Amazing Ice: Glaciers and Ice Ages

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The Theory of Glaciation

Louis Agassiz, a Swiss geologist, observed glaciers.

- He saw glaciers as agents of landscape change.
 - They carried sand, mud, and huge boulders long distances.
 - They dropped these materials, unsorted, upon melting.

He realized that glaciers could explain erratic boulders.



The Theory of Glaciation

Agassiz proposed that an ice age had frozen Europe.

- Ice sheets covered land.
- Ice carried and dropped:
 - Fine-grained unsorted sediment.
 - Erratic boulders.



The Theory of Glaciation

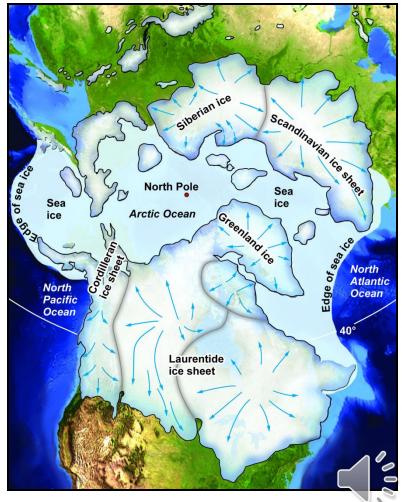
- When first proposed, Agassiz's idea was criticized.
- By the 1850s, many geologists agreed that he was right.
- Agassiz saw evidence for a North American ice age.



Glaciers and Ice Ages

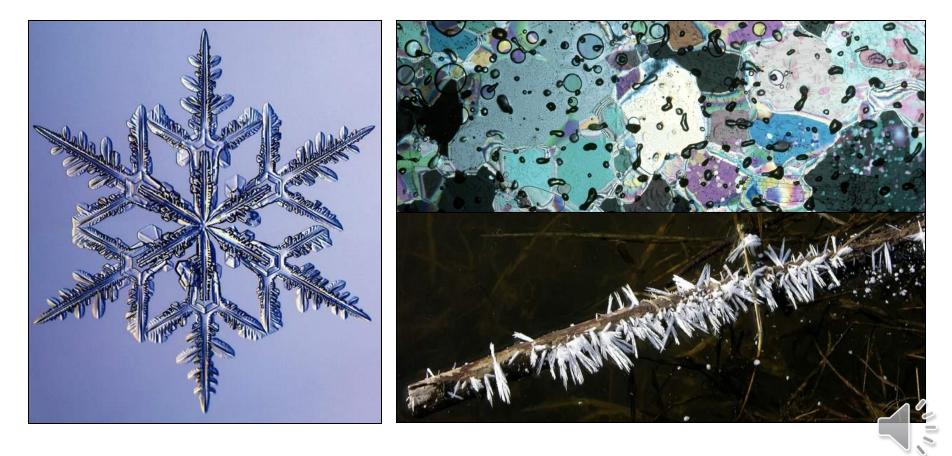
Glaciers

- Thick masses of recrystallized ice
 - Last all year long
 - Flow via gravity
 - Mountain and continental
- Presently cover ~10% of Earth
 - During ice ages, coverage expands to ~30%.
- The most recent ice age "ended" ~11 ka.
- Covered New York, Montreal, London, and Paris.
- Ice sheets were hundreds to thousands of meters thick.



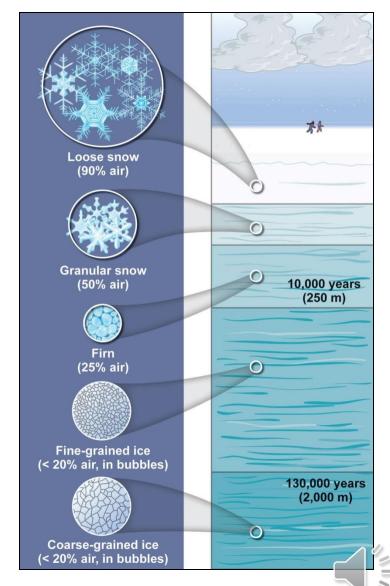
Ice: The Water Mineral

- Ice is solid water (H₂O).
- Forms when water cools below the freezing point.
- Natural ice is a mineral; it grows in hexagonal crystals.



How a Glacier Forms

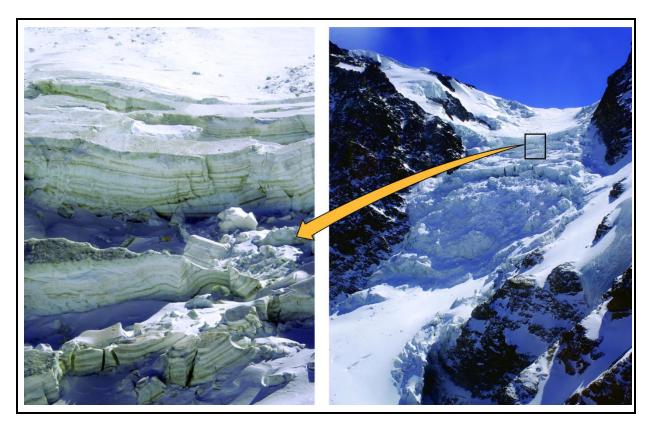
- Snowfall accumulates and survives the following summer.
- Snow is transformed into ice.
 - Snow is buried by later falls.
 - Compression reduces volume.
 - Burial pressure causes melting and recrystallization.
 - Snow turns into granular firn.
 - Over time, firn becomes interlocking crystals of ice.
- Glacial ice may form:
 - Quickly (tens of years)
 - Slowly (thousands of years)



Forming a Glacier

Conditions that are necessary to form a glacier:

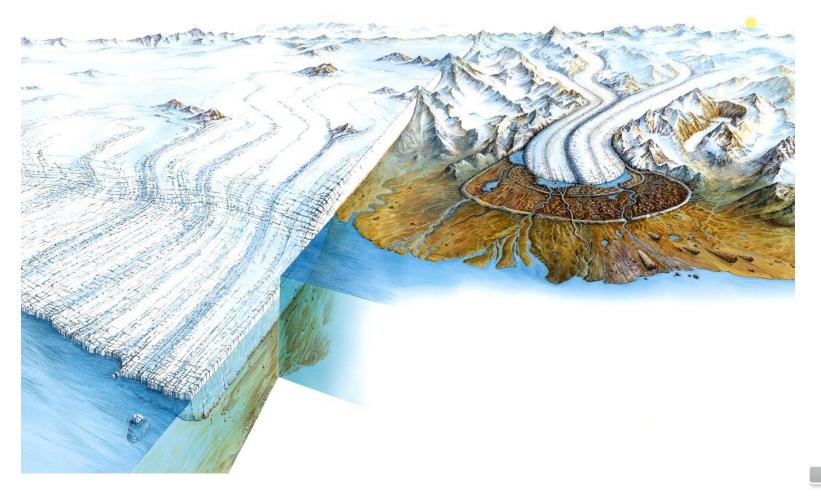
- Cold local climate (polar latitudes or high elevation).
- Snow must be abundant and accumulate in winter; more snow must fall than melts in the subsequent summer.





