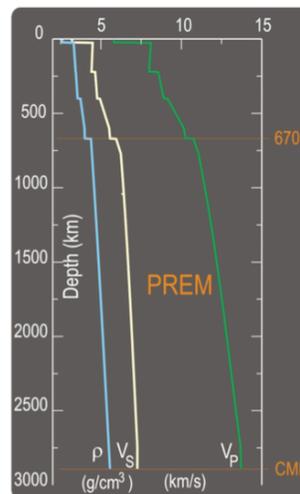
$$\Delta S\mathcal{L}(\theta, \psi, t) = \Delta S\mathcal{L}(\Delta I, \Delta S, \Delta\omega)$$

EARTH FORCING

Ice location and thickness
(inferred from modeling
GIA data, e.g. sea level)

EARTH MODEL

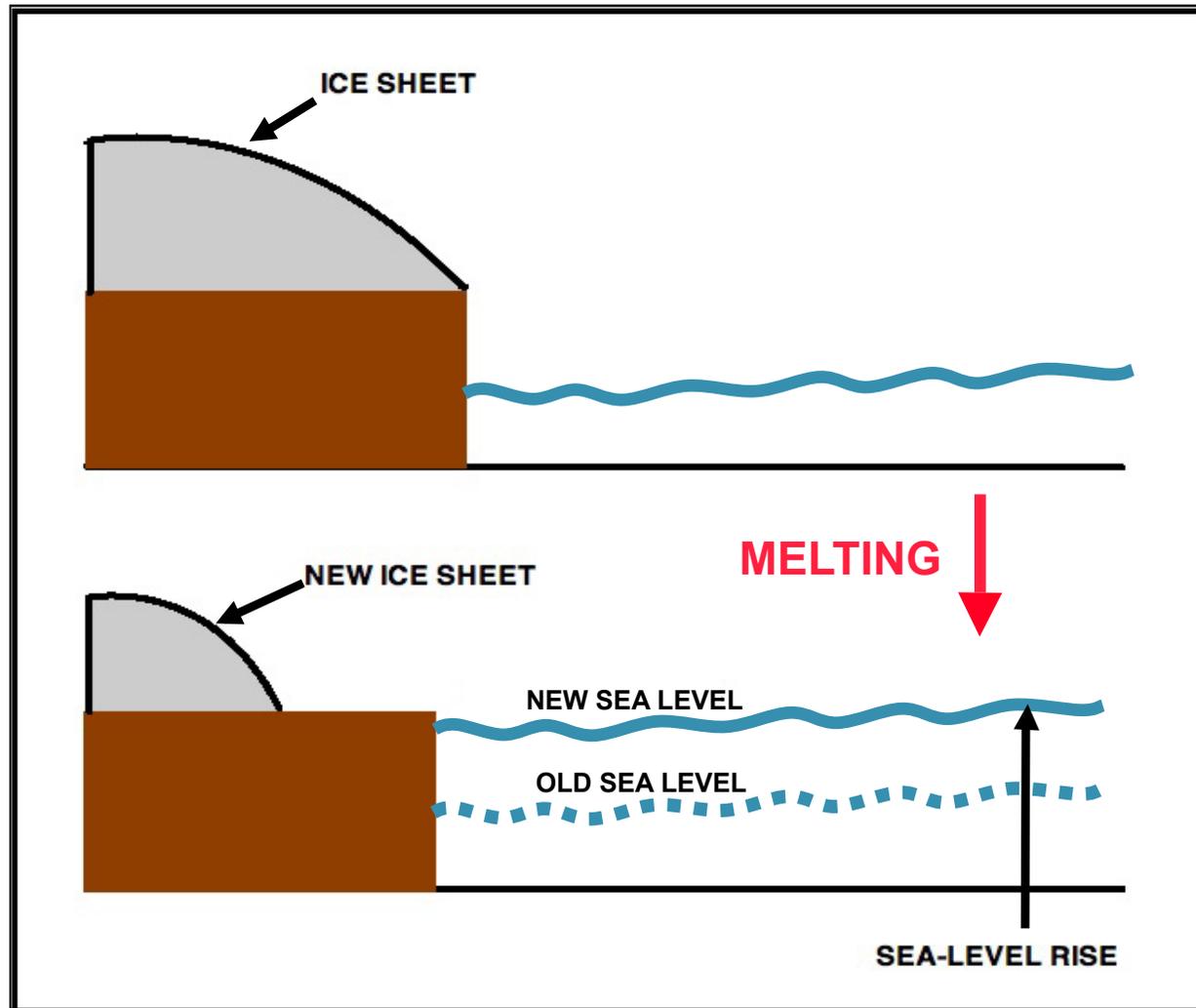
Earth geometry and elastic,
density and viscosity**
structure (PREM, GIA data) Physics...



**strong
temperature
dependence,
not well
constrained

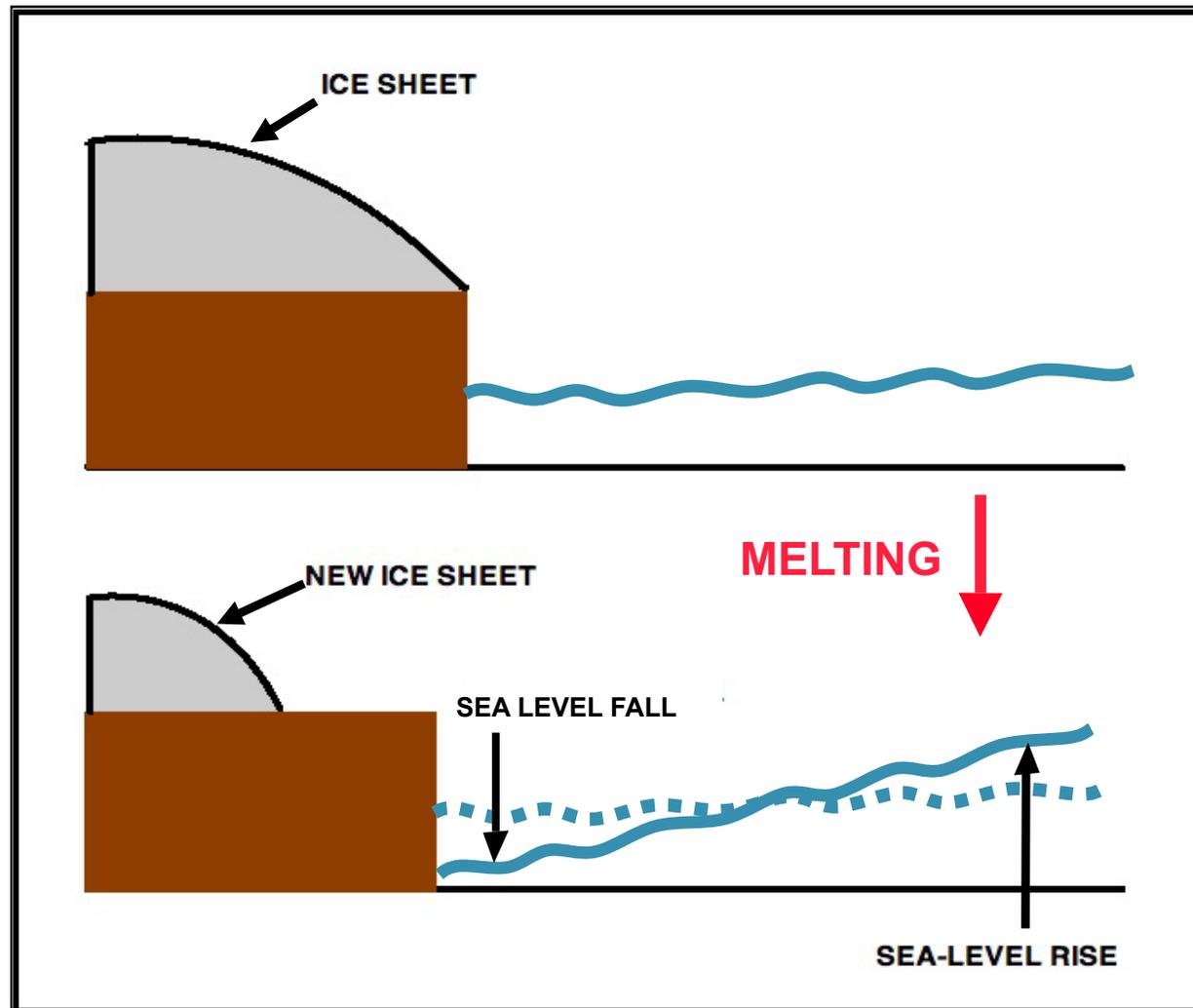
2b. Sea Level Physics

Eustatic Sea Level Change
(a.k.a. the “bath tub” model)



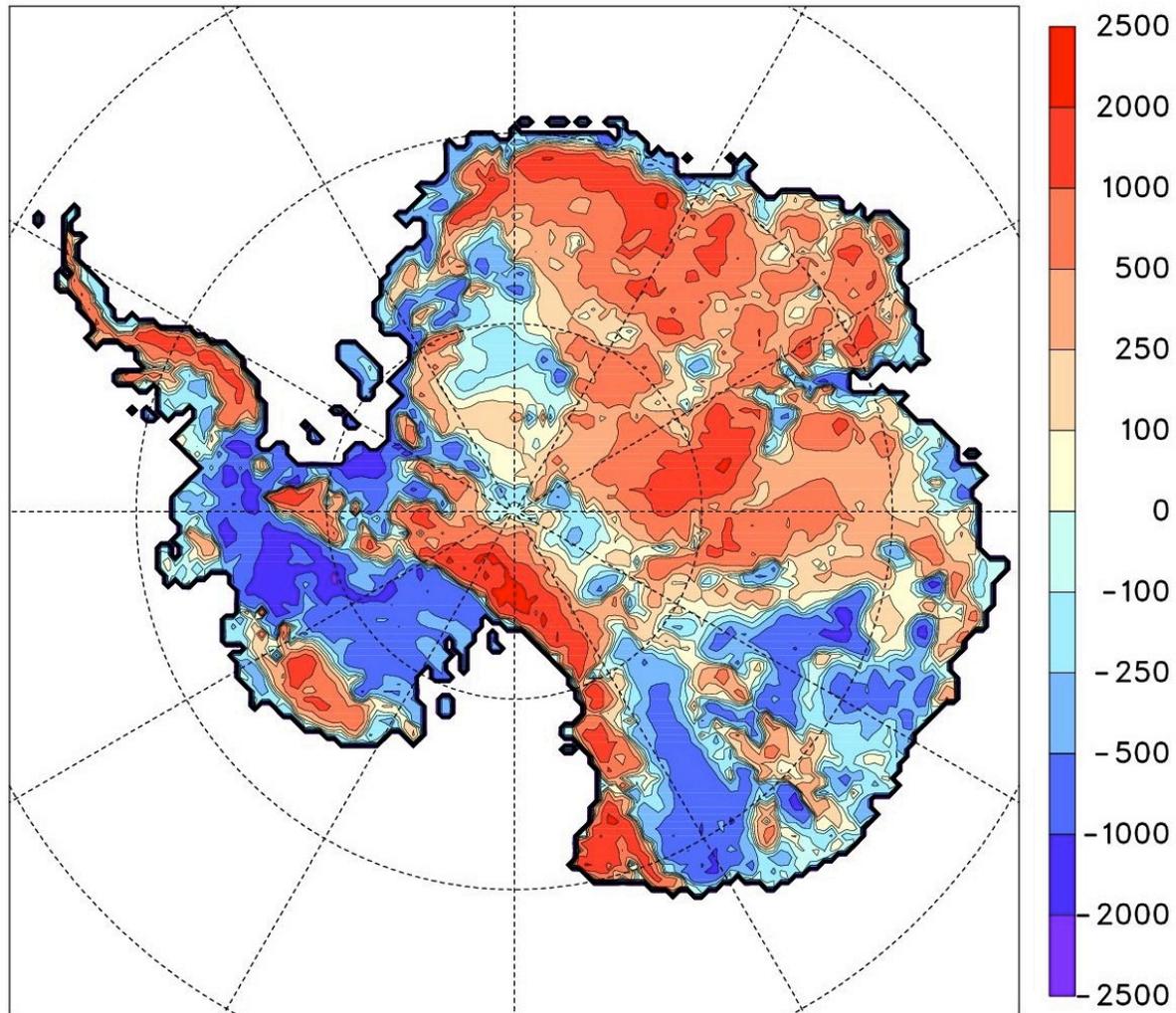
2b. Sea Level Physics

Gravitational Effect on Sea Level Change

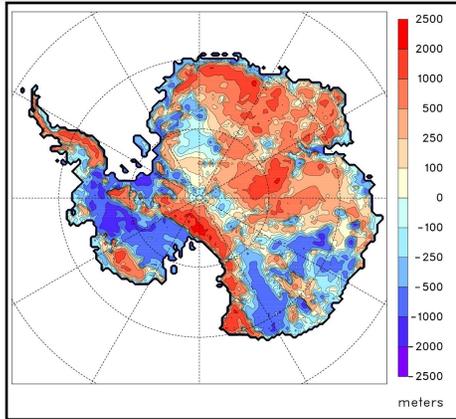


2b. Sea Level Physics

Bedrock Elevation (BEDMAP)



2b. Sea Level Physics: short timescales



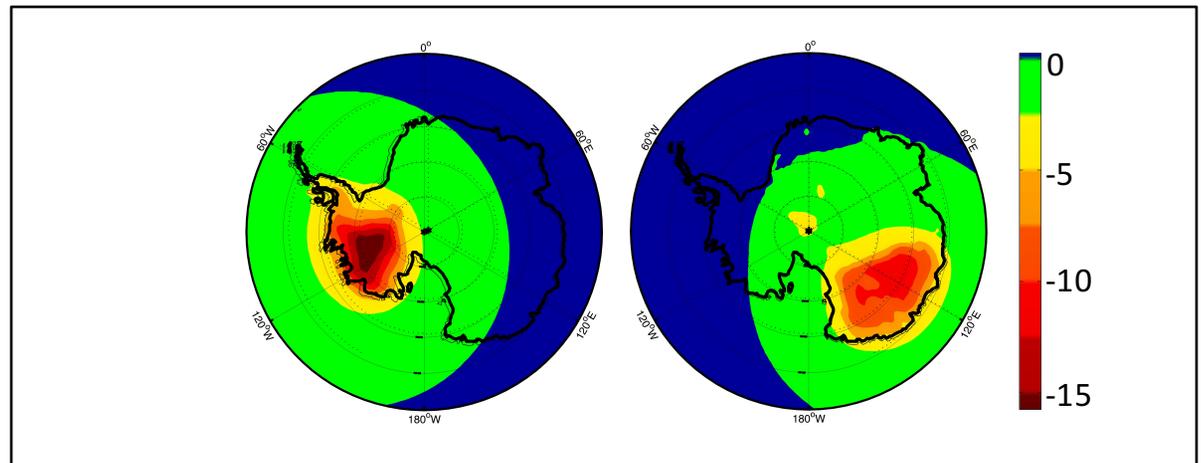
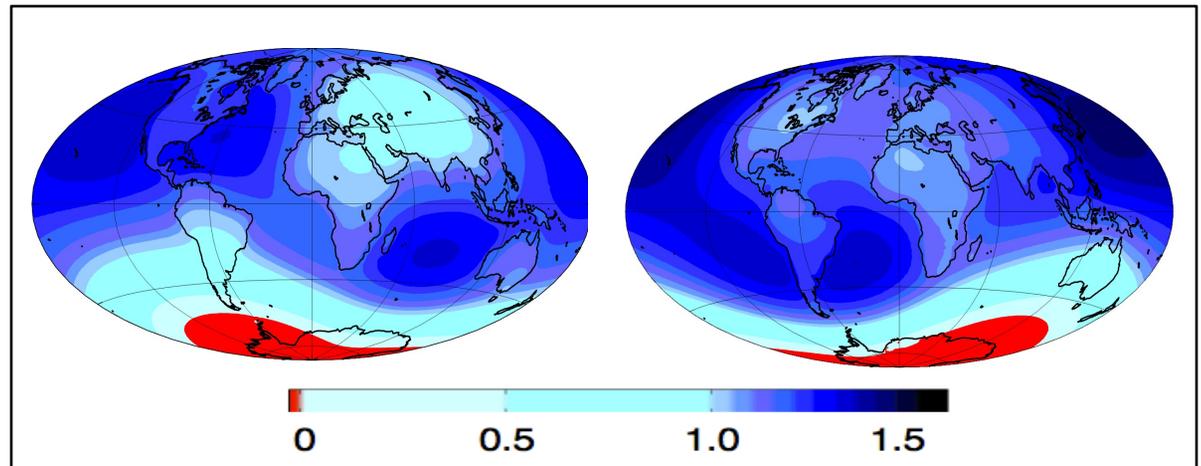
Bedrock Elevation (BEDMAP)