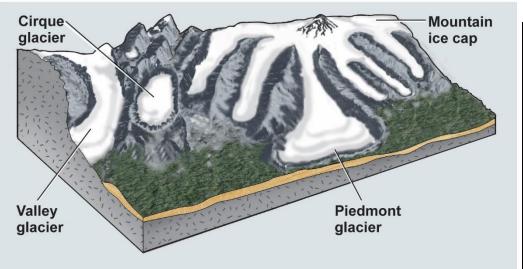
Glaciers

- Thick masses of recrystallized ice
- Two categories of glaciers: mountain and continental



Mountain Glaciers

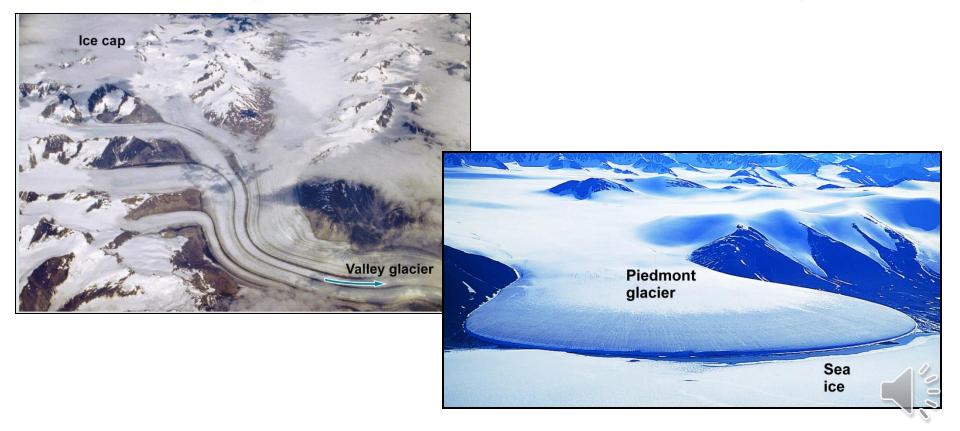
- Flow from high to low elevation in mountain settings.
- Include a variety of types:
 - Cirque glaciers fill mountain-top bowls.
 - Valley glaciers flow like rivers down valleys.
 - Mountain ice caps cover peaks and ridges.
 - Piedmont glaciers spread out at the end of a valley.





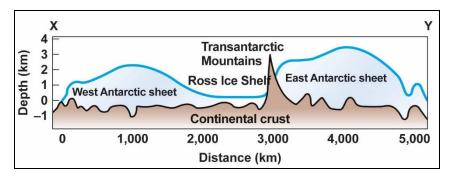
Mountain Glaciers

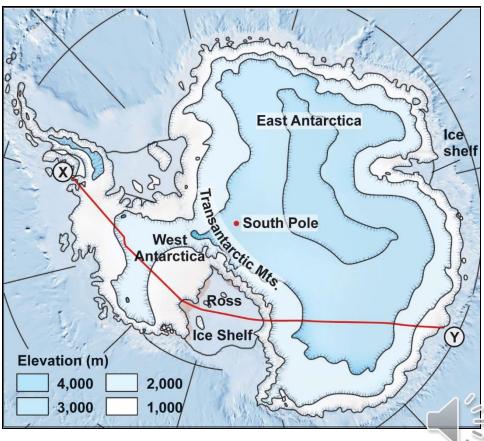
- Flow from high to low elevation in mountain settings.
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Continental Glaciers – Ice Sheets

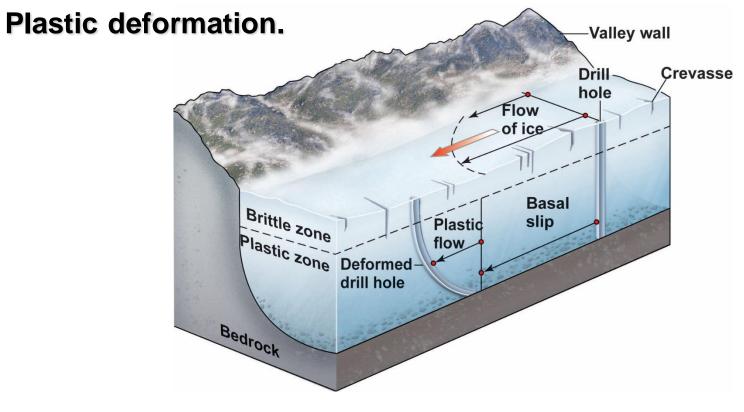
- Vast ice sheets covering large land areas.
- Ice flows outward from thickest part of sheet.
- Two major ice sheets remain on Earth:
 - Greenland
 - Antarctica





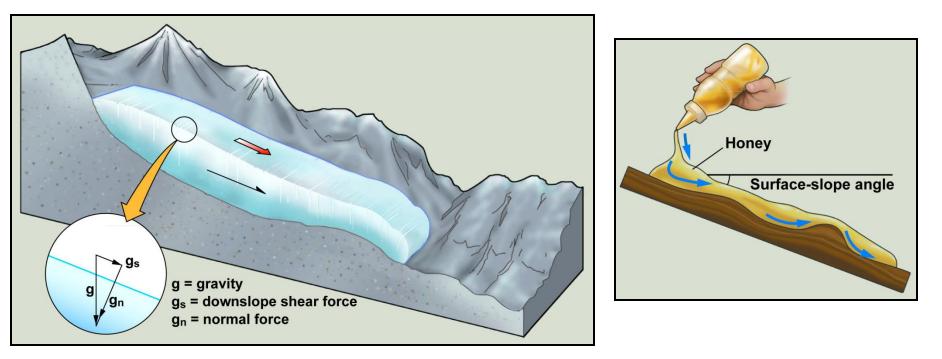
Movement of Glacial Ice

- How do glaciers move?
 - Basal sliding.
 - Significant quantities of meltwater forms at base of glacier.
 - Water decreases friction, ice slides along substrate.



Movement of Glacial Ice

- Why do glaciers move?
 - The pull of gravity is strong enough to make ice flow.
 - A glacier moves in the direction of its surface slope.
 - > The ice base can flow up a local incline.



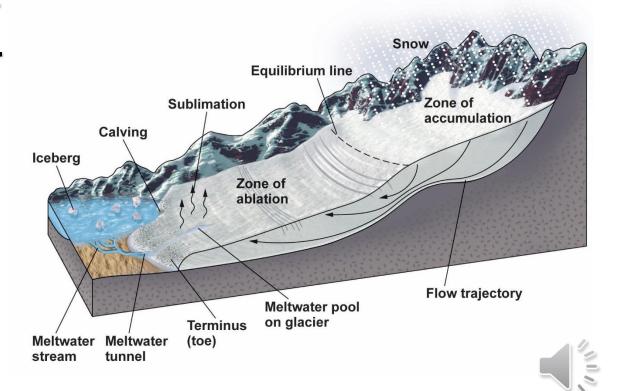


Glacial Advance and Retreat

- Glaciers behave like bank accounts for water.
- Zone of accumulation—area of net snow addition.
 - Colder temperatures prevent melting.
 - Snow remains across the summer months.
- Zone of ablation—

area of net ice loss.

Zones meet at the equilibrium line.



Glacial Advance and Retreat

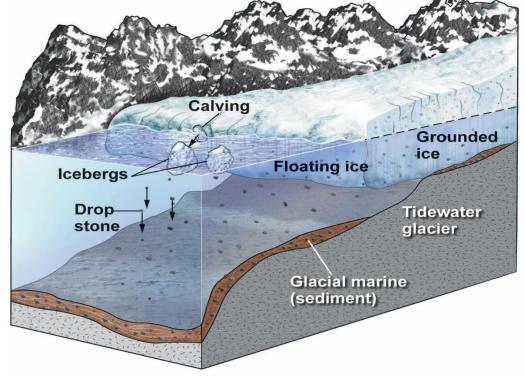
Toe—the leading edge of a glacier
 Ice always flows downhill, even during toe retreat.



Ice in the Sea

In polar regions, glaciers flow out over ocean water.

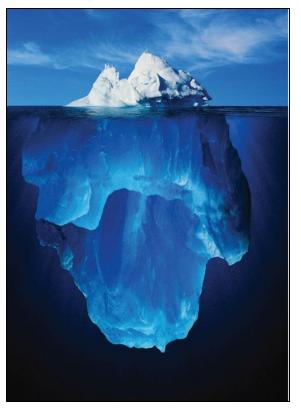
- Tidewater glaciers—valley glaciers entering the sea
- Ice shelves—continental glaciers entering the sea
- Sea ice (Arctic Ocean) nonglacial ice formed of frozen seawater

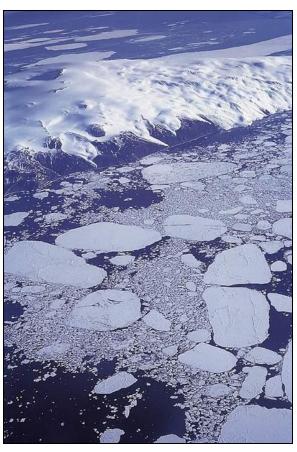




Ice in the Sea

- Floating ice is normally four fifths beneath the waterline.
- Floating ice exhibits a variety of shapes and sizes.
 - Iceberg—greater than 6 m above water
 - Ice shelves yield tabular bergs.

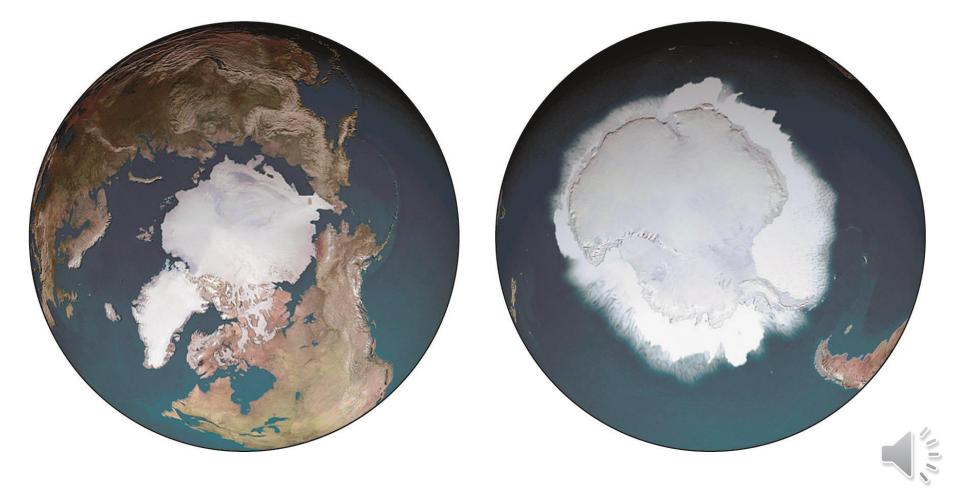






Ice in the Sea

- Large areas of the polar seas are covered with ice.
- Global warming is causing a reduction in <u>sea ice cover</u>.



Carving and Carrying by Ice

- Glaciers are important forces of landscape change.
 - Erosion
 - Transport
 - Deposition

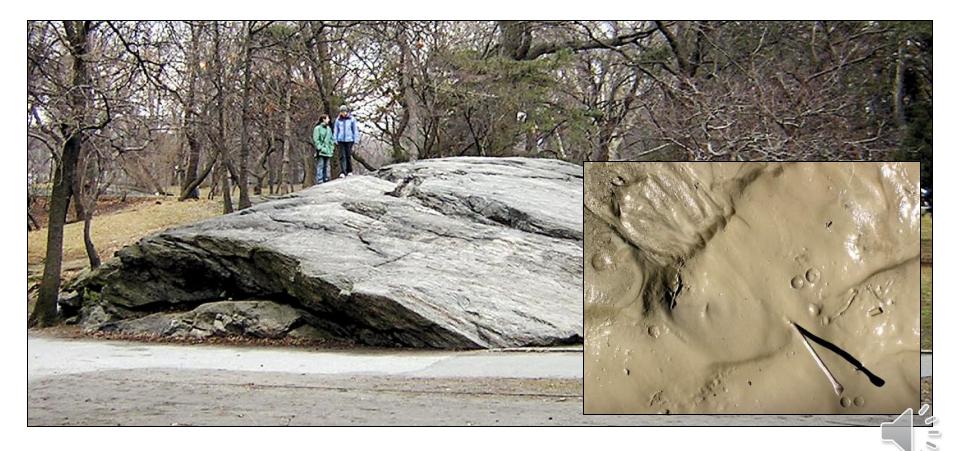


- Glaciers carve deep valleys, such as Yosemite Valley.
 - Polished granite domes and vertical cliffs are the result of glacial erosion.



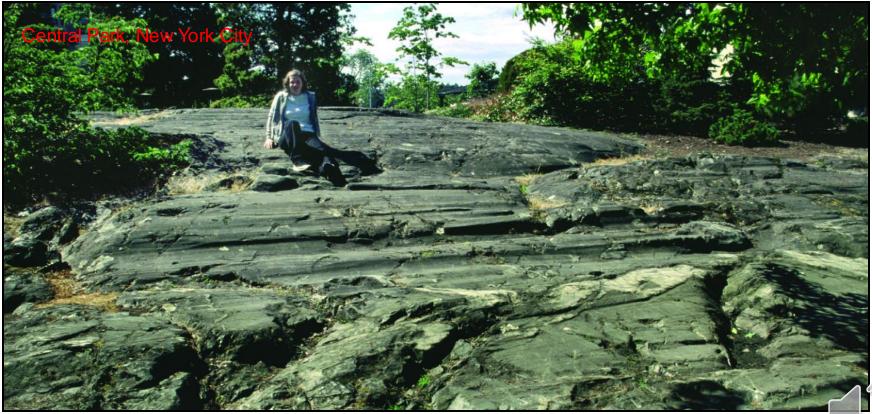
Glacial abrasion—a "sandpaper" effect on substrate

- Substrate is pulverized to fine "rock flour."
- Sand in moving ice abrades and polishes bedrock.