Normalized sea-level change following collapse of marine-based sectors of the Antarctic Ice Sheet

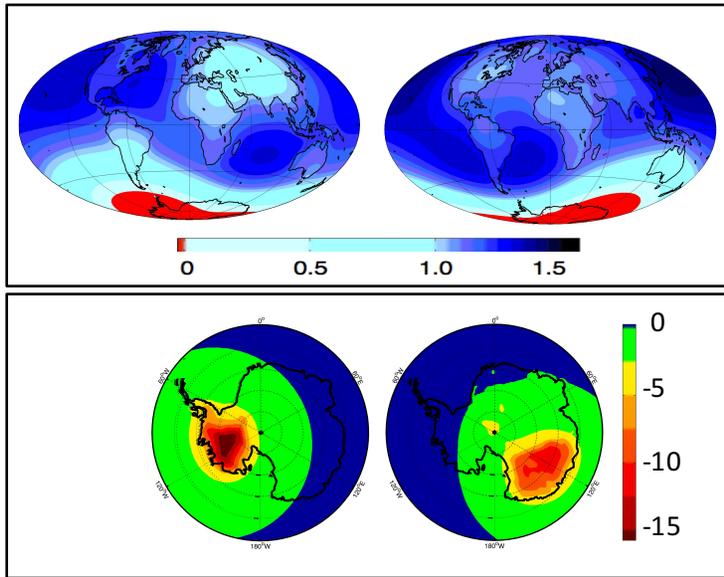
Polar projection focusing on sea level fall in the near-field

West

East

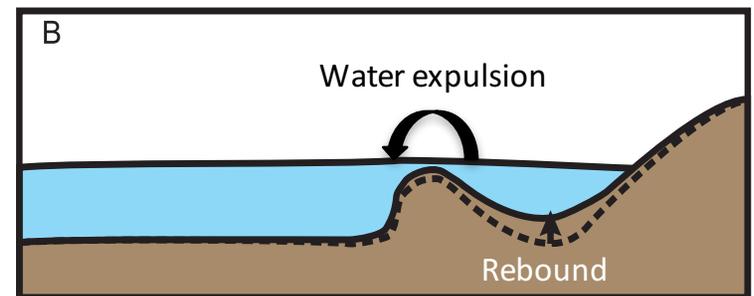
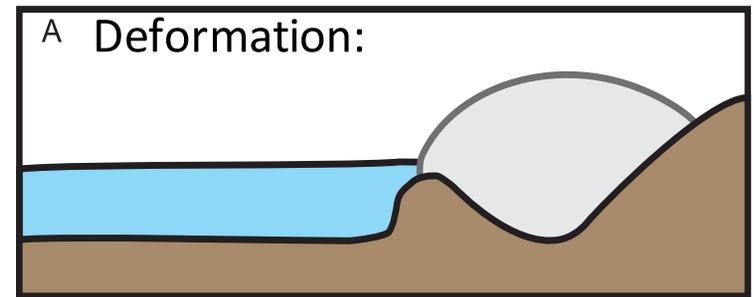
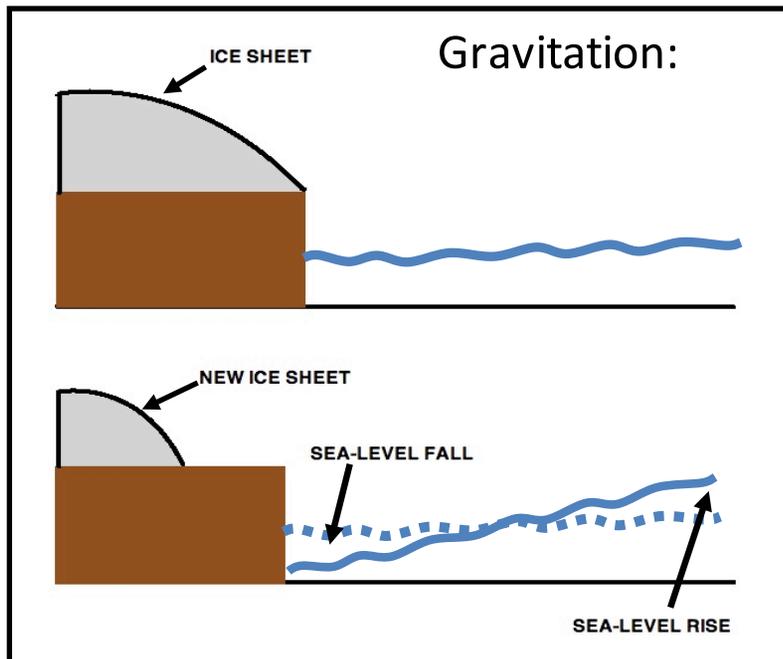
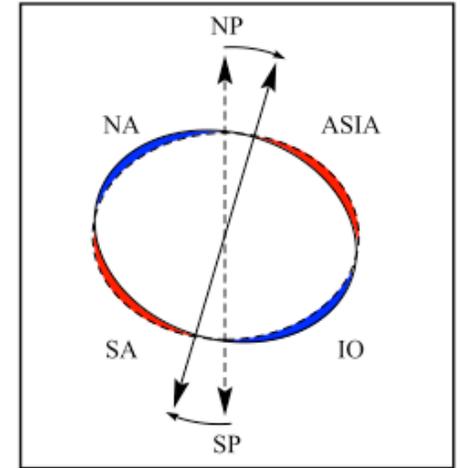
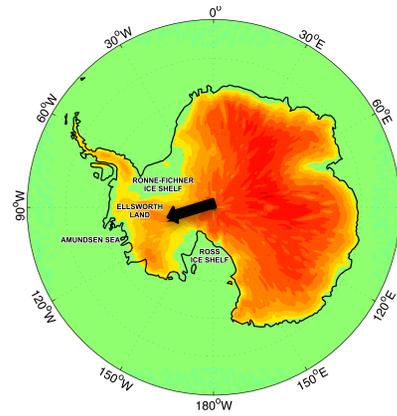


2b. Sea Level Physics: short timescales

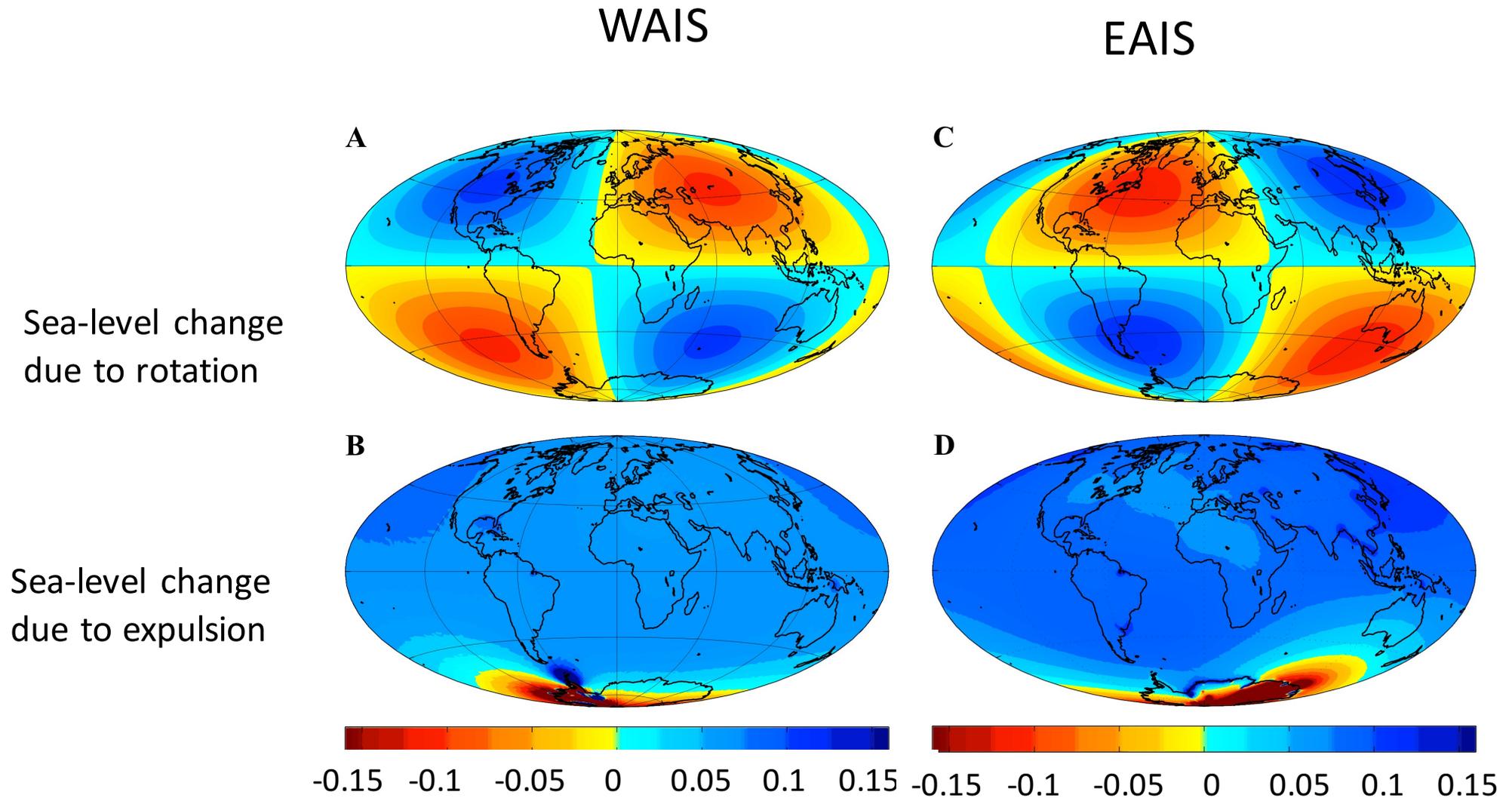


Gomez et al., GJI, 2010

Rotation:

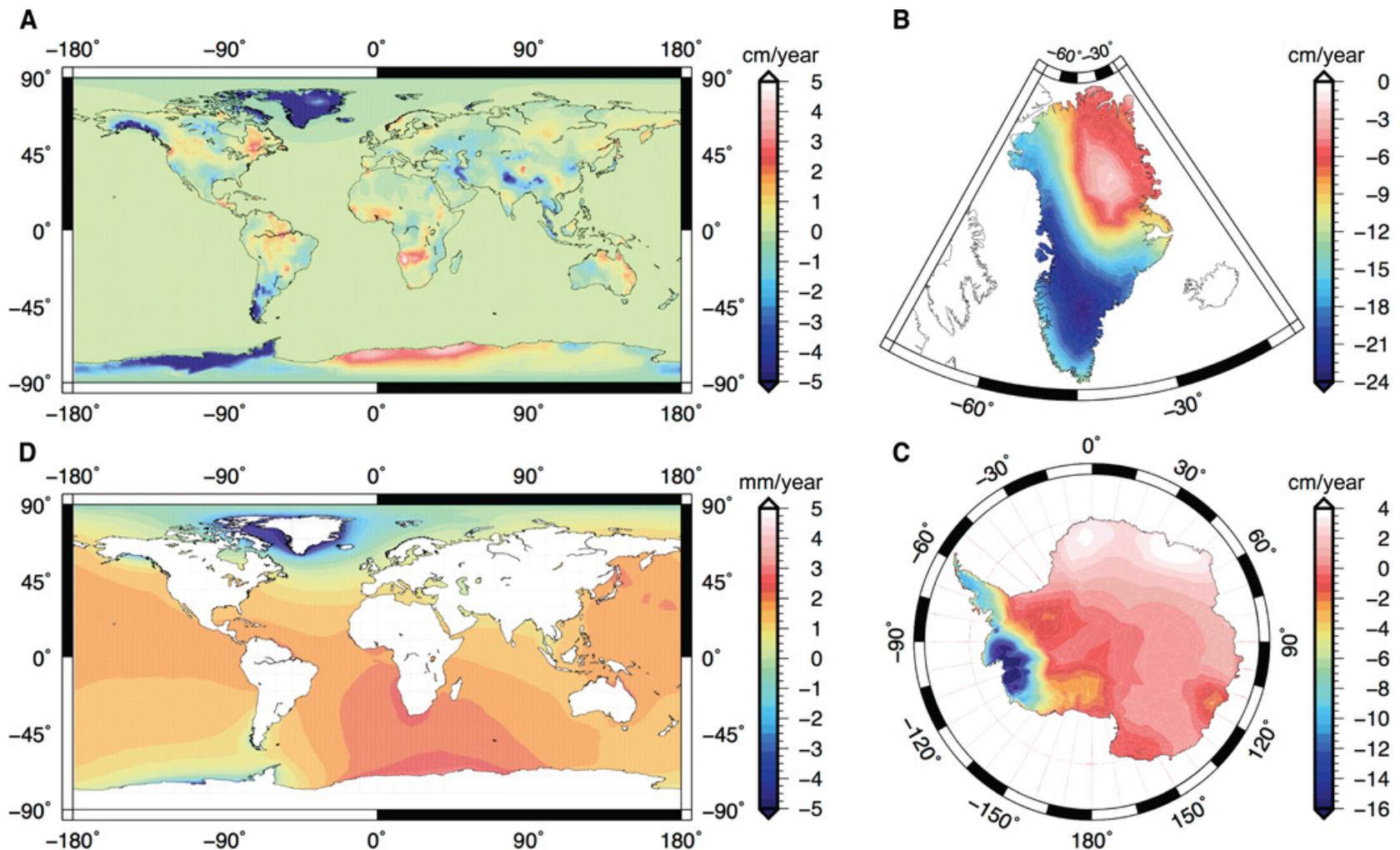


2b. Sea Level Physics: short timescales



2b. Sea Level Physics: short timescales

- <https://sealevel.nasa.gov/resources/78/cumulative-sea-level-change-since-april-2002>

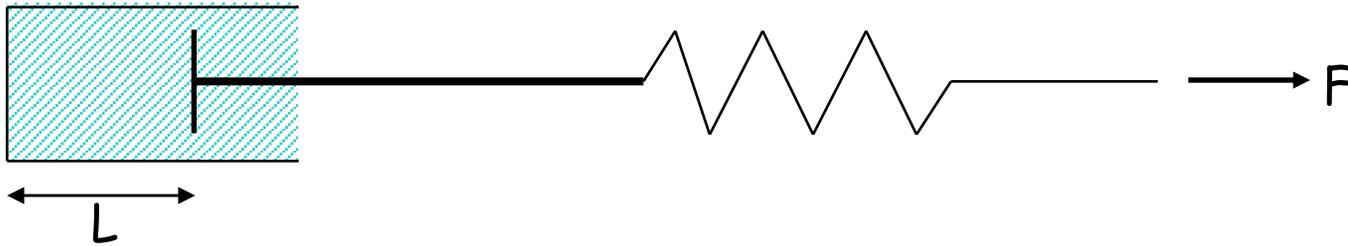


2b. Sea Level Physics: Ice-Age Timescales

The Earth behaves like a visco-elastic body

-many ways to define such a material

-assume the following 1-D linear response: “Maxwell Body”