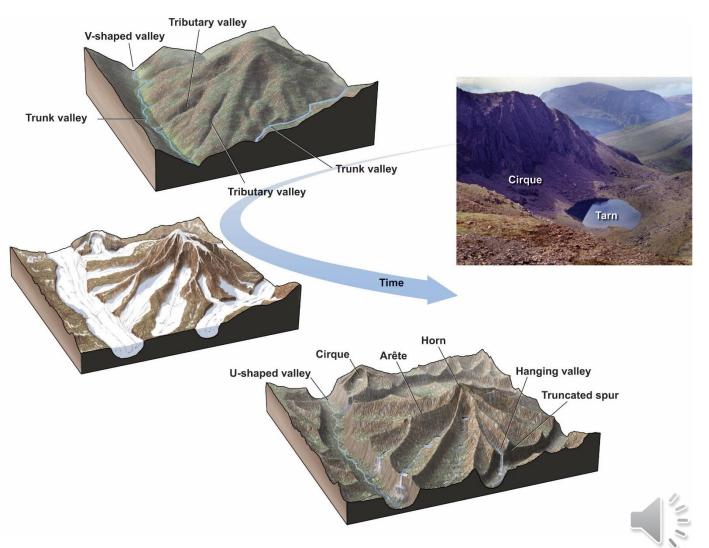
Glacial abrasion—a "sandpaper" effect on substrate

- Large rocks are dragged across bedrock gouge striations.
- Striations run parallel to direction of ice movement.



Erosional features of glaciated valleys:

- Cirques
- Tarns
- Aretes
- Horns
- U-shaped valleys
- Hanging valleys
- Fjords



Cirques—bowl-shaped basins high on a mountain

- Form at the uppermost portion of a glacial valley.
- Freeze-thaw mass wasting chews into the cirque headwall.
- After ice melts, the cirque often becomes a tarn (lake).



Arete—a "knife-edge" ridge

Formed by two cirques that have eroded toward one another

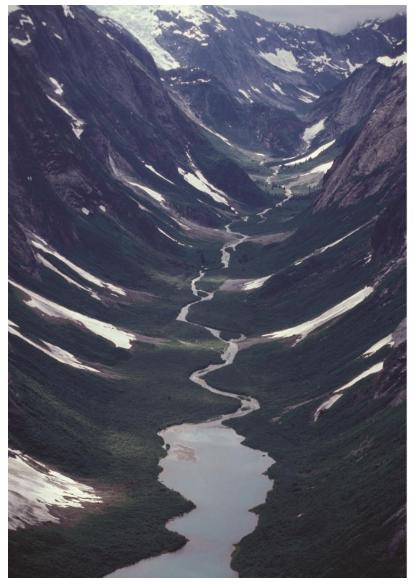


- Horn—a pointed mountain peak
 - Formed by three or more cirques that surround the peak



U-shaped valleys

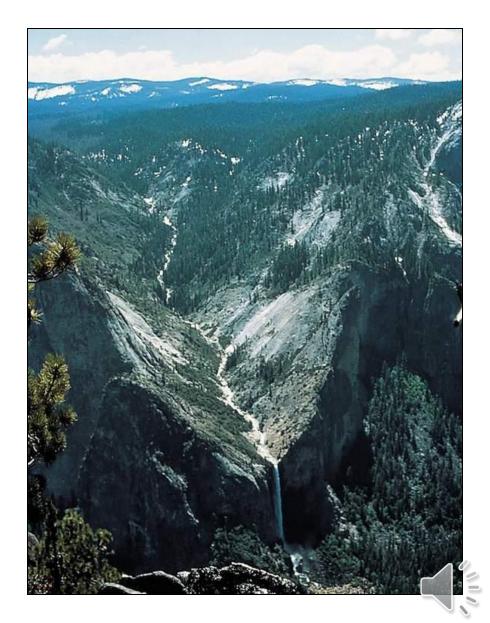
- Glacial erosion creates a distinctive trough.
- Compare to V-shaped fluvial valleys.





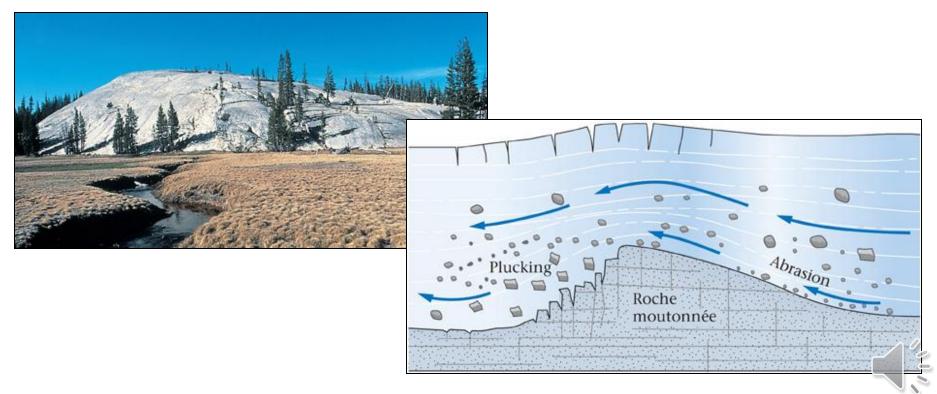
Hanging valleys

- The intersection of a tributary glacier with a trunk glacier
- Trunk glacier incises deeper into bedrock.
- Troughs have different elevations.
- A waterfall results.



Glaciers can also erode by plucking.

- Ice freezes around bedrock fragments and plucks chunks as glacier advances.
- It forms a distinctive asymmetric hill called a roche moutonée.



- U-shaped glacial troughs flooded by the sea
- Accentuated by isostatic rebound





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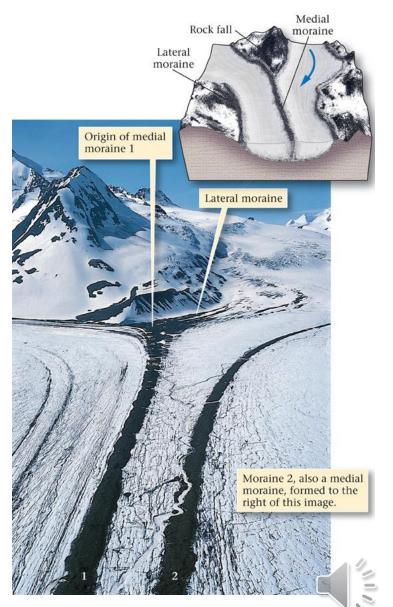


- U-shaped glacial troughs flooded by the sea
- Accentuated by isostatic rebound



Deposition Associated with Glaciation

- Moraines—unsorted debris deposited by a glacier
 - Lateral—forms along the flank of a valley glacier
 - Medial—mid-ice moraine from merging of lateral moraines



Types of Glacial Sedimentary Deposits

- Many types of sediment derive from glaciation.
- Called glacial drift, these include:
 - Glacial till
 - Erratics
 - Glacial marine sediments
 - Glacial outwash
 - Loess (aeolian)
 - Glacial lake-bed sediment
- Stratified drift is water sorted; unstratified drift is not sorted.



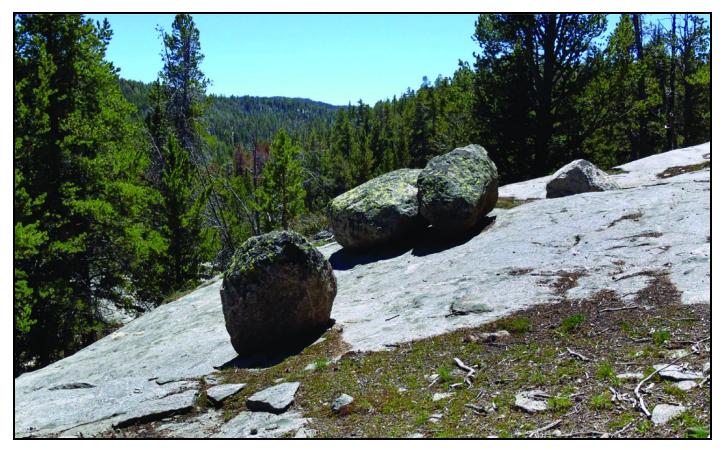
Glacial till—sediment dropped by glacial ice

- Consists of all grain sizes—boulders to clay.
- Unmodified by water, hence:
 - Unsorted
 - Unstratified
- Accumulates:
 - Beneath glacial ice
 - At the toe of a glacier
 - Along glacial flanks



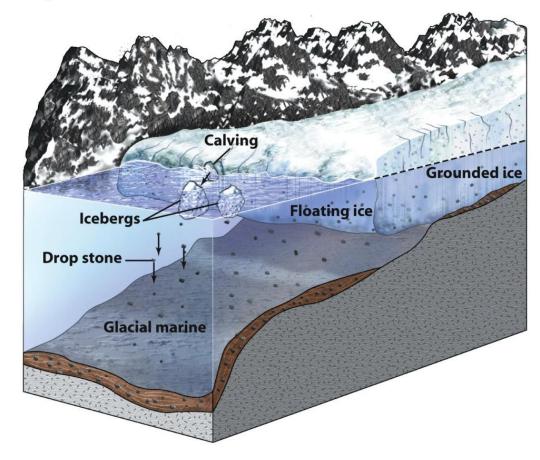
Erratics—boulders dropped by glacial ice.

- These rocks are different from the underlying bedrock.
- Often, they have been carried long distances in ice.



Glacial marine—sediments from an oceanic glacier

- Calving icebergs raft sediments away from the ice.
- Melting icebergs drop stones into bottom mud.





Glacial outwash—sediment transported by meltwater

- Muds are removed.
- Sizes are graded and stratified.
- Grains are abraded and rounded.
- Outwash is dominated by sand and gravel.





Glacial lake-bed sediment

- Lakes are abundant in glaciated landscapes.
- Fine rock flour settles out of suspension in deep lakes.
- Muds display seasonal varve couplets.
 - Finest silt and clay are from frozen winter months.
 - Coarser silt and sand are from summer months.





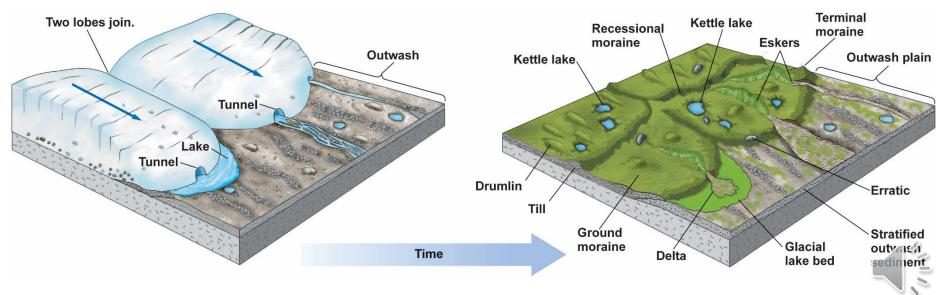
Loess—wind-transported silt. Pronounced "luss."

- Glaciers produce abundant amounts of fine sediment.
- Strong winds over ice blow the rock flour away.
- This sediment settles out near glaciated areas as loess deposits.
- Deposits are unstratified and distinct in color.

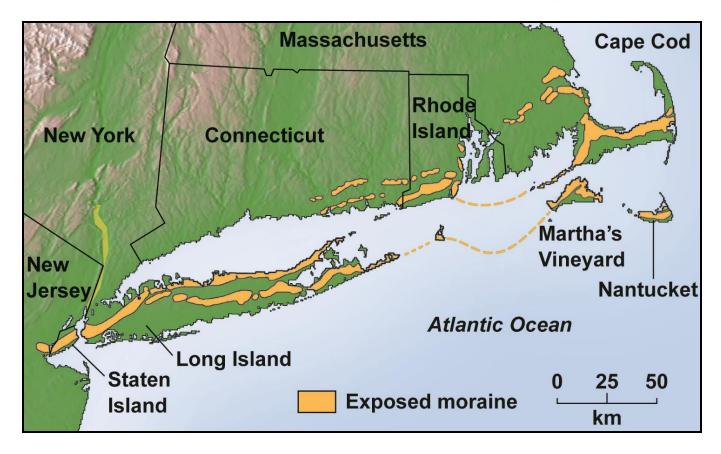


Glacial Depositional Landforms

- Glacial sediments create distinctive landforms:
 - End moraines and terminal moraines
 - Recessional moraines
 - Ground moraine
 - Drumlins
 - Kettle lakes
 - Eskers



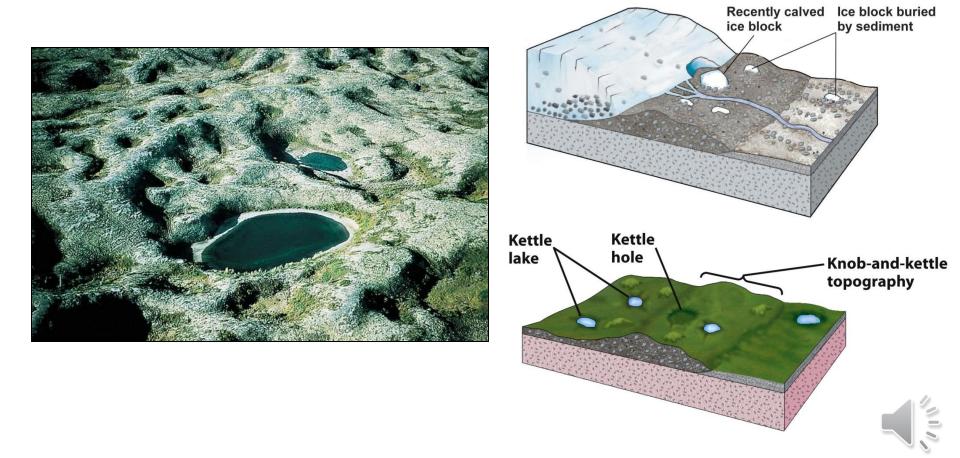
- End moraines form at the stable toe of a glacier.
- Terminal moraines form at the farthest edge of flow.
- Recessional moraines form as retreating ice stalls.