-apply stress: immediate elastic (E) response and then subsequent viscous (v) response:
-for a Maxwell Body:

$$\frac{\Delta L}{L} = \left(\frac{\Delta L}{L}\right)^E + \left(\frac{\Delta L}{L}\right)^v \Rightarrow$$

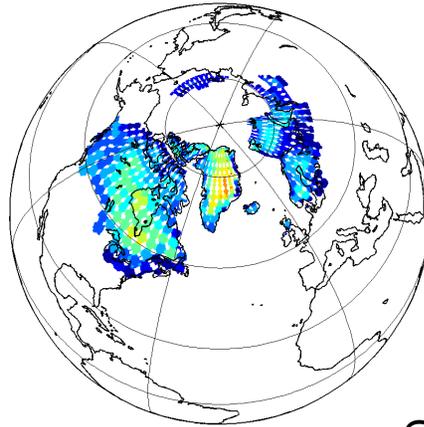
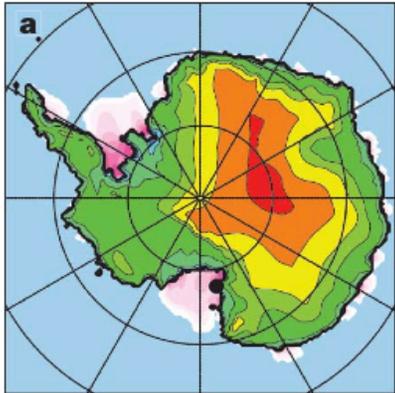
$$\frac{d}{dt}\left(\frac{\Delta L}{L}\right) = \frac{d}{dt}\left(\frac{\Delta L}{L}\right)^E + \frac{d}{dt}\left(\frac{\Delta L}{L}\right)^v \Rightarrow \frac{d}{dt}e = \frac{d}{dt}\left(\frac{F}{k}\right) + \frac{F}{\eta}$$

$$\dot{e} = \frac{1}{k}\dot{F} + \frac{1}{\eta}F$$

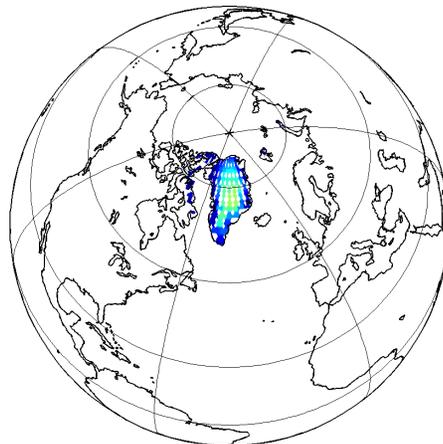
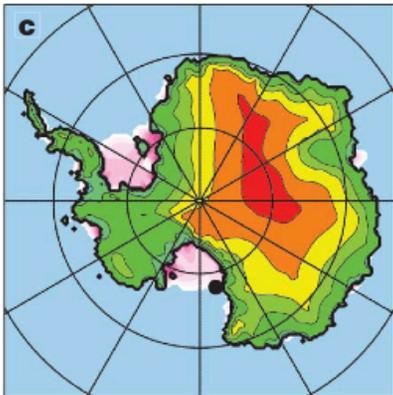
Constitutive equation (i.e. equation relating stress & strain) for a Maxwell visco-elastic 1-D body

2b. Sea Level Physics: Ice-Age Timescales

Last Glacial Maximum



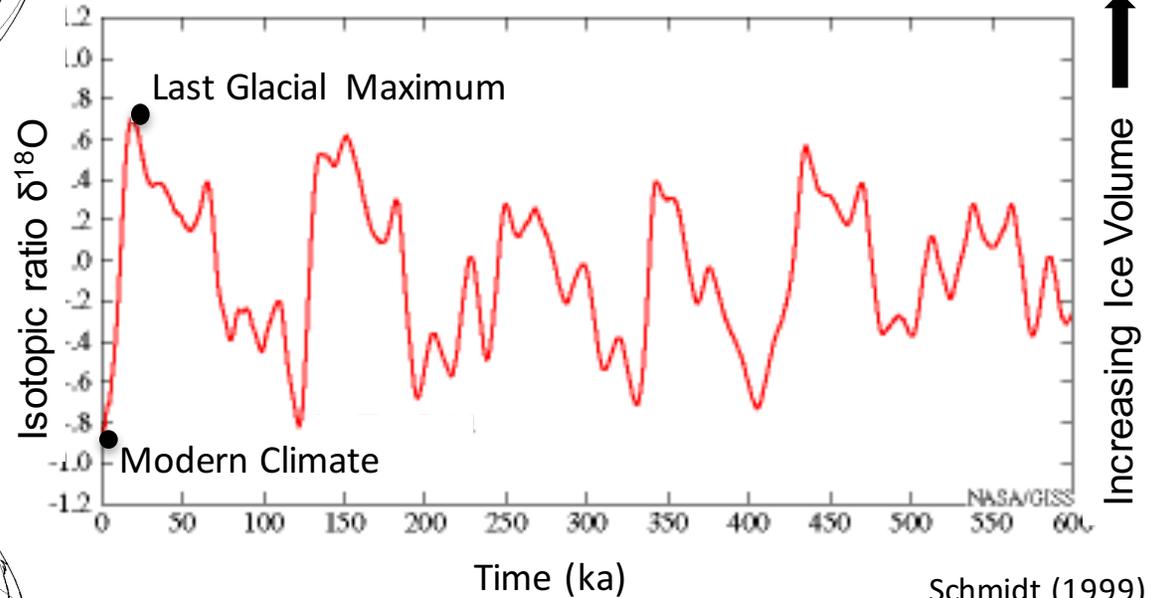
Modern Climate



Pollard & DeConto (2009)

Peltier (2004) ICE5G

Averaged $\delta^{18}\text{O}$ in deep sea sediment carbonate



Schmidt (1999)

2b. Sea Level Physics: Ice-Age Timescales

MORE
ICE

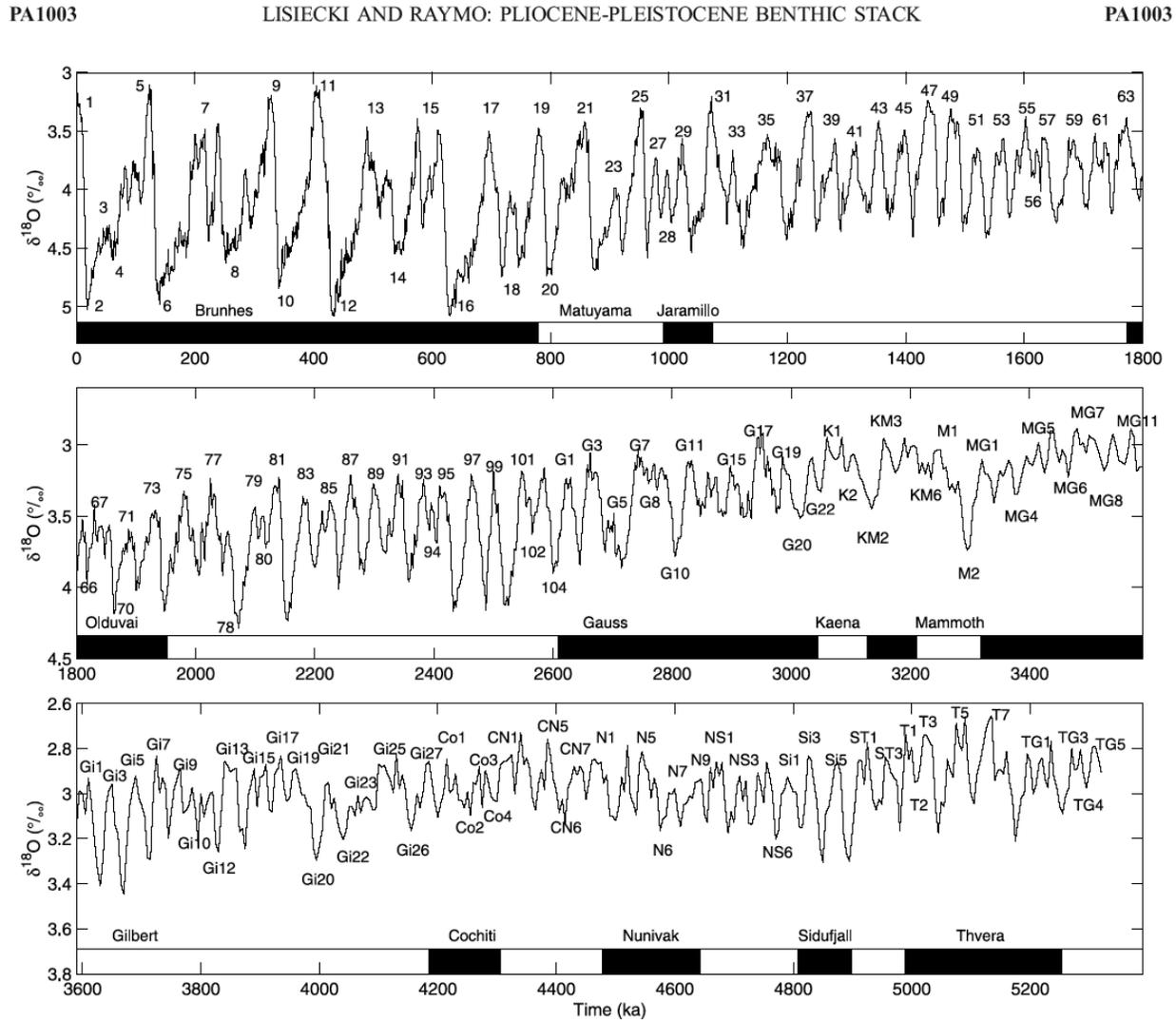
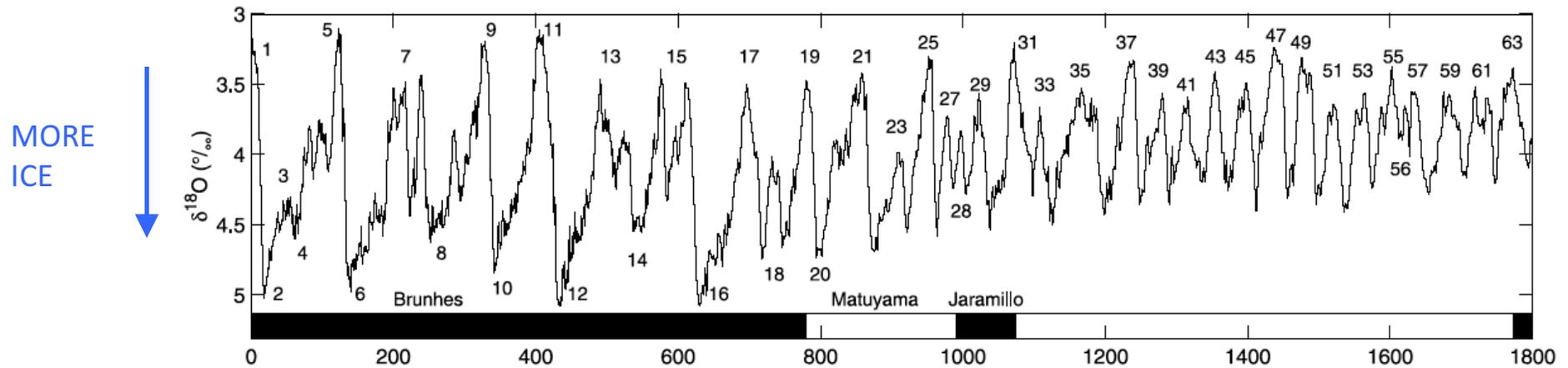
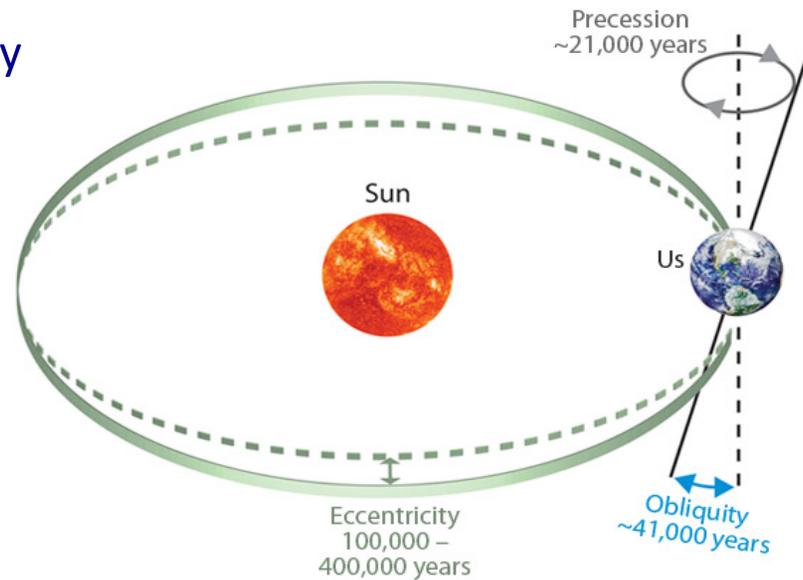


Figure 4. The LR04 benthic $\delta^{18}\text{O}$ stack constructed by the graphic correlation of 57 globally distributed benthic $\delta^{18}\text{O}$ records. The stack is plotted using the LR04 age model described in section 5 and with new MIS labels for the early Pliocene (section 6.2). Note that the scale of the vertical axis changes across panels.

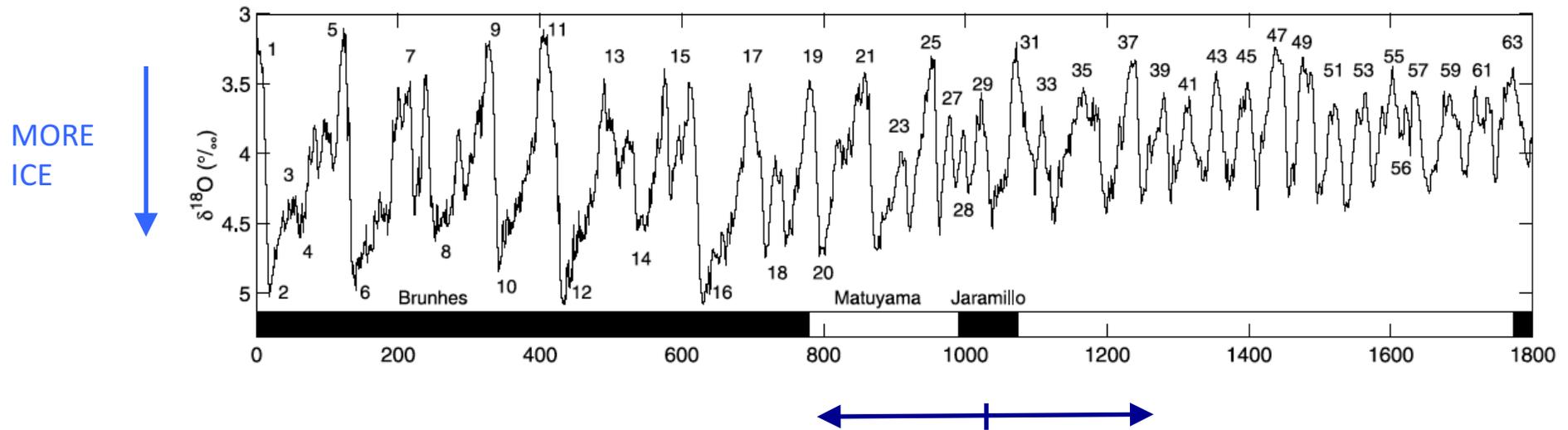
2b. Sea Level Physics: Ice-Age Timescales



Milankovitch Theory



2b. Sea Level Physics: Ice-Age Timescales

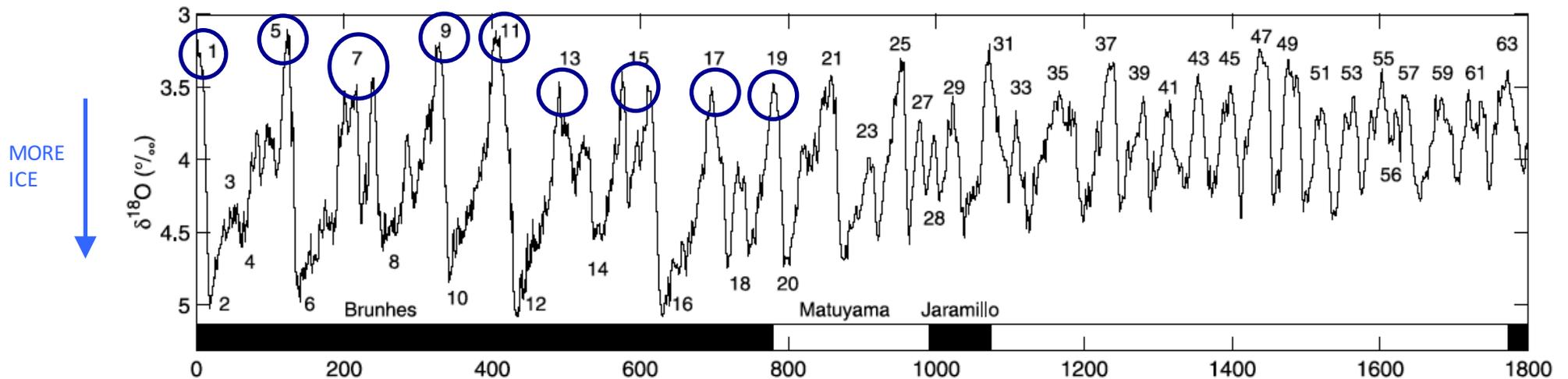


Transition from ice age cycles
with a dominant 40 kyr period to
cycles with a 100 kyr period.