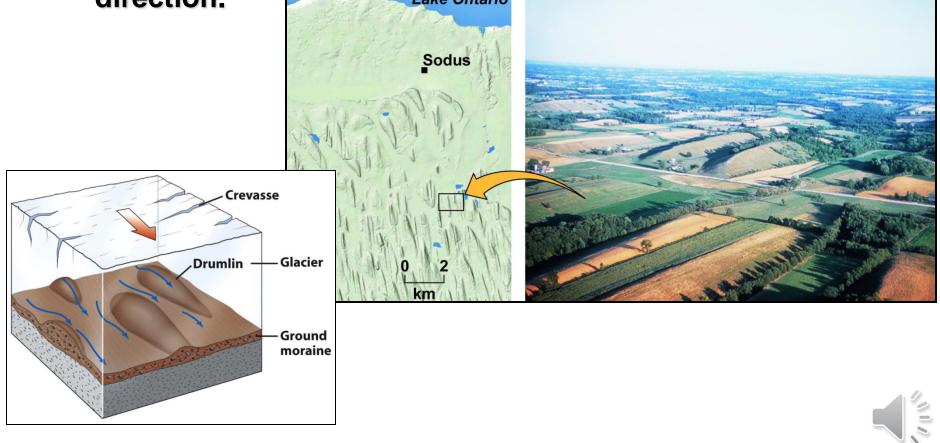
Ground moraine is till left behind by rapid ice retreat.

- Creates a hummocky, mostly flat land surface.
- Kettle lakes form from stranded ice blocks.

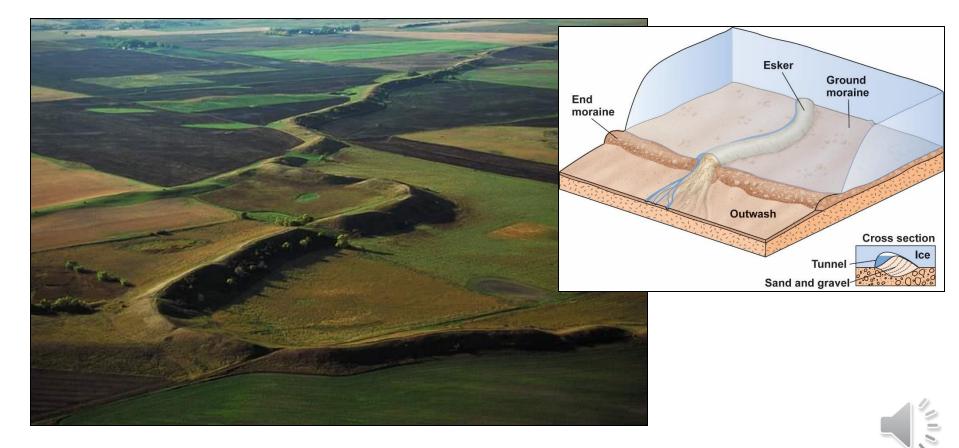


Drumlins—long, aligned hills of molded lodgment till

- Asymmetric form—steep up-ice; tapered down-ice.
- Commonly occur as swarms aligned parallel to ice-flow direction.

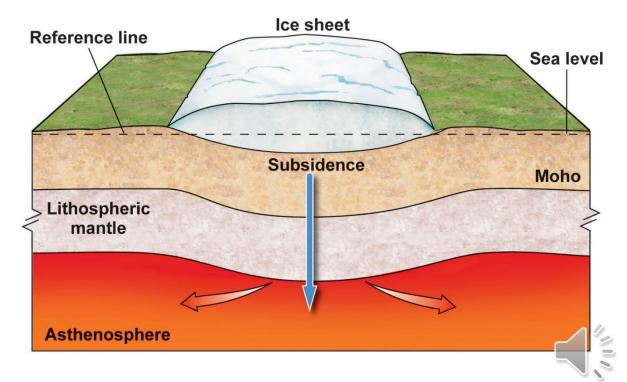


- Eskers are long, sinuous ridges of sand and gravel.
- They form as meltwater channels within or below ice.
- Channel sediment is released when the ice melts.



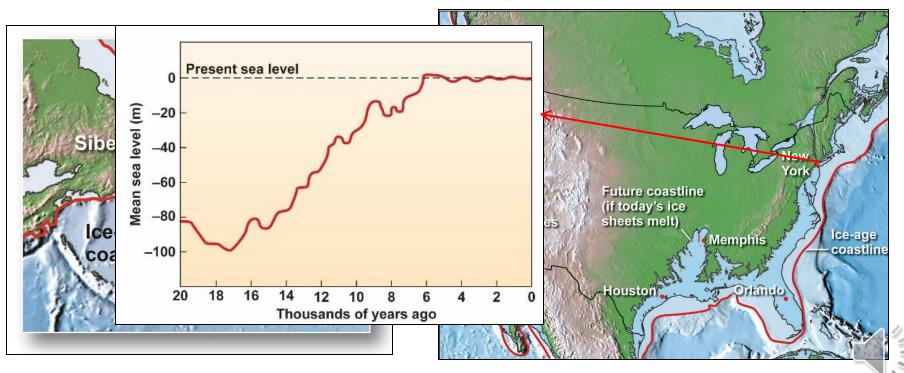
Consequences of Continental Glaciation

- Ice loading and glacial rebound
 - Ice sheets depress the lithosphere into the mantle.
 - Slow crustal subsidence follows flow of asthenosphere.
 - After ice melts, the depressed lithosphere rebounds.
 - The last ice-age glacial rebound continues slowly today.



Sea level—ice ages cause sea level to rise and fall.

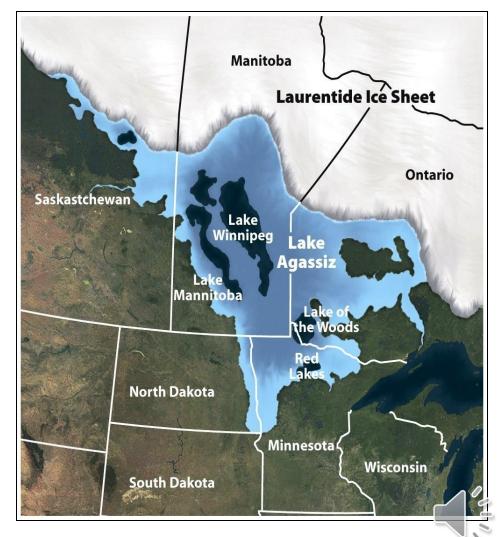
- Water is stored on land during an ice age; sea level falls.
- Deglaciation returns water to the oceans; sea level rises.
- Sea level was ~100 m lower during the last ice age.
- If ice sheets melted, coastal regions would be flooded.