Clark et al. (Paleoceanography, 1998)
Huybers (Science, 2006)
Raymo et al. (Science, 2006)

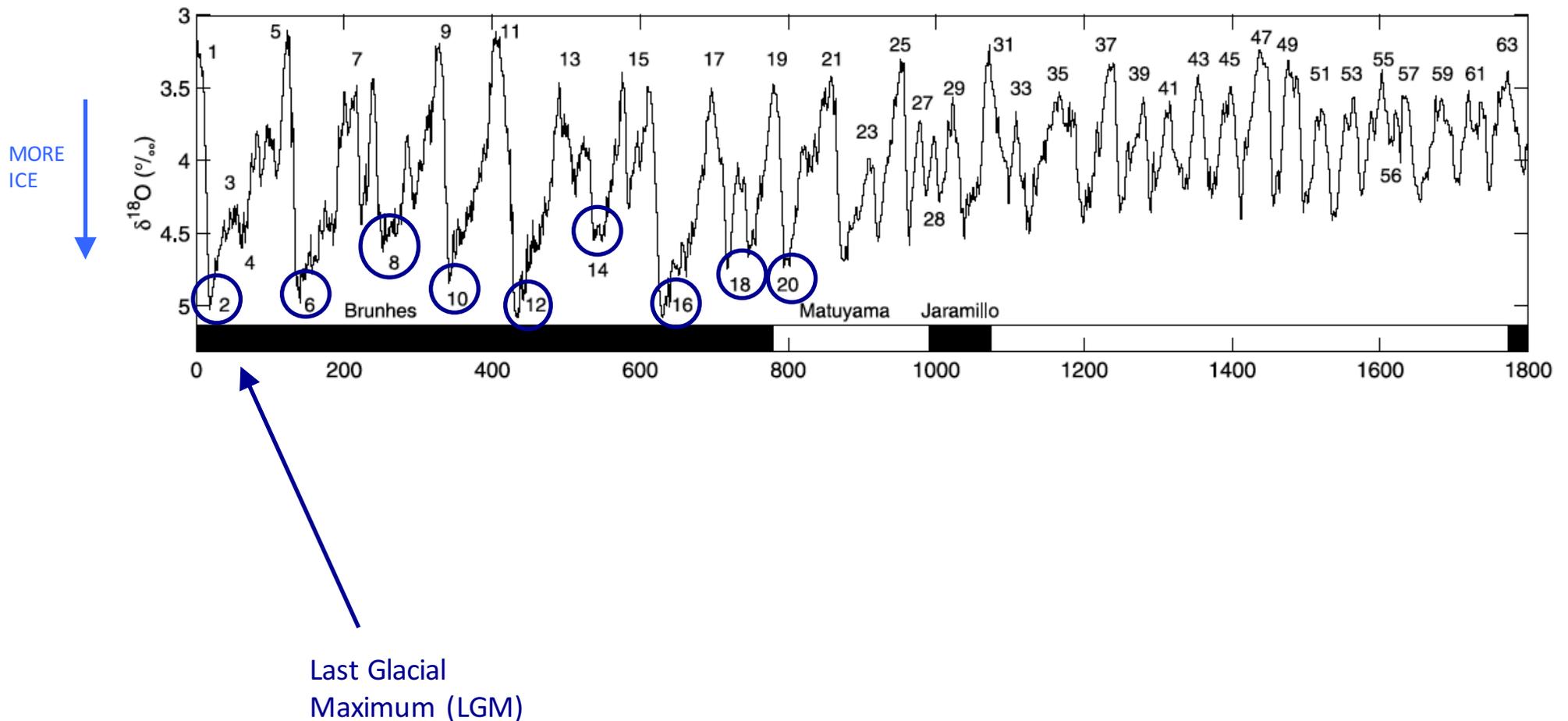
2b. Sea Level Physics: Ice-Age Timescales

Interglacials ...



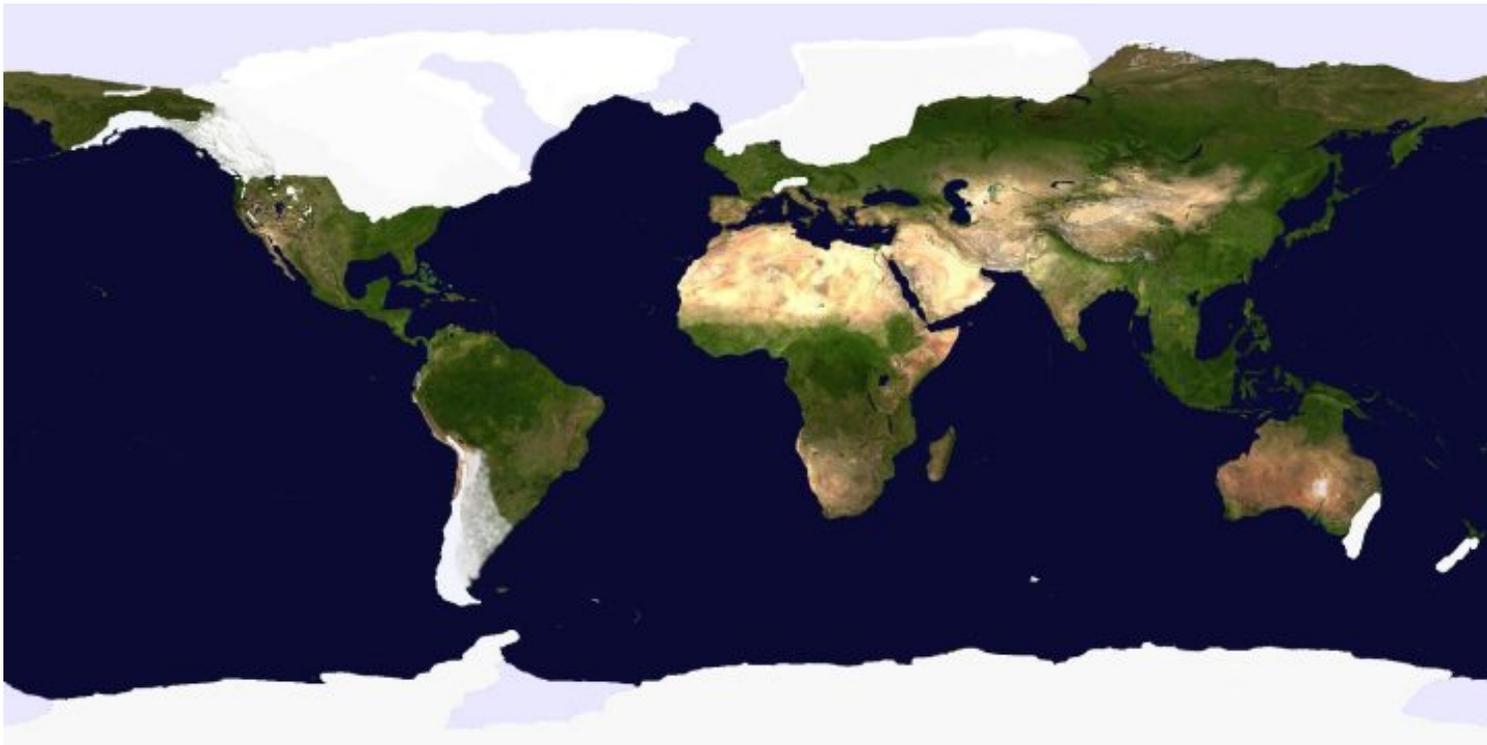
2b. Sea Level Physics: Ice-Age Timescales

Glacial Maxima ...



2b. Sea Level Physics: Ice-Age Timescales

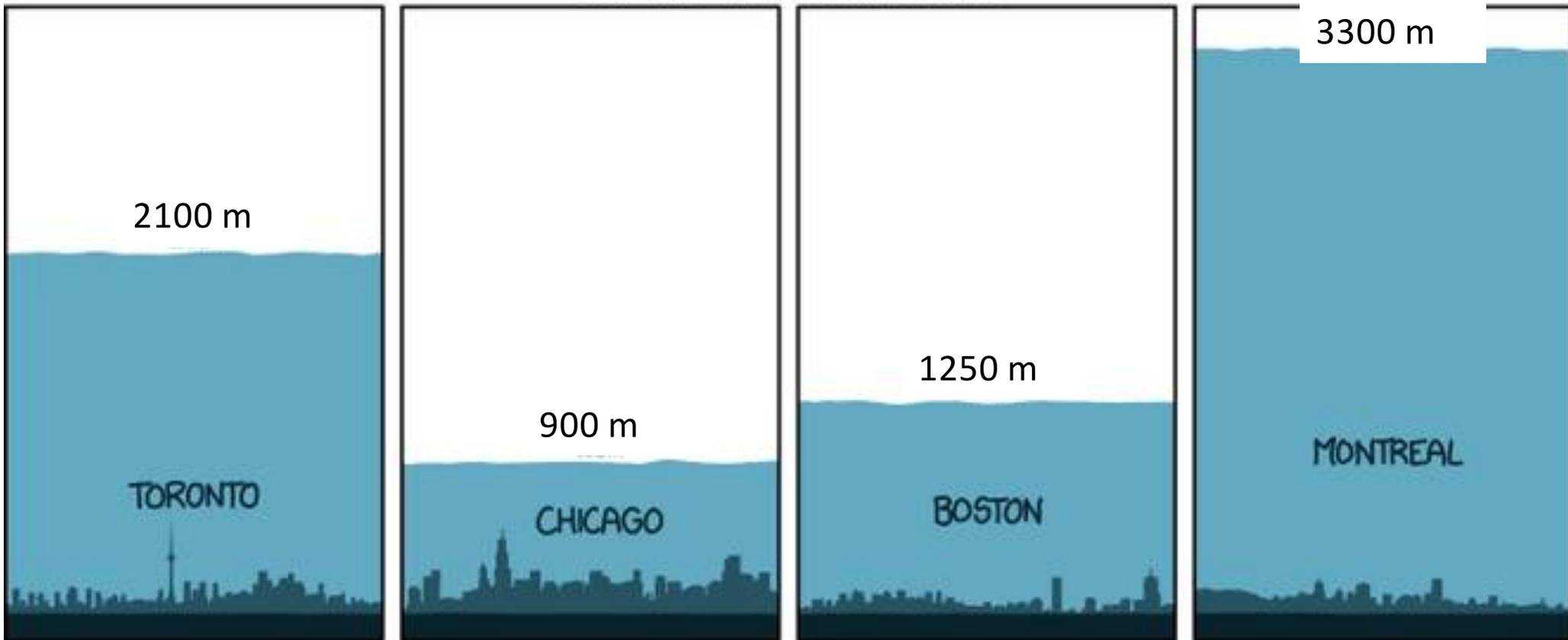
LGM



... difference in ice volume relative to the present-day is sufficient to raise globally averaged sea level by ~ 130 m ($\sim \frac{1}{2}$ of this is associated with the Laurentide ice

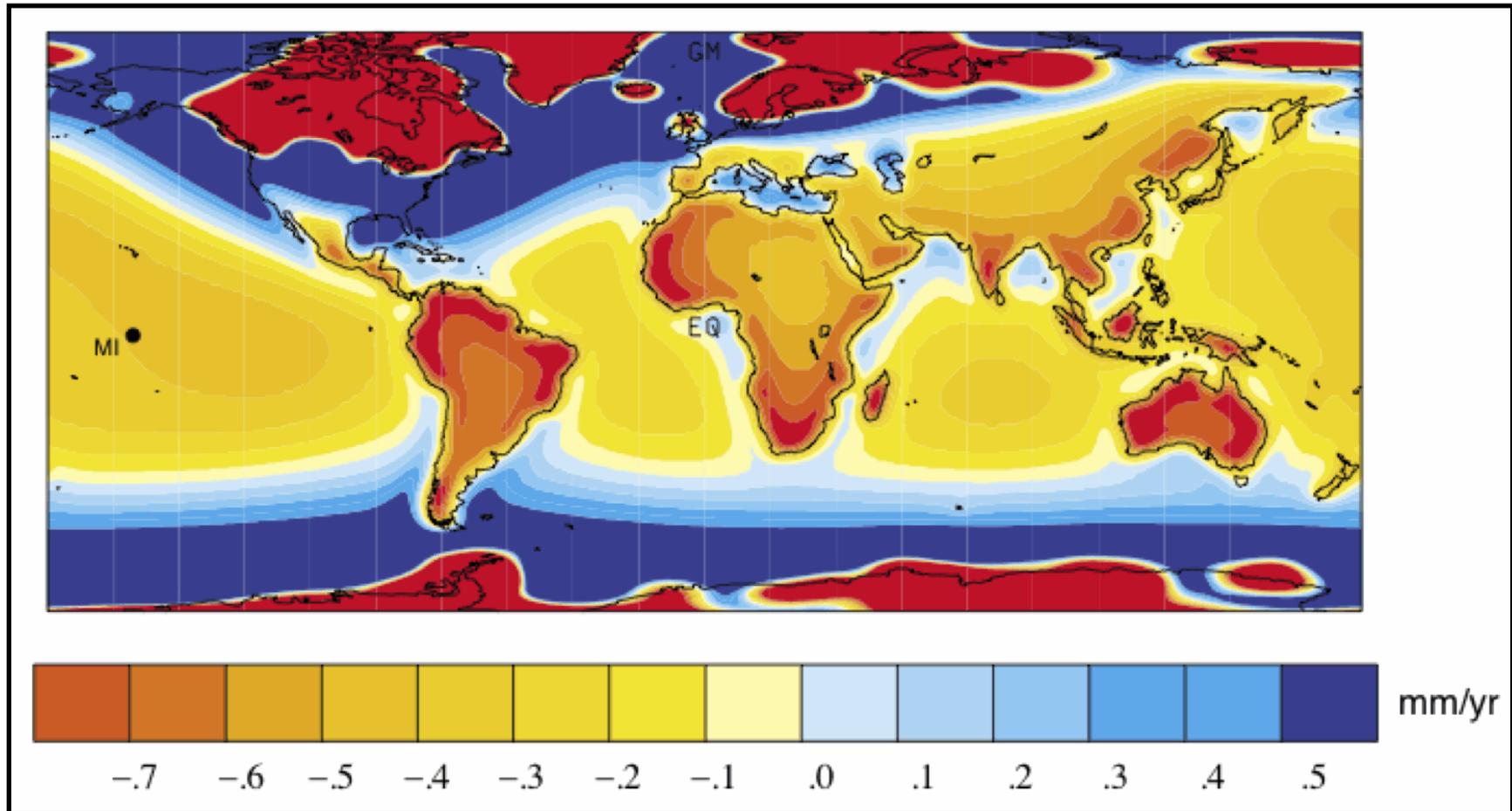
THICKNESS OF THE ICE SHEETS

AT VARIOUS LOCATIONS
21,000 YEARS AGO
COMPARED WITH MODERN SKYLINES



2b. Sea Level Physics: Ice-Age Timescales

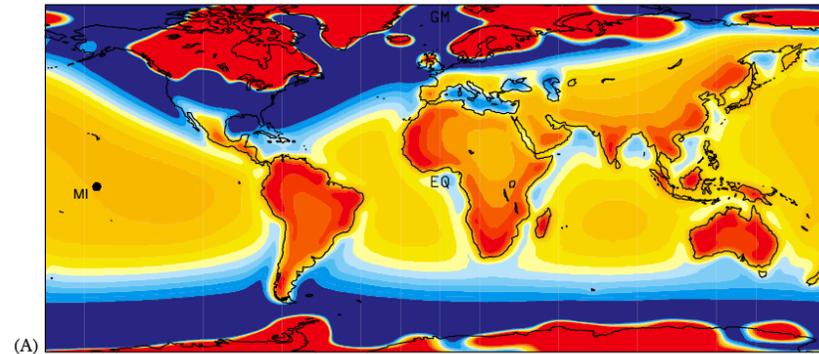
Numerical prediction of the present-day rate of change of global sea level due to ongoing GIA effects from the last ice age