Gigantic proglacial lakes formed near the ice margin.

Glacial Lake Agassiz

- Covered a huge area.
- Existed for 2,700 years.
- Drained abruptly.
- Exposed mud-rich, extremely flat land.



Pluvial features—large lakes formed during ice age.

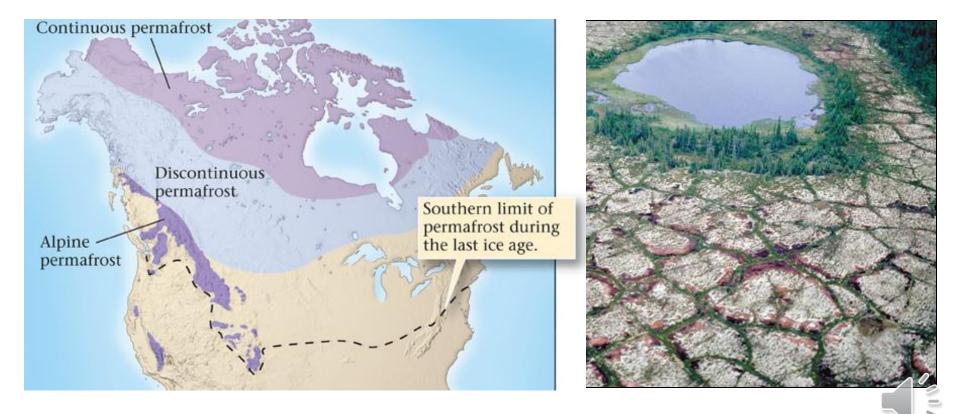
- The American Southwest was much wetter.
 - Large lakes occupied today's deserts.
 - Lake Bonneville (remnant is Great Salt Lake).





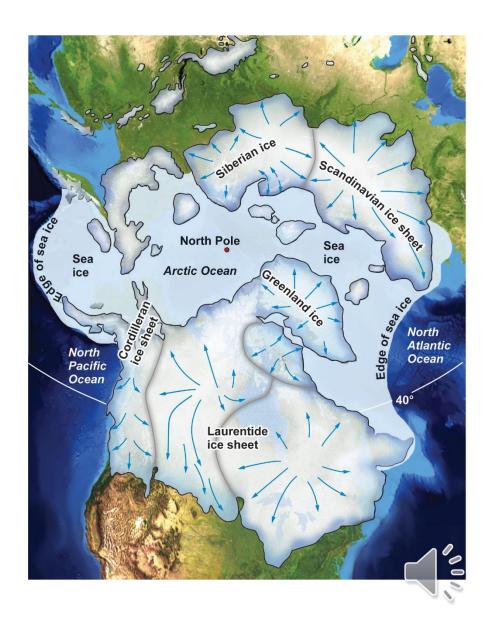
Periglacial (near-ice) environments are unique.

- Characterized by year-round frozen ground (permafrost).
- Freeze-thaw cycles generate unusual patterned ground.



Pleistocene Ice Ages

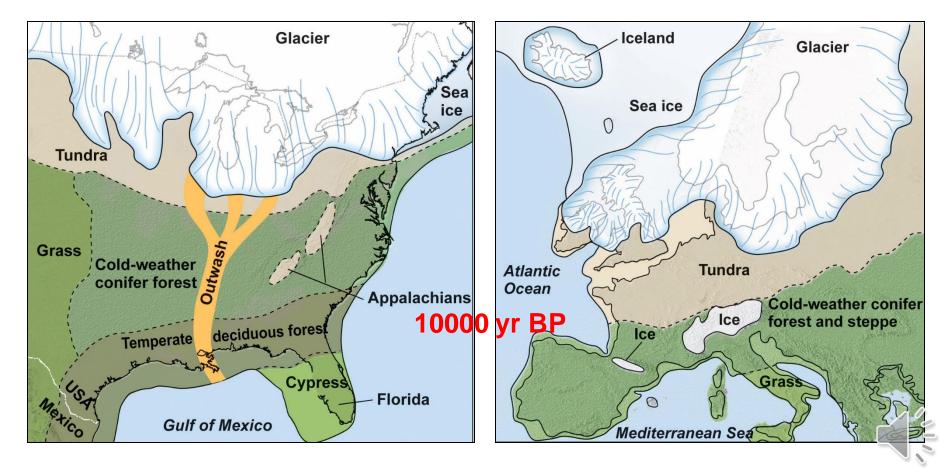
- Young (<2.6 Ma) glacial remnants are abundant.</p>
 - Northern North America
 - Scandinavia and Europe
 - Siberia
- Landscapes in these regions are clearly glacial.



Pleistocene Life and Climate

All climate and vegetation belts were shifted southward.

- The tundra limit was ~48° N. Today, it is above 68° N.
- Vegetation evidence is preserved as pollen found in bogs.



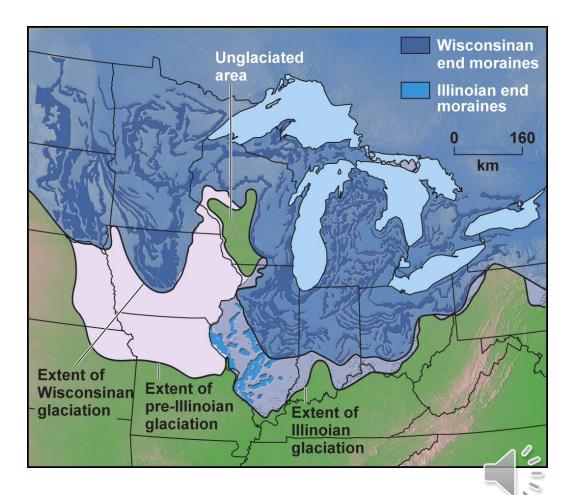
Pleistocene Life and Climate

- Pleistocene fauna were well adapted.
- Mammals included now-extinct giants:
 - Giant beaver
 - Giant ground sloth
 - Mammoths and mastodons
- Modern humans proliferated.



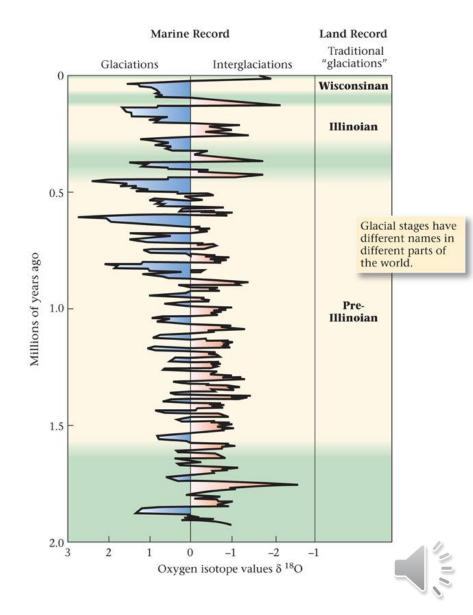
Timing of the Pleistocene Ice Age

- In North America, multiple Pleistocene glacial advances are recognized. Youngest to oldest:
 - Wisconsinan
 - Illinoian
 - Pre-Illinoian
- Ice ages separated by interglacials intervals



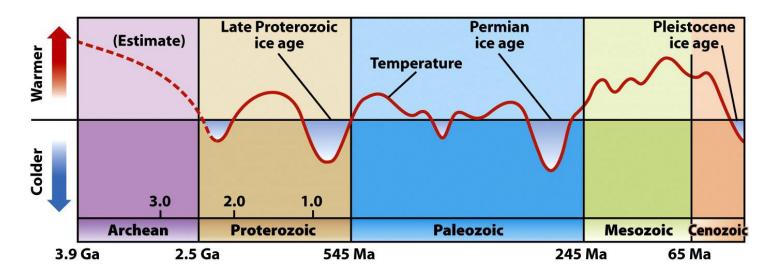
Timing of the Pleistocene Ice Age

- Oxygen isotopes suggest twenty or more glaciations throughout Earth history.
 - Higher ¹⁸O/¹⁶O = colder.
 - Lower ¹⁸O/¹⁶O = warmer.
- The "original four" ice ages may simply have been the largest.



Earlier Glaciations

- Glaciations have occurred before in Earth history.
- They are indicated by fossil tills and striated bedrock:
 - Pleistocene (since 2.5Ma ago)
 - Permian
 - Ordovician
 - Late Proterozoic—tillites at equatorial latitudes suggest an ice-covered world ("Snowball Earth").

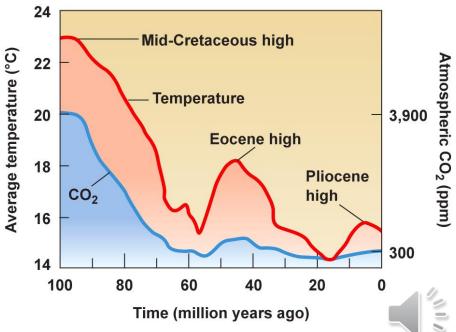




Causes of Glaciation

Long-term causes—set the stage for ice ages.

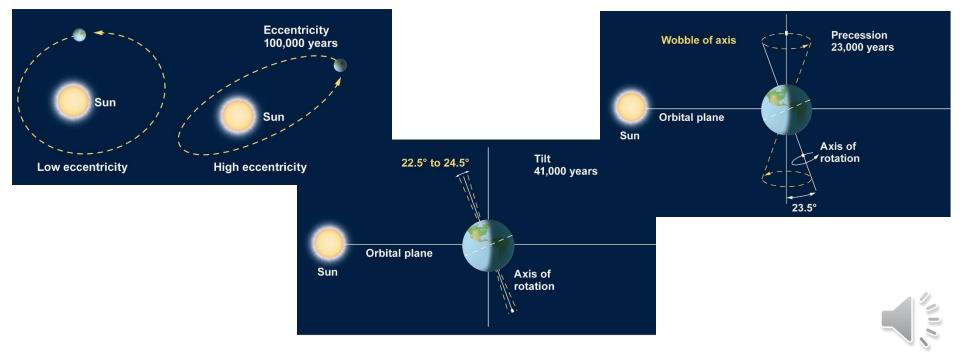
- Plate tectonics control factors that influence glaciation.
 - Distribution of continents toward high latitudes
 - Sea-level flux by mid-ocean-ridge volume changes
 - Oceanic currents
- Atmospheric chemistry
 - Changes in greenhouse gas concentrations
 - Carbon dioxide (CO₂)
 - Methane (CH₄)
- Global average temperature today: ~14C



Causes of Glaciation

Short-term causes—govern advances and retreats

- Milankovitch hypothesis—climate variation over 100 to 300 ka predicted by cyclic changes in orbital geometry.
 - > The shape of Earth's orbit varies (~100,000 year cyclicity).
 - ▶ Tilt of Earth's axis varies from 22.5° to 24.5° (~41,000 years).
 - Precession—Earth's axis wobbles like a top (23,000 years).

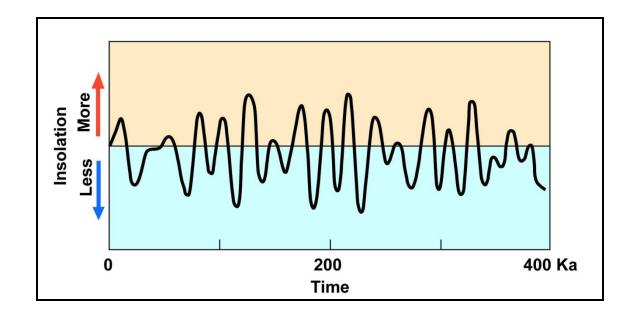


Causes of Glaciation

Short-term causes – govern advances and retreats

Milankovitch cycles drive global climate and glacial cycles.

- Stage 1: average temperature drops, glaciers are born.
- Stage 2: glaciers grow, albedo causes further cooling.
- Stage 3: temperatures warm, glaciers shrink, interglacial begins.

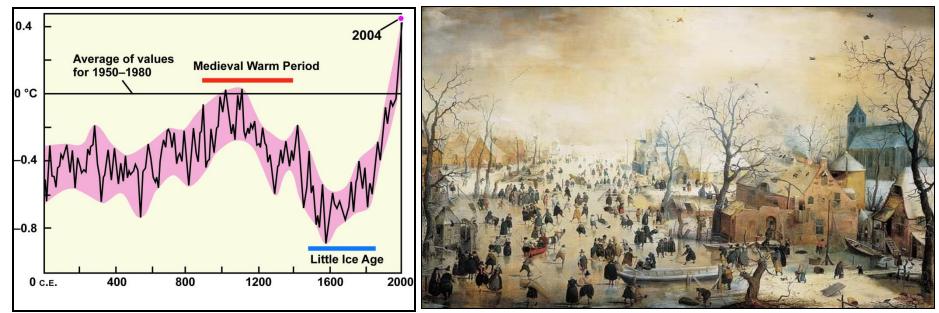




Will There Be Another Glaciation?

We are living in an interglacial. Ice will return! but when?

- Recent interglacials have lasted ~10,000 years.
- But, it has been ~11,000 years since the last deglaciation.
- A cool period (1300–1850) resulted in the Little Ice Age.
- We may have forestalled the next glaciation.... How?





Will There Be Another Glaciation?

- During the last 150 years, temperatures have risen and most mountain glaciers have dramatically retreated.
- Earth's climate could now be in "super-interglacial" period.
- This current interglacial might be in extension because of human-induced warming!