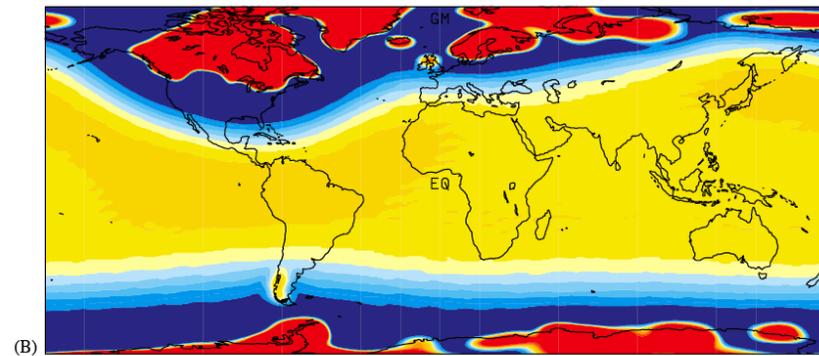
Mitrovica and Milne (2002)

2b. Sea Level Physics: Ice-Age Timescales

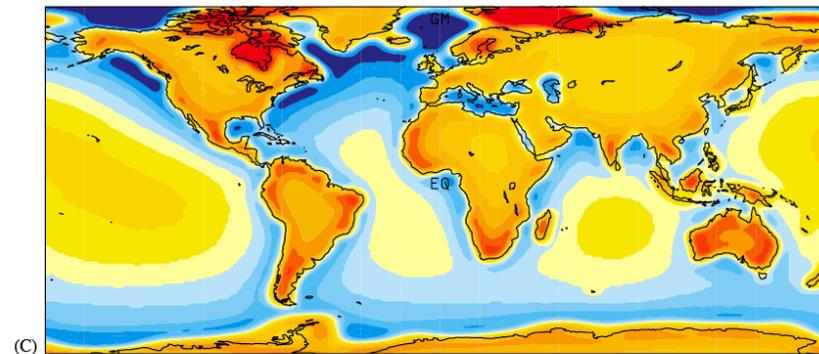
Numerical Prediction of Present-Day Rate of Global Sea-Level Change Due to Ongoing GIA



TOTAL



ICE



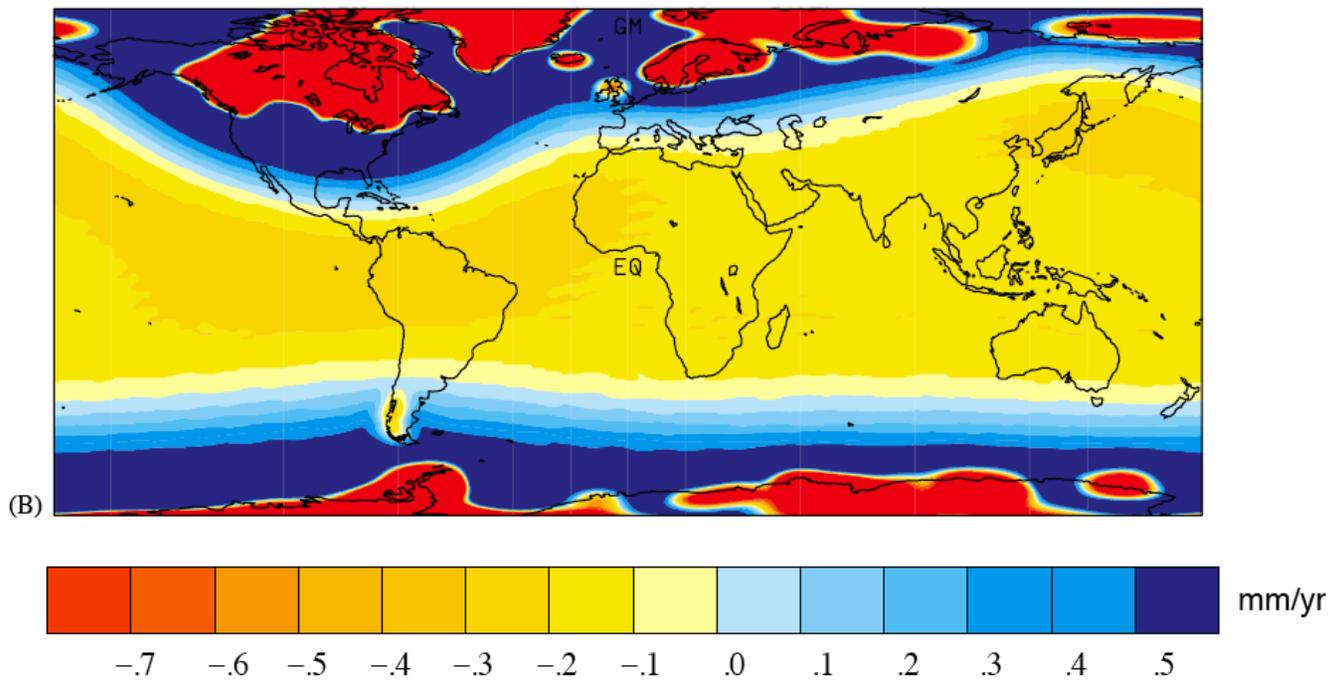
OCEAN



Mitrovica and Milne (2002)

2b. Sea Level Physics: Ice-Age Timescales

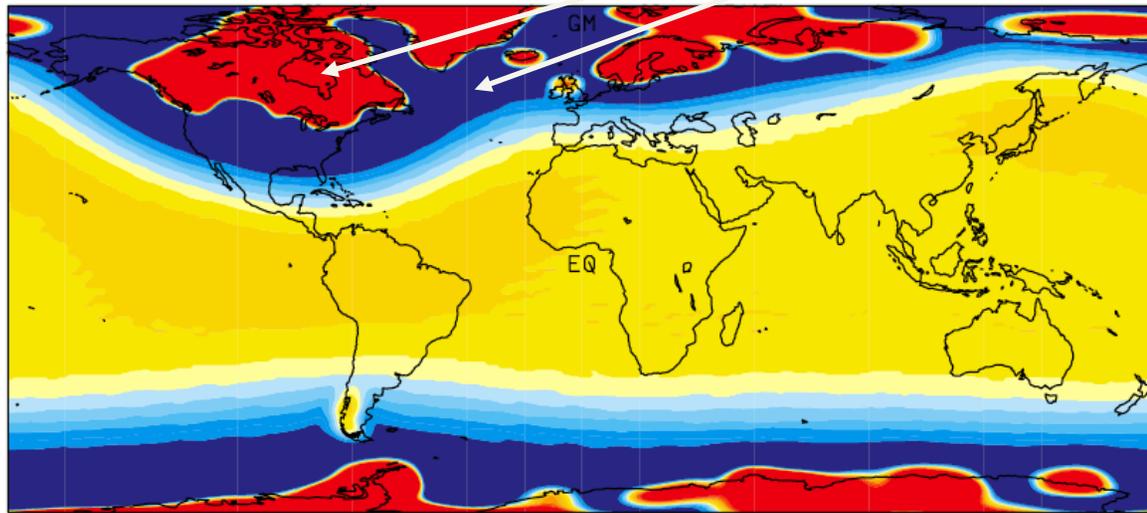
Numerical Prediction of Present-Day Rate of Global Sea-Level Change Due to Ongoing GIA



Ice Signal

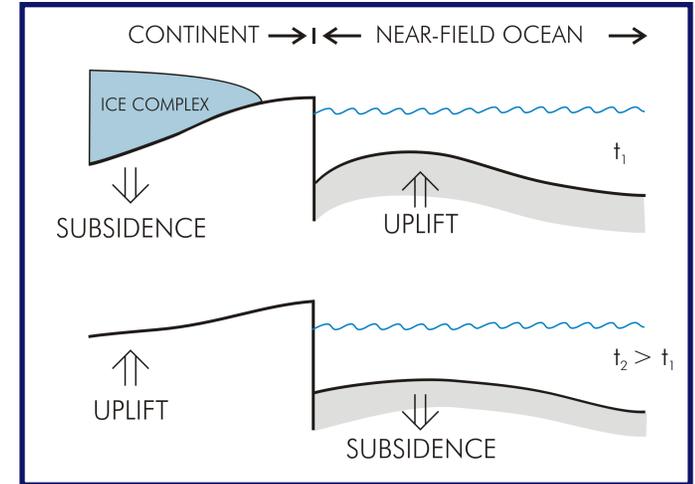
2b. Sea Level Physics: Ice-Age Timescales

Numerical Prediction of Present-Day Rate of Global Sea-Level Change Due to Ongoing GIA



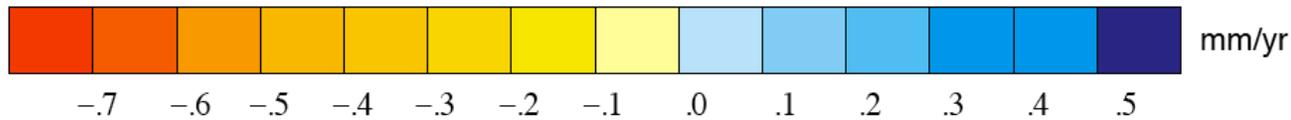
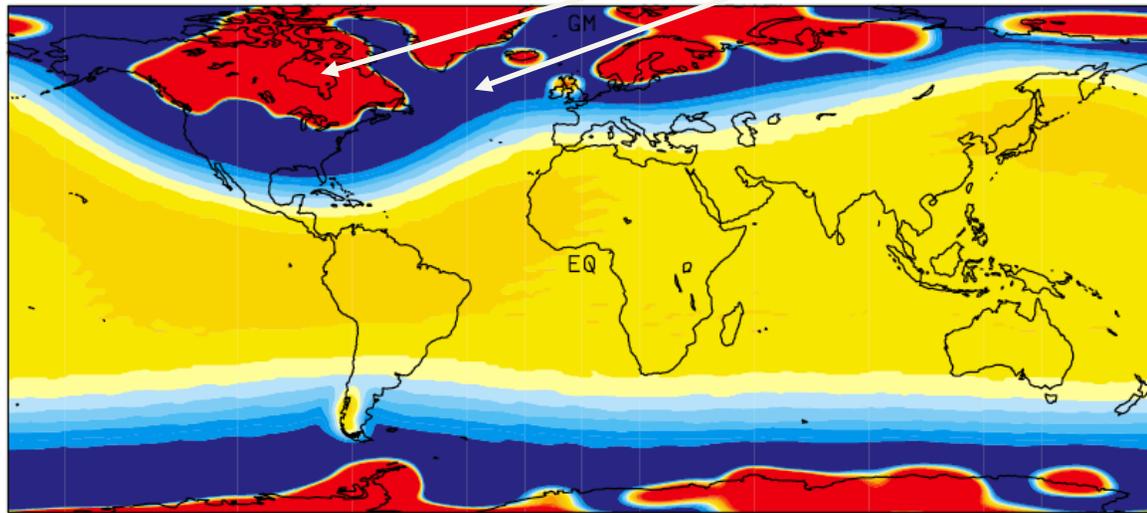
Ice Signal

NEAR FIELD



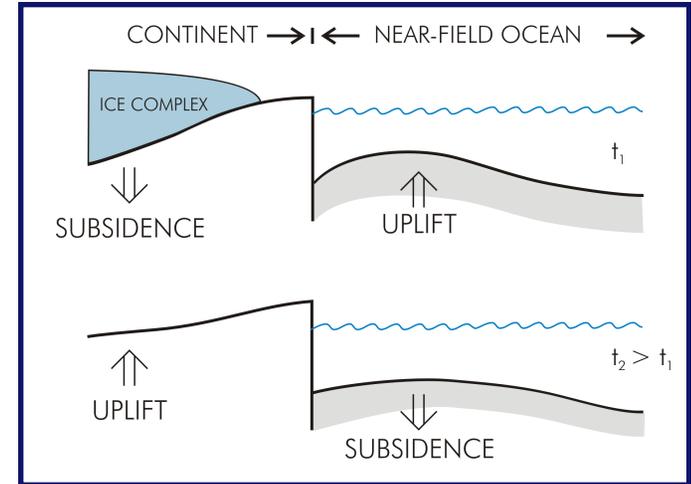
2b. Sea Level Physics: Ice-Age Timescales

Numerical Prediction of Present-Day Rate of Global Sea-Level Change Due to Ongoing GIA



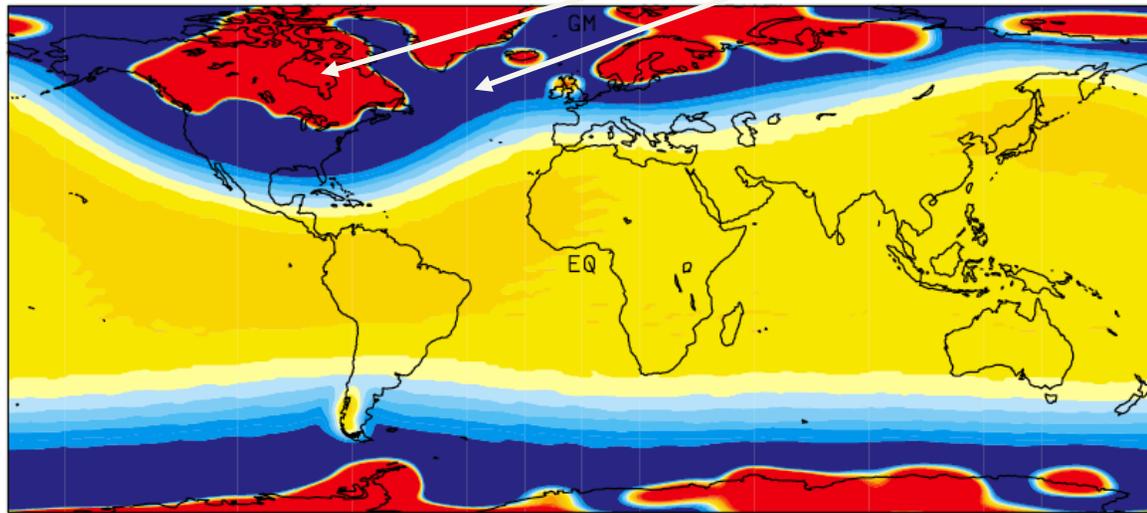
Ice Signal

NEAR FIELD

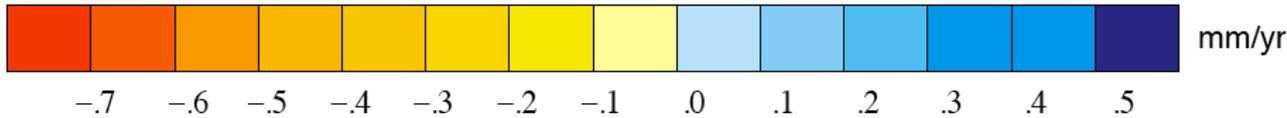


2b. Sea Level Physics: Ice-Age Timescales

Numerical Prediction of Present-Day Rate of Global Sea-Level Change Due to Ongoing GIA

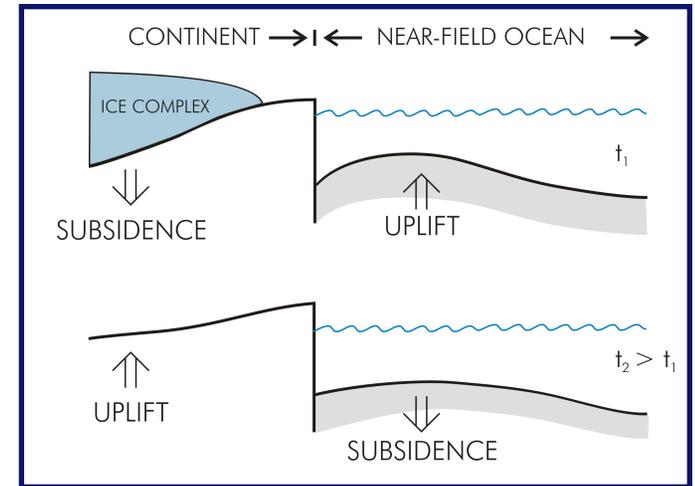


(B)



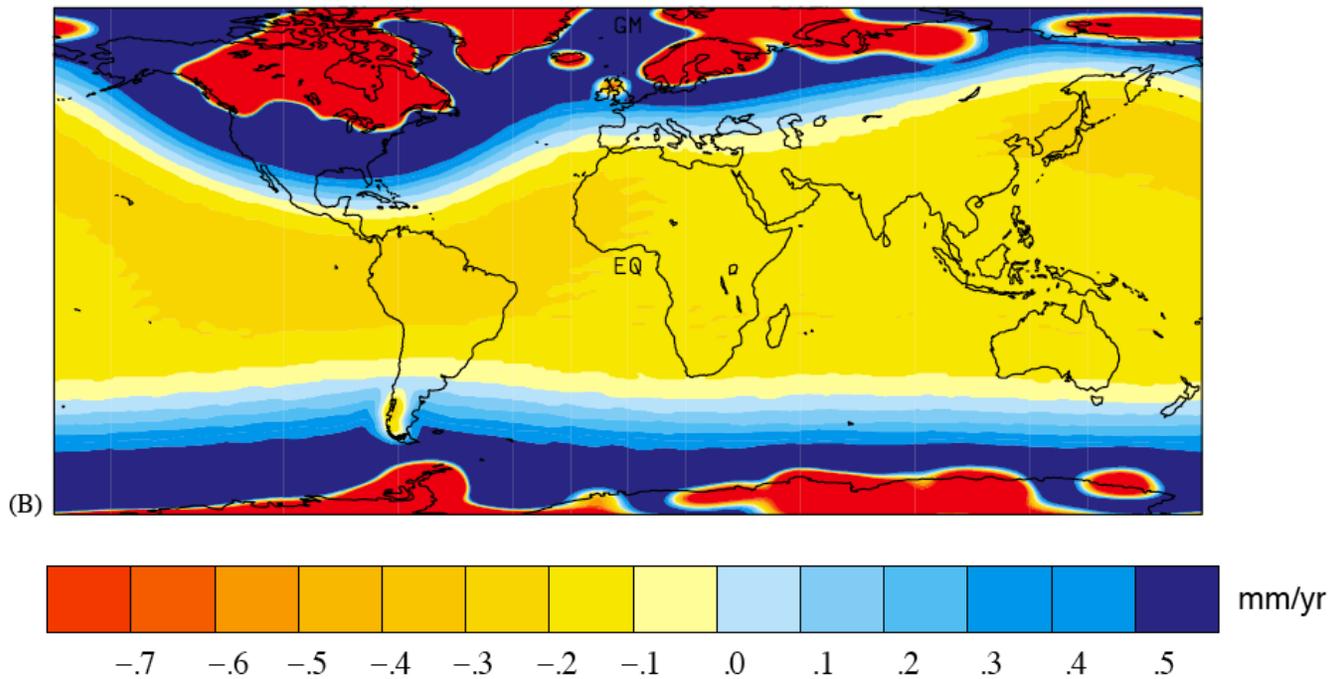
Ice Signal

NEAR FIELD



2b. Sea Level Physics: Ice-Age Timescales

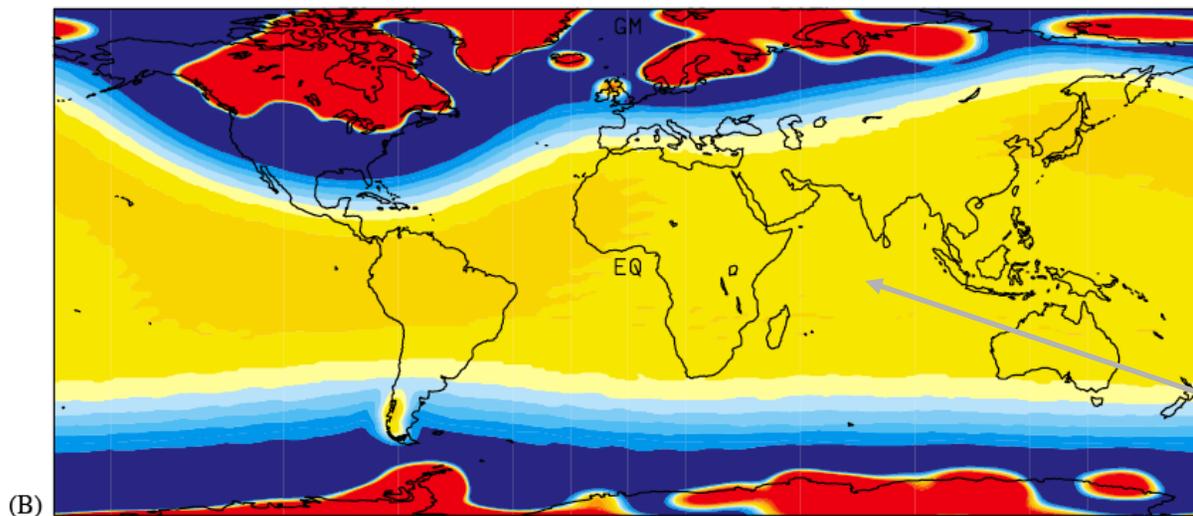
Numerical Prediction of Present-Day Rate of Global Sea-Level Change Due to Ongoing GIA



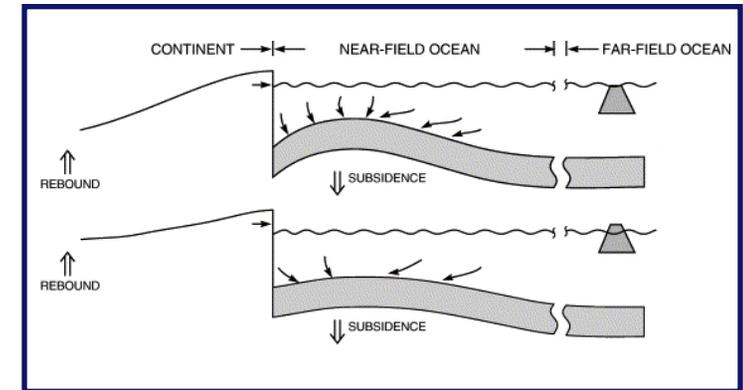
Ice Signal

2b. Sea Level Physics: Ice-Age Timescales

Numerical Prediction of Present-Day Rate of Global Sea-Level Change Due to Ongoing GIA



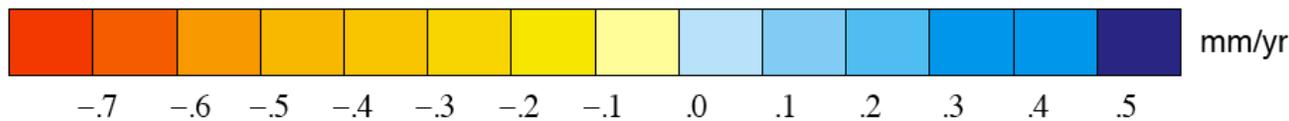
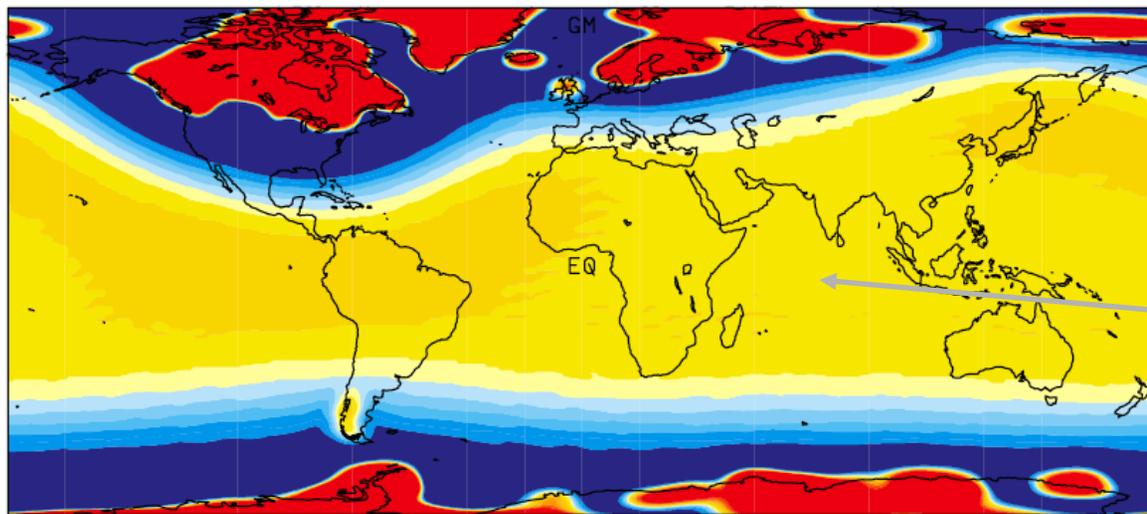
Ice Signal



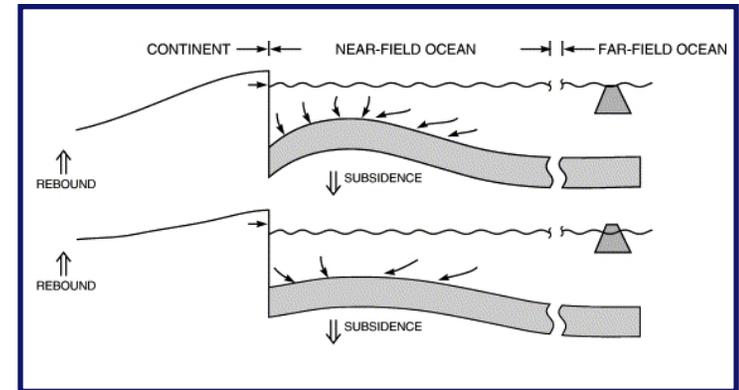
FAR FIELD

2b. Sea Level Physics: Ice-Age Timescales

Numerical Prediction of Present-Day Rate of Global Sea-Level Change Due to Ongoing GIA



Ice Signal



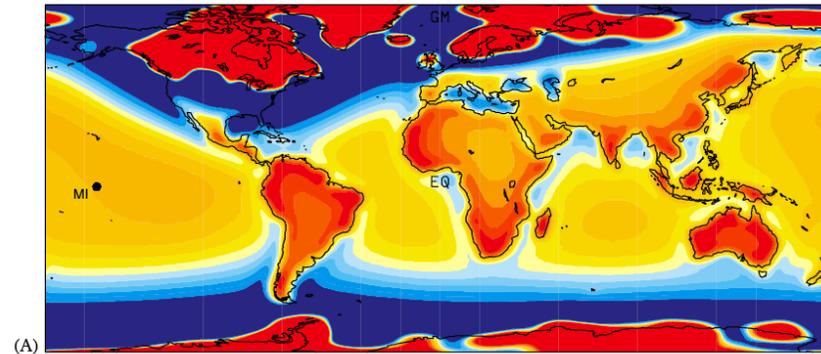
FAR FIELD



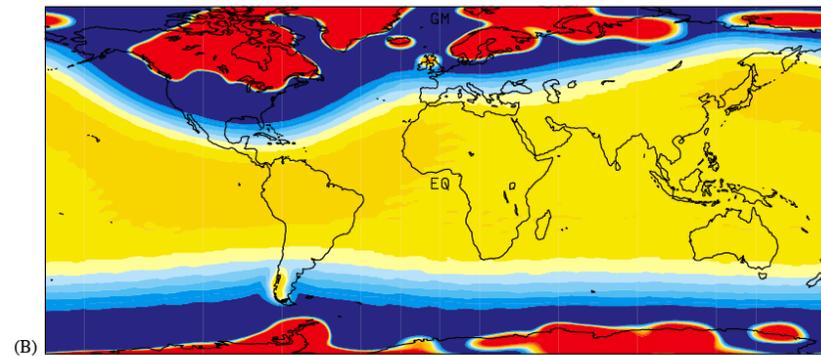
Equatorial Ocean
Syphoning!

2b. Sea Level Physics: Ice-Age Timescales

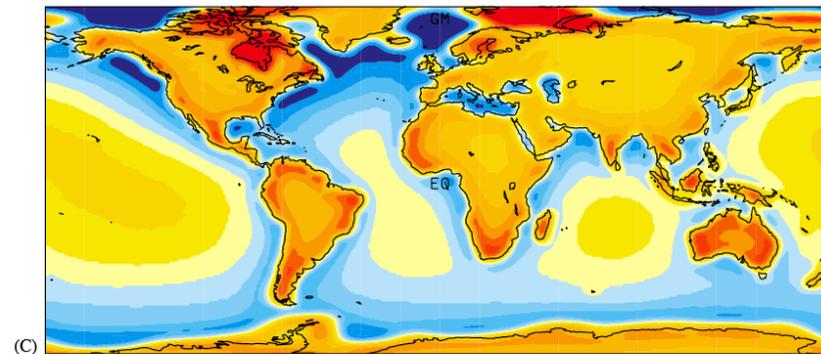
Numerical Prediction of Present-Day Rate of Global Sea-Level Change Due to Ongoing GIA



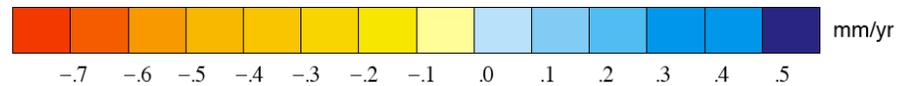
TOTAL



ICE

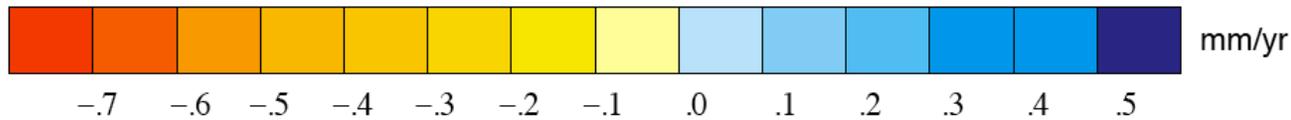
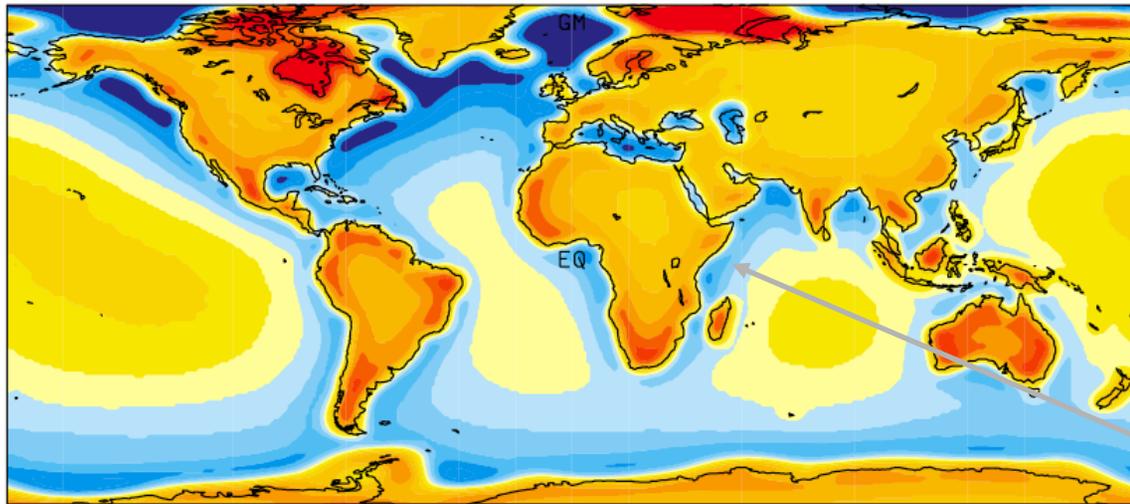


OCEAN

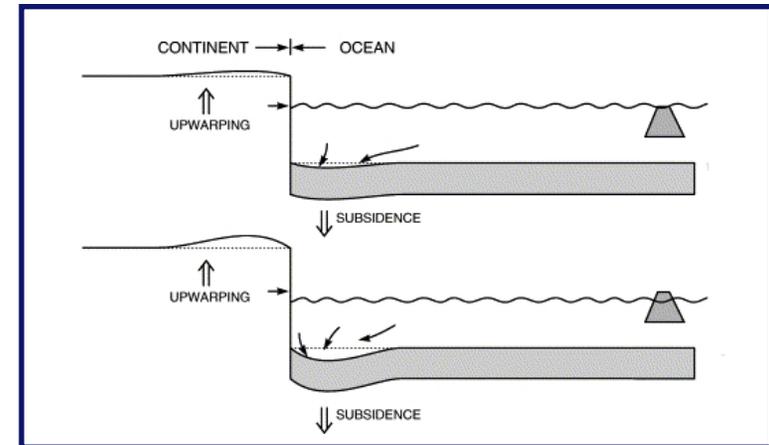


2b. Sea Level Physics: Ice-Age Timescales

Numerical Prediction of Present-Day Rate of Global Sea-Level Change Due to Ongoing GIA



Ocean Signal

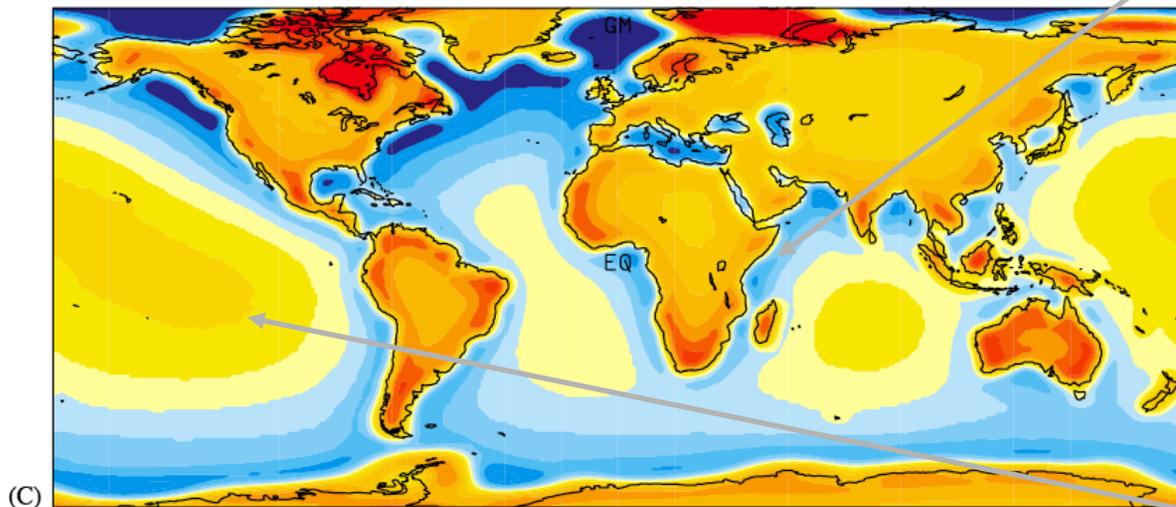


Continental Levering!

FAR FIELD

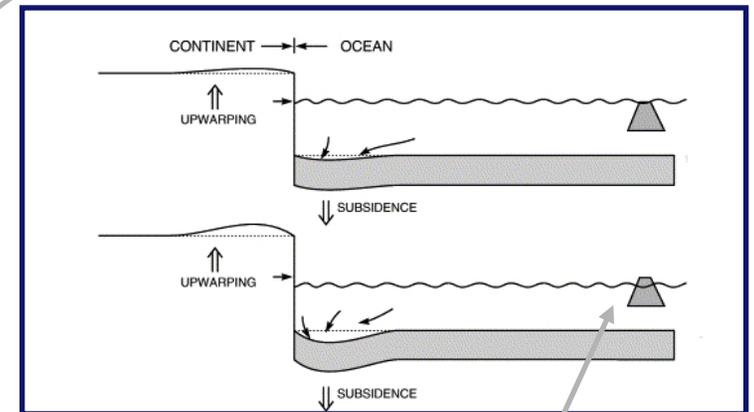
2b. Sea Level Physics: Ice-Age Timescales

Numerical Prediction of Present-Day Rate of Global Sea-Level Change Due to Ongoing GIA



Ocean Signal

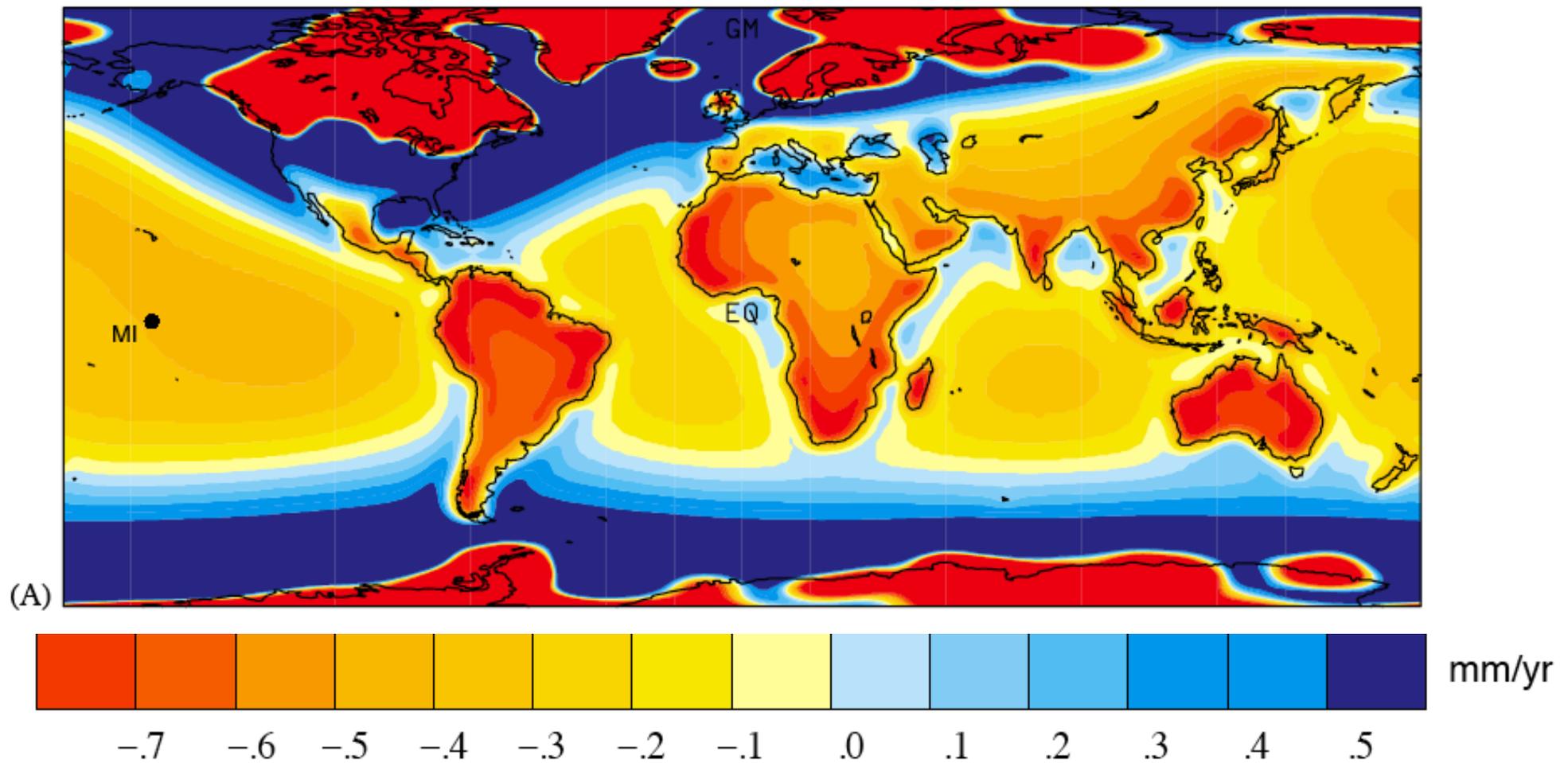
FAR FIELD



2nd contribution to equatorial ocean syphoning

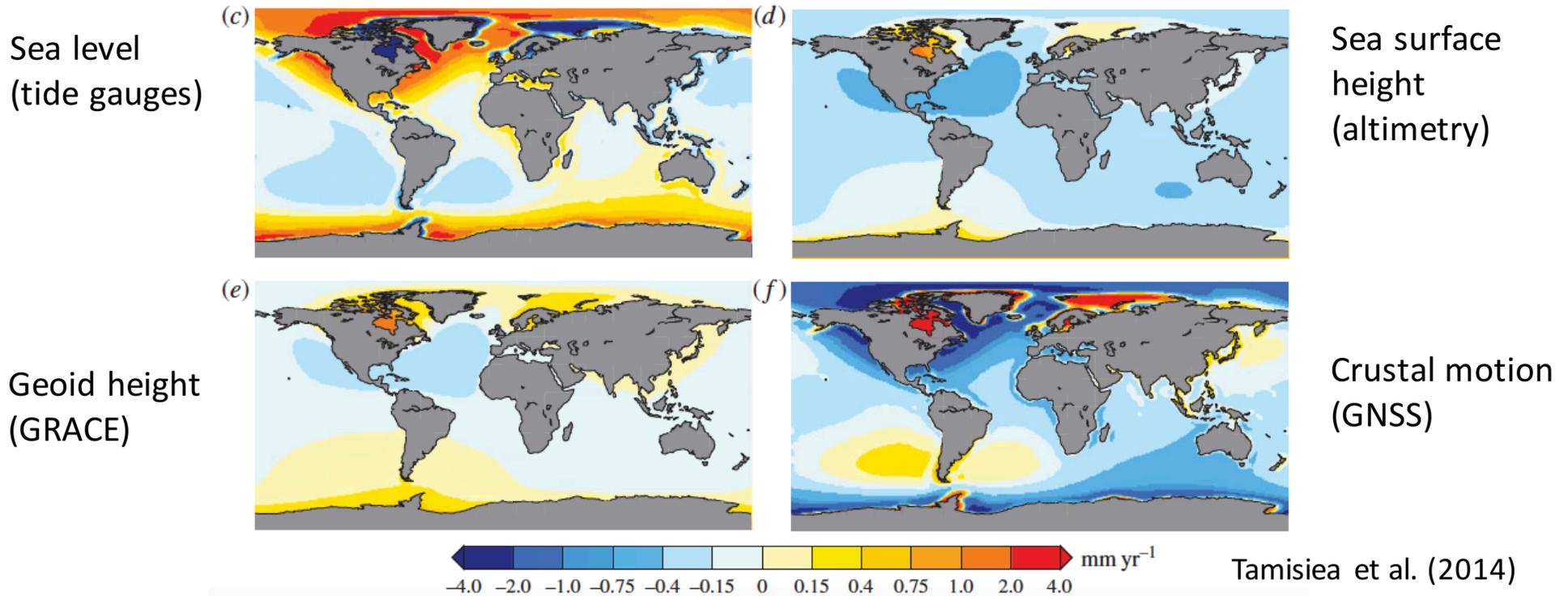
2b. Sea Level Physics: Ice-Age Timescales

Numerical Prediction of Present-Day Rate of Global Sea-Level Change Due to Ongoing GIA



2b. Sea Level Physics: Ice-Age Timescales

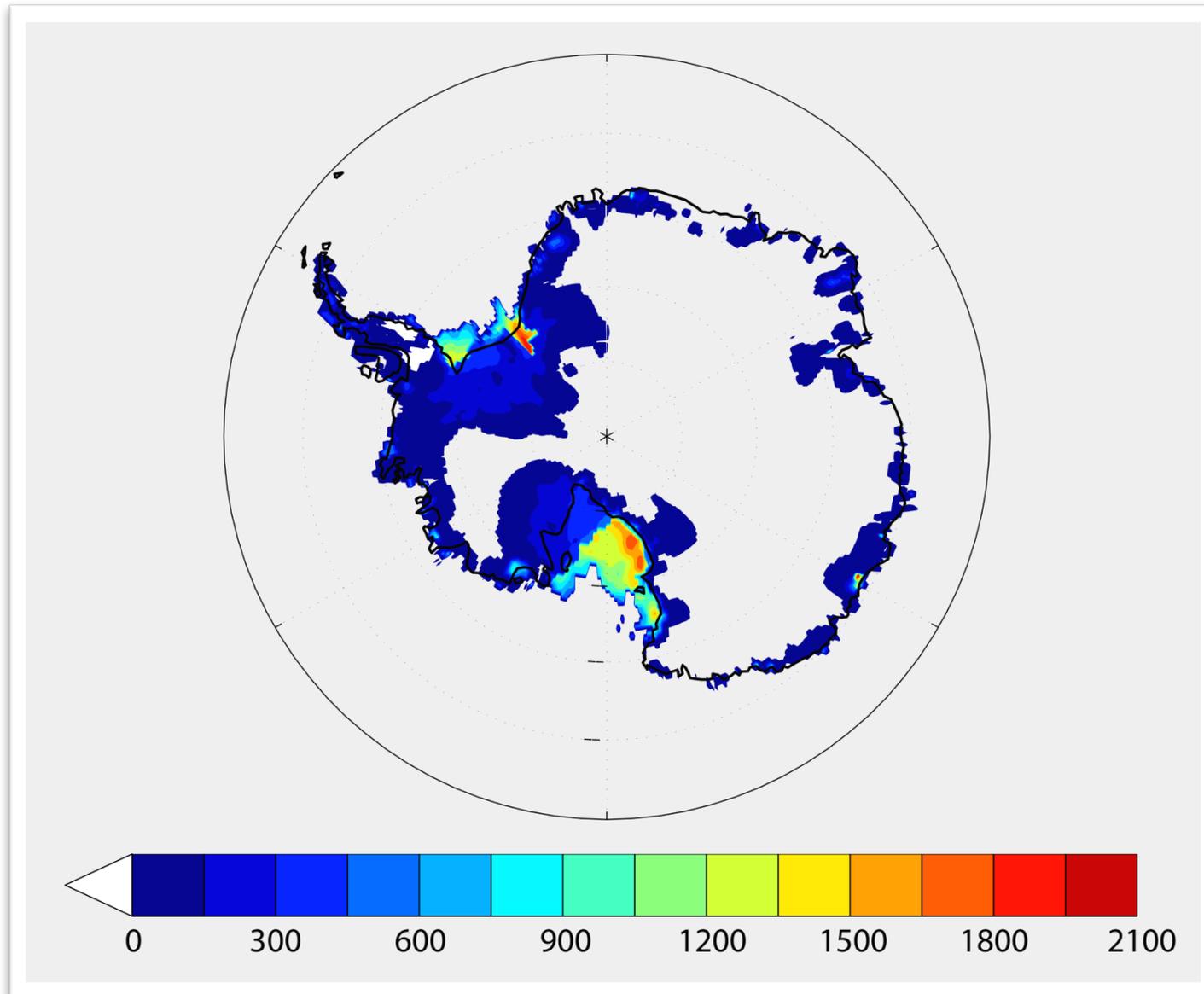
Expressions of GIA in modern sea-level records



All modern observations of sea-level-related quantities are impacted by past ice and ocean loading changes!

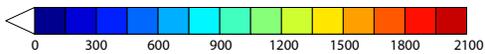
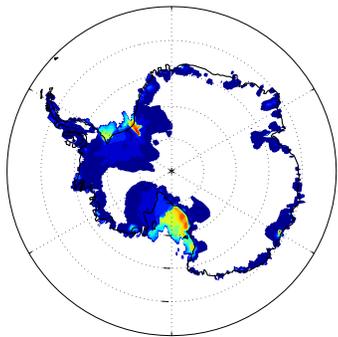
1b. Sea Level Change Example

Ice Loss Scenario:
eustatic value = 1.8 m (after filling the holes)

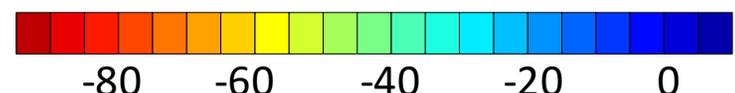
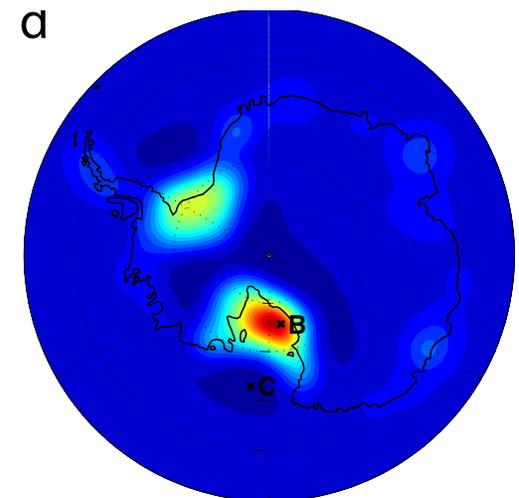
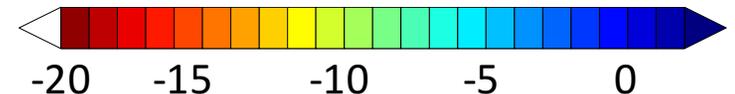
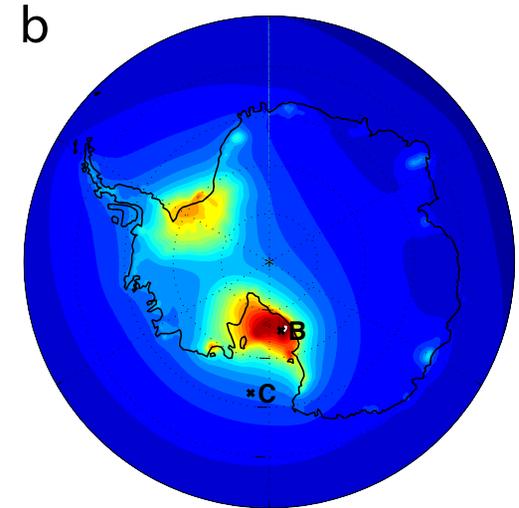
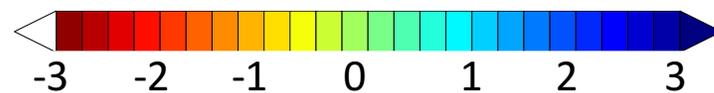
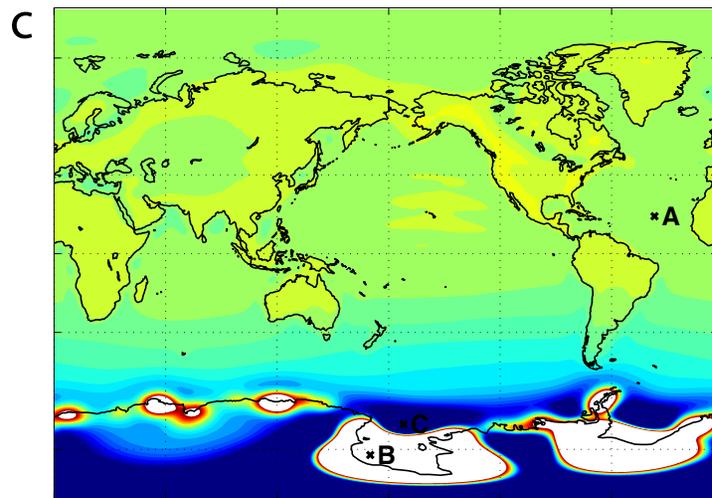
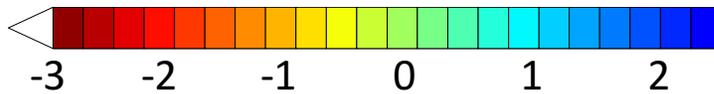
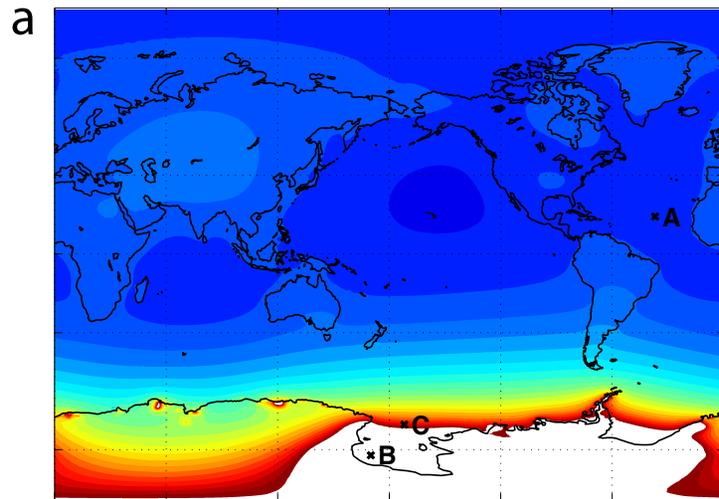


1b. Sea Level Change Example

Elastic SL change immediately after ice sheet retreat

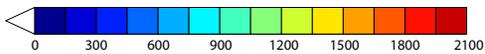
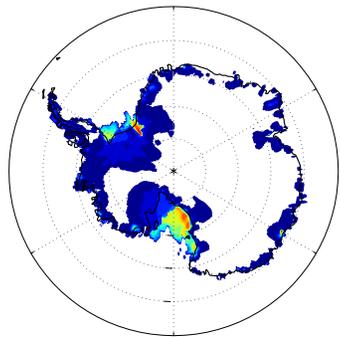
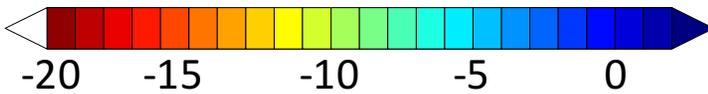
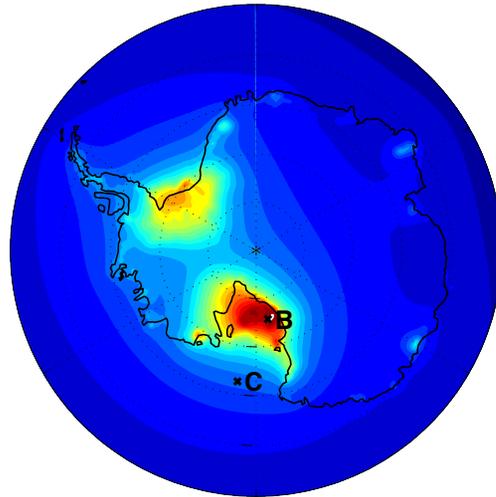


SL change over next 10 ky (ice remains constant)

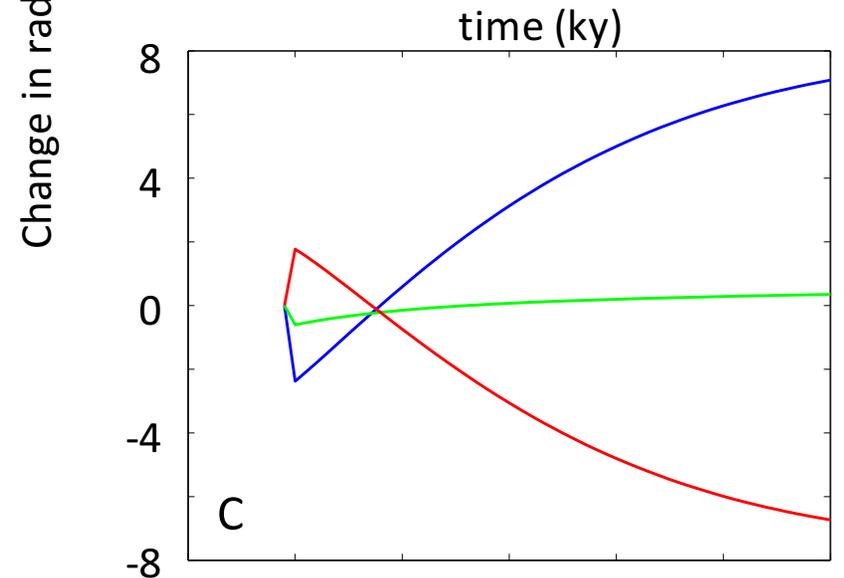
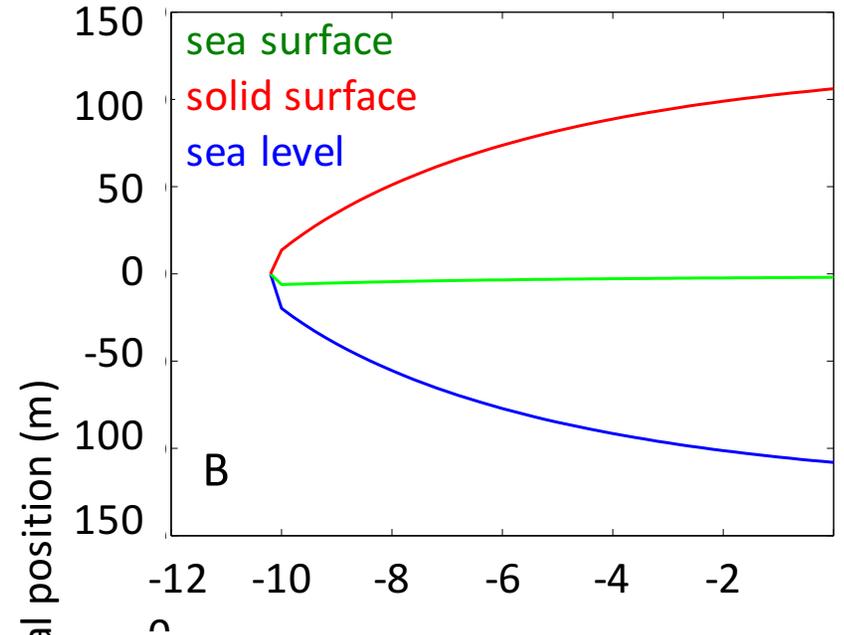
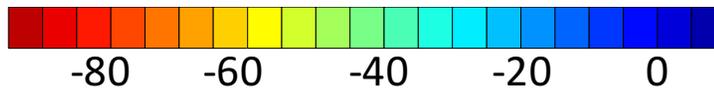
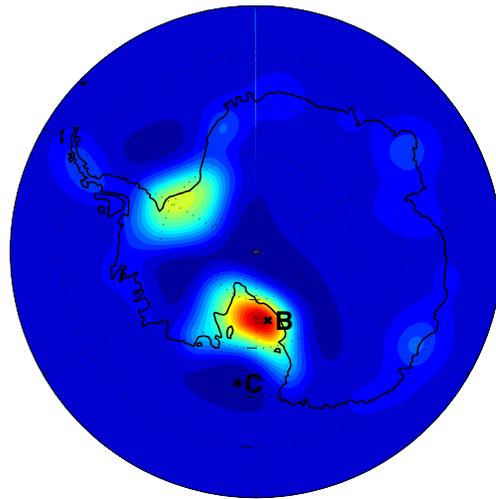


1b. Sea Level Change Example

Elastic SL change immediately after ice sheet retreat

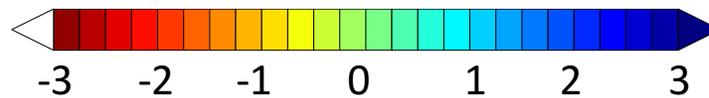
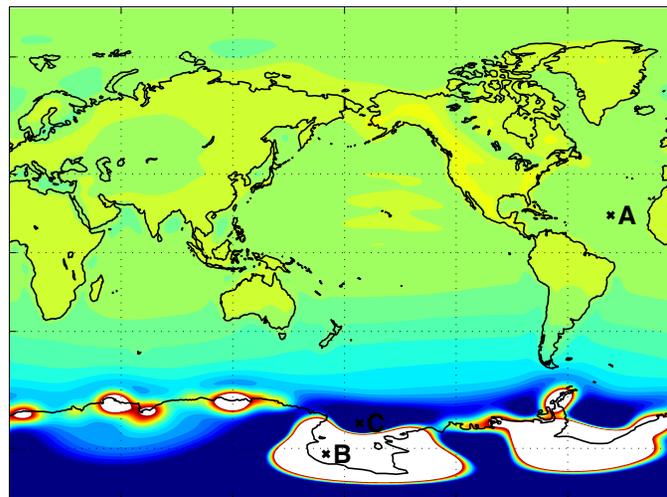
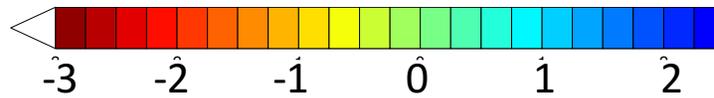
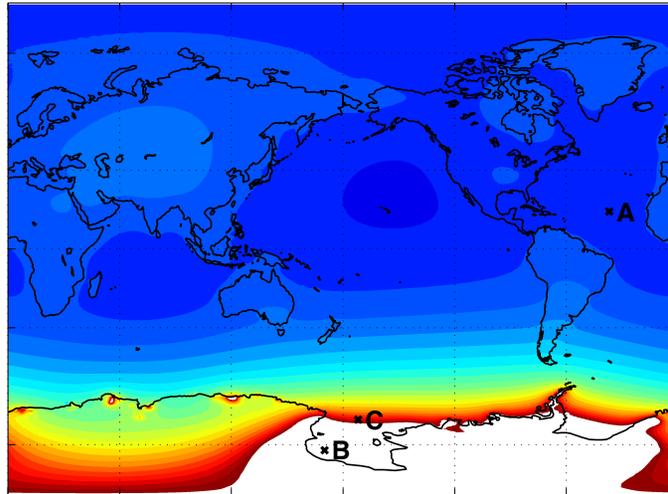


SL change over next 10 ky (ice remains constant)



1b. Sea Level Change Example

Elastic SL change immediately after ice sheet retreat



SL change over next 10 ky (ice remains constant)

